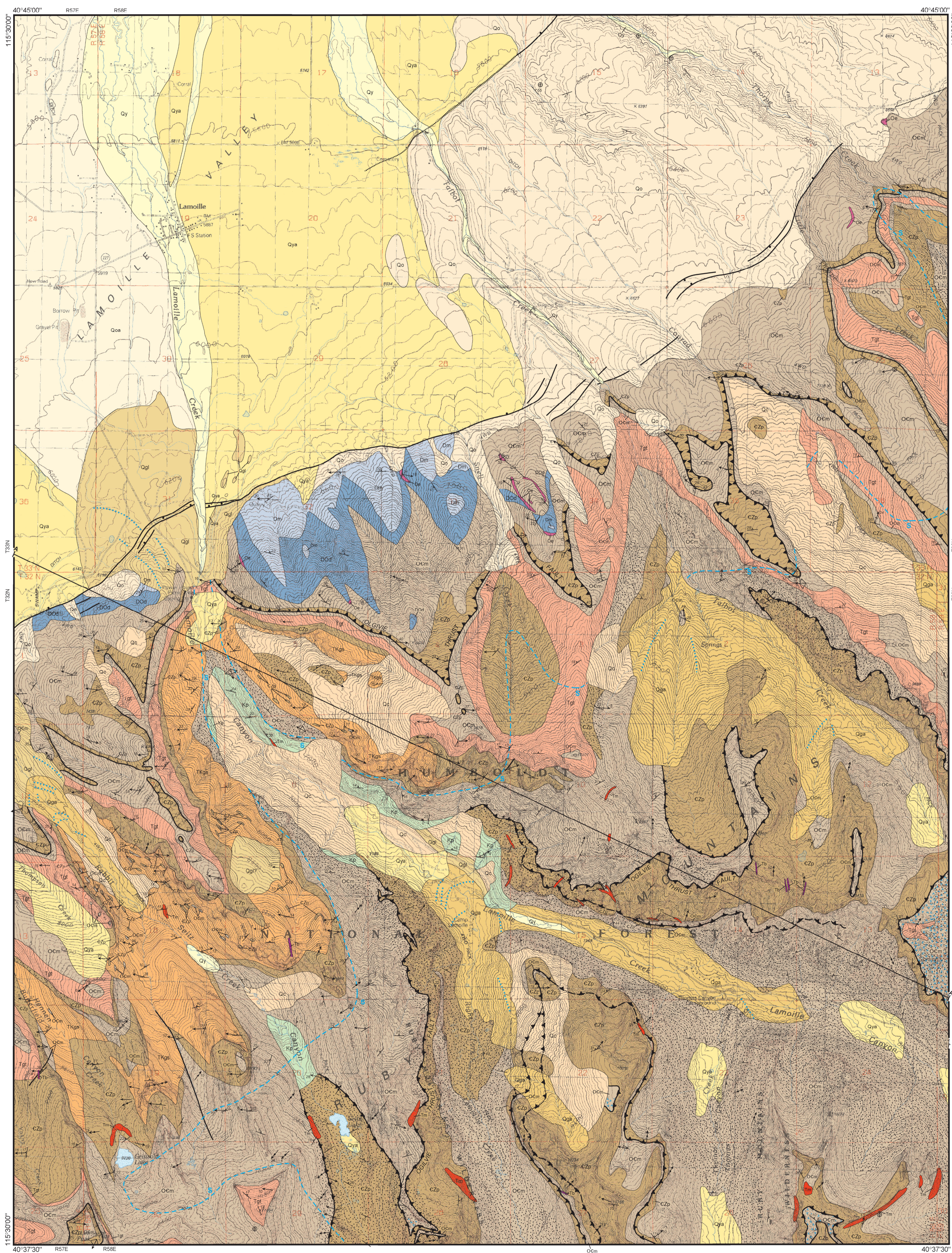


GEOLOGIC MAP OF THE LAMOILLE QUADRANGLE, ELKO COUNTY, NEVADA

Keith A. Howard
2000



Qc Colluvium and steep-slope alluvium (Holocene and Pleistocene) Confined to Ruby Mountains. Largely slope wash, talus deposits, and debris-flow deposits. Includes possible solifluction deposits. An age older than Angel Lake moraines is inferred for most deposits downslope from Angel Lake end moraines (Wayne, 1984). Includes now-breached landslide dam 50 m high across Lamaille Creek 3 km northwest of Camp Lamaille, consisting of house-sized and larger blocks of granite gneiss.

Qy Youngest alluvium (upper Holocene) Gravel, sand, and silt along present streams.

Qya Younger alluvium and outwash (Holocene and Pleistocene) Alluvial fan deposits of gravel and sand. Includes outwash from Angel Lake glacial advance.

Qga Angel Lake moraines Deposited by the Angel Lake glaciation (Angel Lake substage or stage). Boulders on surface and weakly developed soil described by Wayne (1984). Predates a peat carbon-dated at 13,019 ka (Wayne, 1984). Correlated with Tloga and Wisconsin drifts (Wayne, 1984).

Qgl Lamaille moraines Deposited by the Lamaille glaciation (Lamaille substage or stage). Forms prominent lateral and piedmont moraines. Surfaces exhibits fewer boulders than do Angel Lake moraines; thick soil profiles comparable in development to those on Bull Lake drift in the Ruby Mountains (Wayne, 1984). Wayne estimated an age an order of magnitude greater than Angel Lake moraines, and proposed correlation with Illinoian drift.

Qoa Older outwash and alluvium (Pleistocene) Glacial outwash from Lamaille glacial advance and other alluvial fan deposits of gravel and sand. Proximal deposits gravel and sand; distal deposits (northwest of Lamaille) gravel and interbedded calcite-cemented sandstone, siltstone, and claystone. Gravel clasts mostly metamorphosed and typical of Ruby Mountains, and (northwest of Lamaille) very scarce opaline shale, sandstone, and vesicular basalt. Undissected to moderately dissected.

Qo Oldest alluvium (lower Pleistocene) Forms and underlies dissected pediment terraces 50 to 100 m above present stream grade. Sharp (1940) proposed several terrace subdivisions. Alluvium is composed of metamorphic and granitic rocks derived from the Ruby Mountains. Mostly poorly consolidated, but locally includes calcite-cemented conglomerate and sandstone, and calcite-cemented sandstone. Dips suggest terrace is pediment cut on locally tilted beds. Steep 55° dip measured near north-central edge of map suggests landslide or nearby fault. Exposures 2 km north of the quadrangle include megabreccia derived from metamorphosed Eureka Quartzite, granite, brown metaquartzite, and low-grade marble. Thickness at least 100 m. Possibly includes deposits as old as Pliocene.

Tm Basalt dikes (Miocene) 1.5 km NE, 3.5 km SSE, and 2 km SW of Thomas Canyon Campground; also 2 km N of SW corner of quadrangle. Olivine basalt and diabase. Contains plagioclase phenocrysts as long as 2 cm in core of one dike. Locally contains amygdules. Dike margins chilled. Commonly fractured parallel to dike walls. Weathers to form topographic slots or trenches. Most dikes strike north and dip steeply.

Tgt Biotite monzogranite (Oligocene) Medium- to fine-grained, equigranular. Massive to moderately foliated. Map patterns generalized. Forms dikes, sheets, and small irregular bodies; not all are mapped. Age about 20 Ma determined from U-Pb dating of zircon (Wright and Snook, 1993; MacCreedy and others, 1997).

TKgs Granodiorite gneiss of Seltz Canyon (Tertiary or Cretaceous) Medium-grained biotite granodiorite gneiss. Biotite content about 7-10 percent. Locally contains sparse muscovite and garnet. Generally an augen gneiss having coarse feldspar grains (mostly perthite); very coarse-grained feldspar grains and pegmatite clots also common. Foliation defined by flakes and thin clusters of biotite. Texture xenomorphic; relict hypidiomorphic-granular texture locally present. Locally exhibits either west-trending mylonitic stretching lineation or north-trending folds. Contains interbedded thin layers of pegmatitic granite gneiss. Interlayered with adjacent metaquartzite and calc-silicate rock. Cut by pegmatite dikes. Forms massive light-gray cliffs. Forms sheet 700-800 m thick that occupies core of the Lamaille Canyon fold nappe and terminates southeastward to mimic a hinge of the nappe. Envelops tabular rafts of the metamorphosed Prospect Mountain Quartzite and marble of Verdi Peak, which outline a ghost stratigraphy and structure suggestive of folding and deformation before the gneiss was emplaced. Interlayered with and locally crosscut by pegmatitic granite gneiss, which suggests Seltz Canyon unit may overlap in age with the pegmatitic granite unit. As mapped, includes gneiss of fine- to medium-grained biotite monzogranite (about one tenth of unit), which cuts the main granodiorite gneiss phase and may correlate with the biotite monzogranite unit.

Tgt Granite gneiss of Thorpe Creek (Eocene) Garnet-biotite-muscovite monzogranite and leucogranite gneiss. Fine- to medium-grained; pegmatite clots occur locally. Ubiquitous sparse garnet 1 mm across. Locally contains sillimanite. Color index (content of dark minerals biotite and garnet) averages 4-5 percent. Strongly foliated; is mylonitic gneiss in most exposures. Locally exhibits either west-trending or north-trending lineation. White to very light gray, rust-toned weathering bands common. Base interlayered with adjacent metaquartzite. Forms a sill averaging 60 m thick intruded along metaquartzite-marble contact in the upper limb of the Lamaille Canyon fold nappe, where base of sill interlayered with adjacent metaquartzite sill estimated to be southeastward into marble of Verdi Peak in core of the nappe. U-Pb dates on monazite 36(±3) Ma from a sample from lower Lamaille Canyon and 2 samples from the adjacent Verdi Peak 7.5 minute Quadrangle (Wright and Snook, 1993; MacCreedy and others, 1997).

Kp Pegmatitic granite (Late Cretaceous) Well-foliated to massive leucogranite gneiss and leucogranite. Mostly pegmatitic; inequigranular grain size variable but generally coarse to very coarse. K-feldspar locally as coarse as 1 m. Locally includes equigranular granite gneiss. Contains muscovite, biotite, commonly garnet or sillimanite. Mapped where metasedimentary rafts or relict layers or relict pelitic seams nearly absent. Elsewhere in the Ruby Mountains, similar pegmatitic leucogranite and leucogranite gneiss pervasively invades host metasedimentary rocks as pods, sills as thick as 100 m, and dikes; the location and granitic proportions of the resulting magmatite are indicated by overprint pattern on the metasedimentary map units. U-Pb age on monazite approximately 84 Ma determined by J. E. Wright (unpub. data; see MacCreedy and others, 1997) from pegmatitic granite gneiss outcrop 0.5 km east of quadrangle.

Dm Marble of Snell Creek (Devonian protolith age) Fine-grained, layered light- and dark-gray calcite marble. Contains disseminated dolomite and graphite. Calcite megacrysts 4-6 mm across that occur in fine-grained calcite matrix near the base of the unit may represent relict crinoid columns. Brecciated at range front near Snell Creek (north-central part of sec. 23, T. 23 N., R. 58 E.). Maximum exposed structural thickness 50 m. Correlated with Guilmette Formation (middle to upper Devonian).

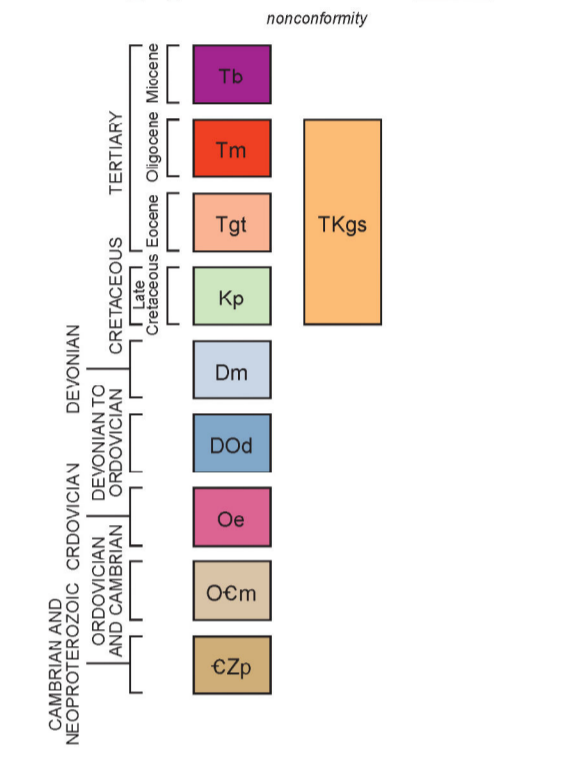
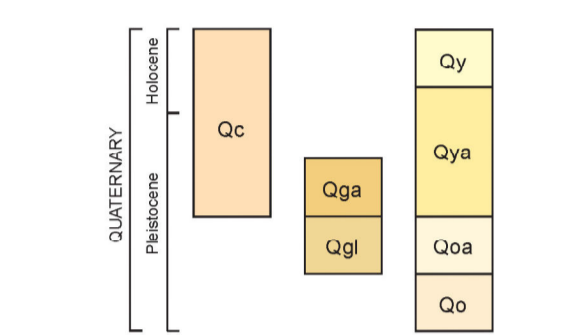
Dd Dolomite (Devonian to Ordovician protolith age) Sugary-textured massive white dolomite marble; weathers white, light gray, or pale buff. Grain size 1/2 mm. Commonly field when freshly broken. Contains small amounts of graphite, tremolite, and diopside. Maximum structural thickness 60 m. Correlated to the sequence Fish Haven, Laketown, Sevy, and Simonson Formations.

See accompanying text for references and a discussion of the geology of the Lamaille Quadrangle.

Oc Metamorphosed Eureka Quartzite (Ordovician) Massive to mylonitic white metaquartzite. Consists nearly entirely of quartz. Maximum thickness 3 m. Thickers northeastward; local absence may in part represent stratigraphic discontinuity.

OCm Marble of Verdi Peak (Ordovician and Cambrian protolith age) Calcite marble, siliceous calcite marble, graphitic marble, calc-silicate rocks, minor mica schist, sillimanite-mica schist, paragneiss, and amphibolite. Mylonitic in northwestern exposures. Diopside and actinolite impart green color to calc-silicate rocks. Rusty-weathering graphitic paragneiss common at the stratigraphic base of the unit where it is inverted, intricately intruded by pegmatitic leucogranite, pegmatitic leucogranite gneiss, and biotite granite gneiss in sills, dikes, and irregular bodies, the proportion indicated by overprint pattern. Correlated to Cambrian shale and limestone formations and Ordovician Pogonip Group. Structural thickness 20 m to 1 km.

CZp Metamorphosed Prospect Mountain Quartzite (Cambrian and Neoproterozoic protolith age) Tan- or brown-weathering, medium-grained micaeous metaquartzite and quartzose schist. Contains 3-10% K-feldspar, less plagioclase, and 4-8% biotite plus muscovite. Sillimanite present southeast of mapped sillimanite isograd. Northwestern and structurally high exposures are fine-grained and taggy, somewhat conspicuous mylonitic foliation and lineation. Deeper and more eastern exposures are nonmylonitic, medium-grained, and form resistant light-brown cliffs. Intricately intruded by pegmatitic leucogranite, pegmatitic leucogranite gneiss, and granite gneiss in sills, dikes, and irregular bodies, the proportion indicated by overprint pattern. Structural thickness 300-400 m.



Lithologic contact

Fault Dotted where concealed. Ball on downthrown side of normal fault.

Prematamorphic thrust fault Dashed where inferred, dotted where concealed. Metamorphosed; concordant to foliation; lacks cataclastic structures. Identified as contact that repeats mapped stratigraphic sequence; dashed where inferred across map units.

Right-side up Sawtooth on upper plate. Places older over younger rocks.

Overturned Sawtooth point into older rocks, into originally higher plate.

Sillimanite isograd In micaeous, muscovite-K-feldspar-bearing metaquartzite. S on sillimanite side; the metaquartzite lacks aluminosilicate minerals on the other side of the isograd. Host-rock-specific isograd; the first appearance of sillimanite in schists occurs outside the host-rock-defined sillimanite zone. Muscovite content of metaquartzite decreases and K-feldspar content increases in going up-grade across the sillimanite isograd.

Strike and dip of bedding
 Inclined ⊕ Horizontal

Strike and dip of foliation
 Inclined ⊕ Horizontal

Bearing and plunge of mylonitic lineation
 Essentially all lineations that strike west-northwest are observed to be stretching lineations in mylonitic textures, so these lineations are distinguished here. May be combined with foliation symbol.

Bearing and plunge of mineral or intersection lineation Defined by elongate minerals such as sillimanite, intersection lineations, and rodding. May be combined with foliation symbol.

Bearing and plunge of mesoscopic fold

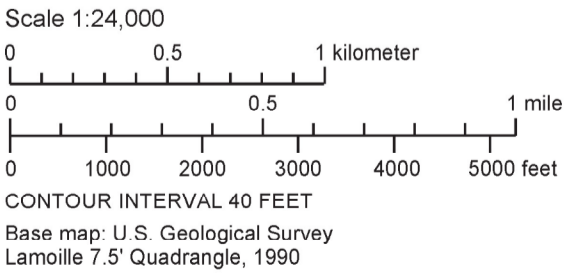
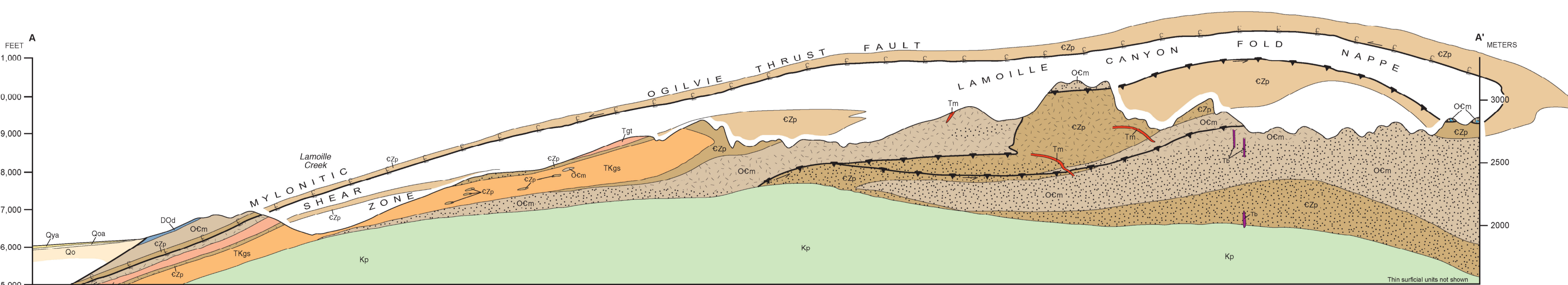
Shear-sense direction of diastrophic folds
 Arrow indicates apparent drag sense of upper rocks over lower as determined from faceted directions of upper short limbs in sets of local, variably oriented mesoscopic folds. Separation arc indicates degree of uncertainty (Howard, 1965; full data plotted in Howard, 1966).

Moraine crest

□ <1/3 □ 1/3 - 2/3 □ >2/3

Proportion of the metasedimentary rock map units in the Ruby Mountains composed of pegmatitic granite Dominantly pegmatitic leucogranite gneiss and pegmatitic leucogranite, lesser biotite granite gneiss, scarce biotite-hornblende granite gneiss.

Areas where small bodies of metamorphosed gabbro occur within mapped units Metamorphosed gabbro sills and dikes containing clinopyroxene, hornblende, cummingtonite, biotite, plagioclase, olivine, plagioclase. Intruded by small sills of leucogranite gneiss. Intrusion mainly or entirely into marble of Verdi Peak unit where in original footwall of Ogilvie thrust fault. May be of Tertiary or Mesozoic age.



Field work done in 1963-65, 1996-1997, assisted by Lee Wilson, 1963.
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