

- Qy** Youngest alluvium (upper Holocene) Arkosic gravel, sand, and silt along present streams. Derived from granitic and metamorphic rocks in Ruby Mountains. Thickness a few meters.
- Oya** Younger alluvium and outwash (Holocene and Pleistocene) Alluvial fan deposits of arkosic gravel and sand derived from granitic and metamorphic rocks in Ruby Mountains. Includes Pleistocene outwash from Angel Lake glacial advance. Thickness more than 5 m.
- Ol** Landslide deposits (Holocene and Pleistocene) Locally derived. Includes lobate rock avalanche deposit of rock in upper Conrad Creek drainage. Deposit near Witherington Creek in northeast part of map area possibly includes glacial till. Confined to Ruby Mountains. Maximum thickness about 20 m.
- Qc** Colluvium (Holocene and Pleistocene) Largely slope wash, talus deposits, landslide, debris-flow, and possible suffocation deposits. Locally derived. Confined to Ruby Mountains. Thickness as much as ~10 m.
- Qga** Angel Lake moraines Deposited by the Angel Lake glaciation (Angel Lake substage or stage). Thickness as much as 50 m.
- Qg** Lamolite moraines Deposited by the Lamolite glaciation (Lamolite substage or stage). Mapped near Hidden Lakes by Sharp (1938); inferred here also along Smith Creek, based on degraded form. Thickness up to perhaps 20 m.
- Pluvial lake deposits (Pleistocene)** Deposits of Pleistocene Lake Franklin in Ruby Valley. Pluvial lakes discussed by Hubbs and others (1974), Milfin and Wheat (1979), and Rehels and others (2002).
- Qps** Silt and clay Locally includes sand and pebbles. At least a few meters thick.
- Qpb** Beach and delta deposits Well sorted to very well sorted arkosic pebble gravel and sand. Grains subangular to subrounded; locally imbricated. Cattle-cemented by groundwater below a depth of about 2 m. Surface hummocky in southeast corner of quadrangle, where fine sand occupies surface of swales between gravelly hummocks. Thickness >4 m.
- Qo** Old alluvium (Pleistocene) Underlies dissected terraces 50 to 100 m above present stream grade. Alluvial-fan deposits of arkosic gravel and sand made of subrounded to subangular clasts of granitic and metamorphic rocks derived from the Ruby Mountains. Mostly poorly consolidated. Thickness 0-50 m. Verses two levels of perched bedrock pediment locally on east side of Ruby Mountains, where higher (eastward?) deposit on terrace surface in Joe Billy Basin along Smith Creek perches higher than and may predate deposit that mantles eastern range-front pediment terrace, itself uplifted by Quaternary fault.
- Tb** Basalt dikes (Miocene) Olivine basalt and diabase. Aphanitic groundmass. Dikes commonly parallel or follow large fractures. Most dikes strike north and dip steeply. Dike thicknesses 0.1 to 5 m. Present west of Verdi Peak, west of Old Man of the Mountains, and 3 km south of Old Man of the Mountains. Similar dikes dated as approximately 13.5 Ma (13.5±0.3 Ma; non-published) in central Ruby Mountains (whole-rock ⁴⁰Ar/³⁹Ar, Hudec, 1990) and about 17 Ma in southern East Humboldt Range (whole rock K-Ar, Snoke and others, 1979).
- Tp** Pegmatite dikes (Oligocene?) Dikes of leucogranite composition containing coarse to very coarse blocks of muscovite and biotite; sparse garnet. Graphic texture common. Plumes or cortical muscovite-quartz intergrowths like those described by Orville (1960) common; locally the cores open upward. Largest dike, in north-central part of quadrangle, contains many zones, across its 50-m thickness, distinguished by variations in grain size (medium-grained to very coarse) and mica and garnet content; a border zone contains resurgite garnets whereas a central zone contains red-brown garnets. Pale as green-brown in outcrop as wide as 30 m present, and intersected, in an unmapped dike in northeast corner of map area near mouth of Robinson Creek. Most dikes strike east-northeast and dip steeply. Possibly consecutive unmapped dikes of this orientation are present at Old Man of the Mountain. Dike thickness 10 to 60 m. Tentatively correlated with muscovite pegmatite dikes described by Lee and Barnes (1997), which intrude the biotite monzogranite unit 2 km south of the southwest corner of the map area.
- Tm** Biotite monzogranite (Oligocene and Eocene?) Medium- to fine-grained, equigranular; massive to moderately foliated biotite monzogranite. Color index (biotite content) 5-10 vol.%. Forms three mapped and numerous unmapped small dikes and irregular-shaped bodies a few meters wide. Similar bodies from several nearby areas in the Ruby Mountains yielded U-Pb zircon ages 29±0.5 Ma (Wright and Snoke, 1983; MacCready and others, 1997). Recent monazite U-Pb dating by J.E. Wright (unpublished) suggests an age no younger than late Eocene for one body near the quadrangle.
- Tgt** Granite gneiss of Thorpe Creek (Eocene) Garnet-biotite-muscovite monzogranite and leucogranite gneiss. Fine- to medium-grained; equigranular; pegmatitic clots present locally. Ubiquitous sparse garnets 1 mm wide. Locally contains sillimanite. Color index (content of mafic minerals biotite and garnet) averages about 5 vol.%. Strongly foliated; present as mylonitic gneiss in most exposures. Locally exhibits either west-trending mineral-along-trend lineation or north-trending mineral-alignment lineation. White to very light gray; rust-colored weathering bands common. Commonly weathers as irregular rock towers. Forms a sill as thick as 80 m intruded into marble of Verdi Peak in core of Lamolite Canyon fold nappe. Monazite U-Pb ages 36-39 Ma determined on one sample from the adjacent Lamolite 7.5-nappe Quadrangle and two samples from the quadrangle (RM-21 and RM-24) (Wright and Snoke, 1983; MacCready and others, 1997). The two samples dated from this quadrangle exhibit a north-trending lineation, whereas the sample dated from the Lamolite Quadrangle exhibits a west-trending lineation.
- Td** Hornblende-biotite diorite (Eocene?) Medium-grained. On eastern front of Ruby Mountains near Dads Creek. May be late Eocene in age like quartz diorite gneiss (40-3 Ma) and quartz diorite intrusions (38-2 Ma) in East Humboldt Range dated by U-Pb method on zircon by Wright and Snoke (1993).
- Kp** Pegmatitic granite (Late Cretaceous) Well-foliated to nearly massive leucogranite gneiss and leucogranite. Pegmatite; inequigranular; grain size variable but generally coarse to very coarse. K-feldspar grains containing graphic intergrowths of quartz are common and may be as large as several tens of centimeters across. Contains muscovite, biotite, commonly garnet or sillimanite. Biotite and sillimanite commonly in folia or mats that wrap around feldspar grains. Characterized chemically by Lee and others (2002), their K₂O unit. Widespread in quadrangle, but mapped only locally (east-central part of quadrangle) where metasedimentary relics absent. The unmapped pegmatitic granite pervasively and intimately invades host metasedimentary rocks as pods, sills, and dikes; location and linear proportions of resulting migmatite are indicated by overprint pattern on the metasedimentary map units. Mapped and unmapped pegmatitic granite makes up approximately 50 vol.% of rocks in this part of Ruby Mountains. Preliminary U-Pb age on monazite approximately 84 Ma determined by J.E. Wright (unpub. data referred to by Snoke and others, 1997) from unmapped pegmatitic granite gneiss in roadcut in southwest part of quadrangle (sample locality TM-4). Tertiary Eocene monazite U-Pb age determined for a pegmatitic granite body in the Lamolite Quadrangle (unpublished data of J.E. Wright).
- Of m** Marble of Verdi Peak (Ordovician and Cambrian protolith age) Calcite marble, siliceous calcite marble, graphitic marble, calc-silicate rocks, minor mica schist, sillimanite-mica schist, paragneiss, and hornblende-biotite schist. Commonly layered. Mylonitic in northwestern exposures. Diopside and actinolite impart green color to calc-silicate rocks. Rusty-weathering graphitic paragneiss interlayered with graphitic marble common at stratigraphic base of the unit. Upper half of unit mostly marble, locally graphitic or containing metachert, minor schist. Initially intruded by pegmatitic granite and equigranular granite veins in sills, dikes, and irregular bodies; the proportion indicated by overprint pattern. Locally (in original lower plate of Ogilvie thrust fault) intruded by sills and irregular bodies of gabbroic rocks, the proportion indicated by separate overprint pattern. Structural thickness 20 m to about 2 km (including roughly 50 vol.% granite); too not exposed. Correlated lithologically on the basis of stratigraphic position to Cambrian limestones and shales and Ordovician Popogrip Group, having a combined probable thickness of about 3 km. Includes:
- Of md** Brown dolomite Massive, rust-brown to white, medium-grained (3 mm) dolomite marble and interstitial calcite marble. Coarse tremolite laths and interstitial calcite common. Dolomite marble typically crops out as rounded knobs surrounded by brown or white sand of disintegrated marble. Intruded by pegmatitic granite; the proportion indicated by overprint pattern. Present in stratigraphically low part of marble of Verdi Peak near Hidden Lakes uplift. Thickness a few meters to about 100 m. Similar rock occurs in lowest part of marble of Verdi Peak 12 to 15 km south and southwest of the quadrangle, and in lowest part of the Cambrian limestone section 30 km south of the quadrangle near Green Mountain (Howard, 1971).

Fz Metamorphosed Prospect Mountain Quartzite (Cambrian and Neoproterozoic protolith age) Tan- to pale brown-weathering, medium-grained micaceous metagranite and quartzite schist. Contains 3-10 vol.% K-feldspar, less plagioclase, and 4-8 vol.% biotite plus muscovite. Sillimanite present southeast of mapped sillimanite isograd. Northwestern and structurally high quartzite exposures are fine-grained and fluggy, exhibit conspicuous mylonitic foliation and lineation. Structurally deep, more eastern quartzite exposures are nonmylonitic, medium-grained, and form resistant light-brown cliffs. The two textural types grade into each other. Intensely foliated by pegmatitic granite, the proportion indicated by overprint pattern. Present in Lamolite Canyon fold nappe. Structural thickness 10 to 150 m; base not exposed.

Fzpm Metamorphosed Prospect Mountain Quartzite (Cambrian and Neoproterozoic protolith age) and McCoy Creek Group (Neoproterozoic protolith age), undivided Tan- to pale brown-weathering, medium-grained micaceous metagranite and quartzite schist; local marble and biotite schist. Rusty mylonitic. Forms resistant light-brown cliffs. Dominant quartzite lithotype contains 3-10 vol.% K-feldspar, less plagioclase, 4-8 vol.% biotite plus muscovite, minor sillimanite. Biotite schist locally present. White calcite marble and interlayered quartz-diorite rock and tremolite-calcite rock present locally. Quenched exposures north of lower Witherington Creek, mostly marble and less schist, may alternatively correlate with marble of Verdi Peak (unit Ofm). Uppermost quartzite locally graphitic and rusty-weathering in southern part of Hidden Lakes uplift. Initially intruded by pegmatitic granite, the proportion indicated by overprint pattern. Present in Hidden Lakes uplift and root zone of Soldier Creek nappes. Presence of upper part of the Neoproterozoic McCoy Creek Group of Meach and Hazzard (1962), stratigraphically below Prospect Mountain Quartzite, indicated by calc-silicate and schist layers similar to those present in upper part of the McCoy Creek Group in the Ruby Mountains 25 km to the south of the quadrangle (Hudec, 1990; Jones, 1999). Structural thickness >700 m. This compares to an approximate undifferentiated stratigraphic thickness elsewhere in Ruby Mountains of Prospect Mountain Quartzite of about 1,200 m plus at least 75 m (down to the calc-silicate horizon) of upper subunit of the underlying McCoy Creek Group (Hudec, 1990).

See accompanying text for references and a discussion of the geology of the Verdi Peak Quadrangle.

Lithologic contact Dashed where inferred (cross section only)

Fault Dashed where inferred, dotted where concealed. Ball on downthrown side of normal fault.

Premetamorphic thrust fault Dotted where concealed. Metamorphosed, concordant to foliation; lacks calcareous structures. Identified as contact that repeats mapped stratigraphic sequences.

Right-side up Sawtooth on upper plate. Places older over younger rocks.

Overturned Sawtooth point into older rocks, into originally higher plate.

Sillimanite isograd Delineates presence of sillimanite in muscovite-biotite K-feldspar bearing metagranite. S on sillimanite side; the metagranite lacks sillimanite minerals on the opposing side of the isograd. Host-rock-specific isograd; sillimanite locally occurs in schists outside this host-rock-defined sillimanite zone.

Upright syncline

Upright anticline Showing plunges direction.

Projected recumbent axial plane Hinge line along fold nose defined by stratigraphic top of metamorphosed Prospect Mountain Quartzite. Drawn and projected based on where the hinge line pierces successive valley walls. Arrow shows plunging direction. Dash-dotted where projected above ground. Dotted where concealed.

Strike and dip of foliation

↖ 32° inclined ⊗ Horizontal ↗ Vertical

↖ Bearing and plunge of lineation in mylonitic zone. Mineral elongation and stretching; local folding in marble. May be combined with foliation symbol. Tail of arrow located at measurement site.

↖ Bearing and plunge of lineation in infrastructure and in rock fabrics transitional between mylonitic zone and infrastructure. Mineral alignment, compaction, rodding, and where oriented mineralite minerals (folded lithologic layering) intersection lineation. Tail of arrow located at measurement site. May be combined with foliation symbol.

↖ Bearing and plunge of mesoscopic fold. Folded layering. Tail of arrow located at measurement site.

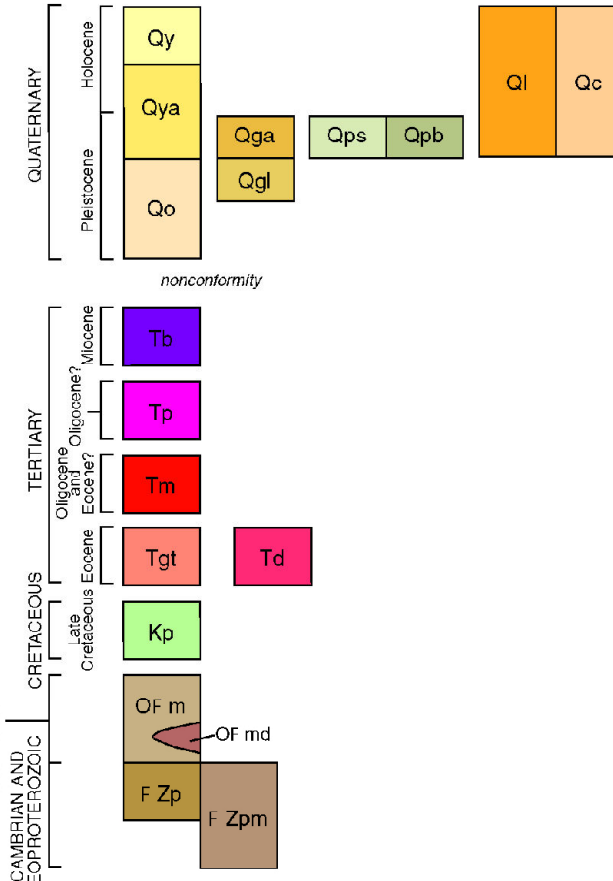
Moraine crest

Groundwater discharge zone and subtle scarp Concentration of lush vegetation and springs on bajada surface on younger alluvium in Ruby Valley. Tracks contour around coarsening alluvial fans at about 6,200-foot elevation. Southern part straddles a subtle scarp or sloping nose approximately 3 m high and tens of meters wide where the bajada slopes as steeply as 3 degrees. 2 degrees steeper than adjacent parts of bajada. Boulders as wide as 2-3 m concentrate on the sloping bajada surface below scarp. On air photo, resembles other constant-elevation groundwater discharge zones on alluvial fans in Nevada (cf. Smith and others, 1979, figs. 12, 19). Scarp morphology suggests possible fault origin. Boulder concentration would be compatible with origin as lag at a wavecut pluvial shoreline, but a wavecut origin here rejected because the scarp rises ~34 m in elevation above clear shoreline features of pluvial Lake Franklin, and 16-40 m above three possible outlet passes where pluvial lakes could have spilled from Ruby Valley northward into Clover Valley and the Humboldt River drainage.

Proportion of granite within mapped metasedimentary rock units Dominantly gneissic pegmatitic leucogranite gneiss and pegmatitic leucogranite of the pegmatitic granite map unit (Kp); lesser medium-grained two-mica granite gneiss and biotite granite gneiss; local medium-grained biotite-hornblende quartz diorite gneiss. Dominantly Cretaceous; may include Tertiary rocks.

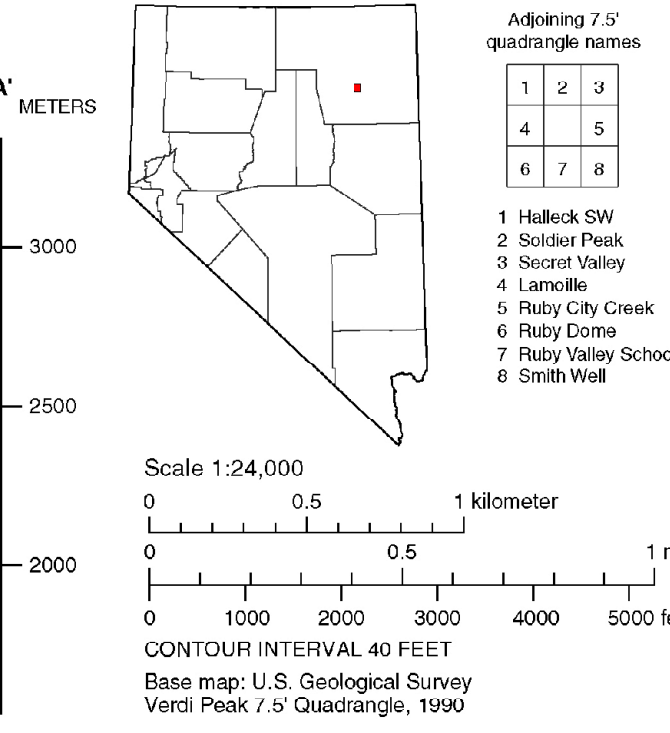
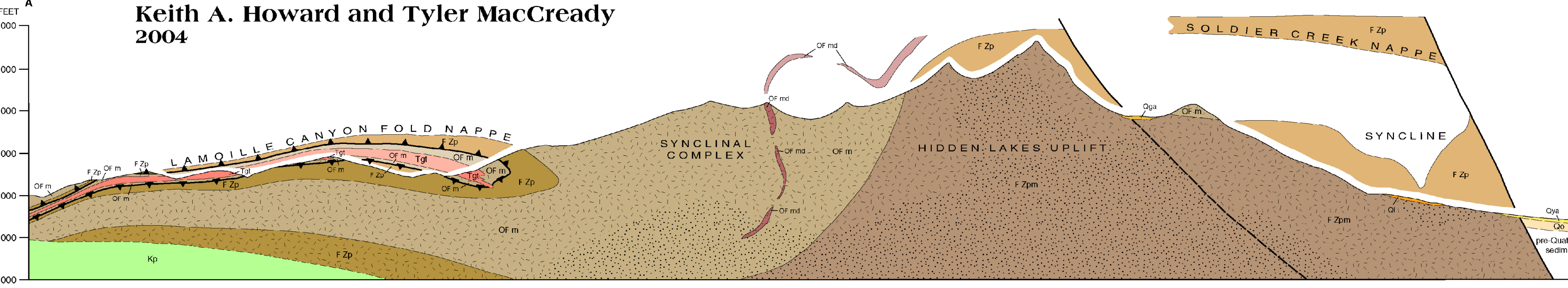
Areas where small bodies of metamorphosed gabbro occur within mapped marble of Verdi Peak In southwest part of quadrangle. Sills and dikes of metamorphosed gabbro. Mineral assemblages include various combinations of clinopyroxene, hornblende, cummingtonite, biotite-phyloppite, olivine, plagioclase, intruded by small sills and agmatite veins of leucogranite gneiss. Gabbro may be of Tertiary or Mesozoic age.

Sample dated by U-Pb method (MacCready and others, 1997) RM-4 dated by J.E. Wright, unpub., cited in Snoke and others (1997).



GEOLOGIC MAP OF THE VERDI PEAK QUADRANGLE, ELKO COUNTY, NEVADA

Keith A. Howard and Tyler MacCready
2004



Field work performed from 1963-1967, 1973, 1983. Mapping largely by Howard (1966). Compiled and revised by Howard, incorporating revisions to northern third of area after MacCready and others (1997) and unpublished field map of Horse Mountain provided by George Dams (California State University at Northridge) from work by his 1973 field class.

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