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INTRODUCTION

The energy shortage has led to accelerated exploration of possible alternatives to US dependence on fossil fuel reserves, and as part of this effort a number of exploratory projects are under way to assess technological and environmental aspects of constructing and operating geothermal power plants. Of the environmental problems that might be associated with the utilization of geothermal resources, the possibility of affecting seismicity has been of particular concern.

The possibility of triggering earthquake activity by the reinjection of geothermal fluids into deep wells is based on known cases of changes in seismicity associated with fluid injection or reservoir loading. For example, Evans (1966) showed that earthquakes near Denver, Colorado, were linked to injection of waste fluid into basement rocks beneath the Army's Rocky Mountain Arsenal, and Raleigh and others (1972) carried out experiments to determine the reservoir pressure threshold needed to trigger earthquakes in an oil field near Colorado. There have also been numerous cases of earthquakes Rangely, associated with the impoundment of water in large reservoirs; six of these earthquakes had magnitude 5-6.5 (Simpson, 1976). Simpson concluded that the potential for induced seismicity appears to be highest in areas of strike-slip or normal faulting and areas of at least moderate strain accumulation. Most reservoir sites where major changes in seismicity have occurred (Kariba, Koyna, Hsinfengkiang, Hoover and Oroville) are close to areas of high seismicity or had low-level seismicity near the reservoir before impoundment. Areas of low strain accumulation or areas characterized by horizontal compression (thrust faulting) appear to have lowest potential for induced seismicity.

The relationship between induced seismicity and pore-pressure change has been clearly shown for earthquake activity associated with high-pressure fluid injection (Healy et al, 1968; Raleigh et al, 1976; Fletcher and Sykes, 1976). The effect of pore pressure on seismicity has been tentatively explained by the work of Hubbert and Rubey (1959), who showed that the fracture strength of rock is proportional to the difference between total normal stress across the fracture and the pressure of pore fluids within the rock. Simpson (1976) notes that increases in pore pressure in the case of fluid injection (a few hundred bars) are much higher than those (a few tens of bars) created by a deep reservoir. In cases where fluid injection has triggered earthquake activity (Denver, Rangely, Matsushiro, Dale) injection has taken place into, or very near a fault zone.

In recent years, Dixie Valley, north-central Nevada, has been the focus of geothermal exploration involving surface and drilling investigations by a number of oil companies. The area of most intensive study is between Dixie Hot Springs on the south and Sou or Seven Devils Hot Springs on the north. This area is a graben, capped by a gabbroic intrusive at a depth varying from approximately 1,500 to 2,000 feet. The investigation was intended to integrate industrial and academic research efforts to produce a comprehensive geothermal reservoir assessment of the central part of Dixie Valley.

Based on Simpson's (1976) criteria, Dixie Valley would have relatively high potential for induced seismicity should geothermal development of the area lead eventually to injection of geothermal fluids into deep wells. Features that make the area favorable for induced earthquakes are the following:

• The Dixie and Sou hot springs are on the surface trace of the Dixie Valley fault zone (Trexler et al, 1978).

• Significant movement, in a normal faulting sense, occurred on this zone in 1954 (Slemmons, 1957).

• The area of interest is just south of a zone that Wallace (1978) believes is a "seismic gap," that could be filled by a single ML 7.4 earthquake.

• The area of interest is at the north end of the aftershock zone of the 1954 Fairview Peak-Dixie Valley earthquakes, a zone characterized by moderate-to-high seismicity (Ryall, 1977).

Because of the high potential for induced seismicity, a detailed microearthquake study was undertaken by the University of Nevada in parallel with other studies in the Dixie Valley area. The primary objective of this study was to provide a baseline for assessment of changes in seismicity associated with future injection of geothermal fluid into deep wells if the Primary objectives of the is developed commercially. geothermal field investigation were to operate a network of seismic stations in Dixie Valley, and to analyze and interpret the network data. The DOE contract supporting this study began on 16 August 1979 and was initially written to cover a one-year The Seismological Laboratory installed ten seismic stations in and effort. around Dixie Valley in December 1979, and the network began full operation in January 1980. Primarily because of high equipment reliability and very low seismicity the rate of expenditure on this project was lower than originally anticipated, and the period of performance was extended from 12 to 27 months at no additional cost to DOE.

SEISMIC CYCLE IN THE WESTERN GREAT BASIN

A number of papers have been published treating various aspects of the seismic risk problem in Nevada. Slemmons et al (1965) compiled a catalog of earthquakes for the Nevada region for the period 1852-1960. Ryall et al (1966) studied seismicity of the western United States based on earthquake catalogs for various regions. Slemmons (1967) used photogrammetric methods to map Pliocene and Quaternary faults within the Great Basin province, and grouped the faults according to degree of weathering and other age-related parameters. Ryall et al (1974) studied microearthquake activity for 1970 and 1971, and compared the distribution of these events with locations of large historic earthquakes and active faults. Douglas and Ryall (1975) used earthquake recurrence statistics for 1932-1969 to determine average acceleration return periods from calculations involving distance to the causative fault and magnitude. Wallace (1977, 1978) studied the morphology of young fault scarps in north-central Nevada to estimate average recurrence times in that region. Ryall (1977) reanalyzed historic and current seismicity patterns to determine the character of the seismic cycle, as well as foreshock and aftershock sequences in western Nevada. VanWormer and Ryall (1980) and Ryall and VanWormer (1980) related earthquake hazard and maximum magnitude to structure and active tectonic processes along the Sierra Nevada-Great Basin boundary zone, and recommended changes in seismic zonation for the western Great Basin.

Large Historic Earthquakes.

During the historic period since about 1840, five major (M > 7) earthquakes have occurred in the western Basin and Range province. Approximate rupture zones for these events are shown on Figure 1, and descriptions are given below.

- <u>1845(?)</u> <u>Stillwater area(?)</u>. At the time of the 1869 Virginia City earthquake the Gold Hill News reported that a Washoe Lake Piute aged about 30 years stated that there had been a great earthquake when he was a little boy, while Indians from the Walker, Carson and Truckee Rivers were fishing at the Carson Sink. The shock knocked people down, river banks were shaken down in the vicinity of Stillwater, and the river changed its course. Slemmons et al (1965) placed this event farther to the west, near Pyramid Lake, and gave it a date of 1852(?) based on an 1865 Indian report of ground failure in the Pyramid Lake area and landsliding along the Sierra Nevada near Reno.
- <u>March 26, 1872, Owens Valley, California</u>. In Lone Pine, out of a total population of 250-300, 23 were killed and 60 injured; 52 out of 59 houses (mainly adobe) were destroyed. Faulting along the east side of the Owens Valley extended for at least 70 km, with scarps up to 7 meters high. The earthquake was felt over 640,000 sq miles, which according to Oakeshott et al (1972) would correspond to magnitude around 8.
- October 2, 1915, Pleasant Valley. This earthquake was accompanied by faulting for a distance of 30-40 miles along the western face of the Sonoma Range, with scarps 2-4 m high. Damage was greatest at Kennedy, where every building was destroyed; in Winnemucca nearly every chimney was toppled, walls were cracked and a few thrown down. The shock was felt over 500,000 sq miles. Richter (1958) estimated that it had magnitude 7.6.



Figure 1. Generalized map of late Cenozoic structural features of the western Great Basin (Wright, 1976), together with epicenters of earthquakes for the period 1969-1978 (dots) and approximate rupture zones of major historic earthquakes (stippled areas, with year of the main shock).

- <u>December 20, 1932, Cedar Mountains</u>. The Cedar Mountains earthquake had magnitude 7.3 (Richter, 1958) and was felt throughout Nevada, California and Utah. Faulting consisted mainly of fissures with occasional evidence of vertical or lateral slip (Gianella and Callaghan, 1934). Vertical scarps were small, usually only a few inches; the zone of surface rupture extended for about 60 km in a northwest-southeast direction, and was 6-14 km wide.
- December 16, 1954, Fairview Peak-Dixie Valley. In this sequence, a magnitude 7.1 shock at Fairview Peak was followed 4 minutes later by a magnitude 6.8 event about 55 km to the north in Dixie Valley (Romney, 1957). These earthquakes were accompanied by surface ruptures in two zones trending slightly east of north. The southern zone, associated with the Fairview Peak shock, was 50 km long and 10 km wide; the Dixie Valley rupture zone to the north was about 40 km long and 5 km wide. The Dixie Valley zone had more than 2 m of dip-slip movement, while the Fairview Peak zone had about 4 m each of vertical and horizontal movement. The two shocks did little damage because the region of maximum shaking was very sparsely populated; however, they were felt over 220,000 sq miles in eastern Oregon, Idaho, Utah, California and northwest Arizona. The earthquakes in December followed by only a few months the Fallon-Stillwater earthquakes of July 6 and August 23, 1954. The first of these, with magnitude 6.6, caused damage at Fallon; the second had magnitude 6.8. The combined rupture zone of the two events had length 40 km and was a few km wide; it was parallel to and about 30 km west of the Dixie Valley zone.

Active Faults.

From the geologic literature, active faulting in the western Great Basin is distributed rather evenly over much of the region and not confined to a single belt analogous to the San Andreas zone in California. Figure 2, based on a paper by Slemmons (1967), shows faults in late Quaternary alluvium, lake sediments of the playa, Bonneville or Lahonton type, or glacial deposits. Approximately a thousand faults are shown on the figure, ranging in length from about 1 to more than 100 km. Presumably faults shown as continuous for several tens of kilometers on the figure were so mapped because they had fresh, continuous scarps that could be clearly identified on aerial photographs. Faulting associated with the 1915 Pleasant Valley and 1954 Dixie Valley-Fairview Peak earthquakes appears in the center of the figure as an almost continuous, 200-km long zone; in contrast, the 1932 Cedar Mountains rupture, just south of the 1954 zone, consists of about 20 small faults or fissures, none of which is more than a few kilometers in length.

In recent studies of fault-scarp morphology, Wallace (1977,1978) estimated the most recent age of displacement for 19 clusters of faults in an area of 17,000 sq km in north-central Nevada. He concluded that no more than seven major events occurred in this area in Holocene time, leading to average re-rupture times of 4,000-30,000 years for the various rupture zones in the study, based on the assumption that a typical rupture zone has area of about 1,000 sq km, and that such rupture zones comprise the entire 225,000 sq km region containing Holocene scarps in the western Great Basin. Wallace also observed that fault-scarp morphology in the rupture zones of major historic earthquakes does not suggest faulting at a greater rate than in surrounding areas.



Figure 2. Map of faults that show photogrammetric evidence of Quaternary displacement (adapted from Slemmons, 1967).

In the Dixie Valley area, Wallace (1981, personal communication) has concluded that the 1954 earthquake was preceded by faulting concentrated in the north end of the rupture zone, and that this area acted as a nucleation point from which the rupture propagated to the south at the time of the main shock. In support of this conclusion he observes that a series of 4-5 scarps in the vicinity of Dixie Hot Springs had recurrent movement during Holocene time, while scarps to the south did not, that the 1954 shock produced only small displacement in the northern part of the valley and large displacements to the south, and that Romney's (1957) epicenter for the Dixie Valley earthquake was at Dixie Hot Springs (39.8 deg N, 118.1 deg W, depth about 40 km).

Wallace (1978) has also identified a seismic gap, between the 1915 and 1954 rupture zones, which he describes as follows:

"Between the southern end of the 1915 Pleasant Valley scarp and the northern end of the 1954 Dixie Valley scarp is a segment of fault-generated range front which has not broken in historic time. The gap is about 40 km long, and includes the precipitous east flank of the northern end of the Stillwater Range."

"This precipitous segment of range front is about 20 km long and rises approximately 1,000 m from the floor of Dixie Valley. A slope of between 30 and 35 degrees is maintained over much of the scarp height. The steepness of this scarp suggests a rate of uplift much higher than the regional average and thus accelerated uplift in the last few million years. If uplift of 1,000 m is assumed to have occurred in 2 million years, the rate of uplift is 0.5 m per 1,000 years. If a 3-m displacement event produces an M7 event, the average interval of time between such events would be 6,000 years, discounting the effect of small events and tectonic creep."

"The age of the latest scarp at the base of the east flank of the Stillwater Range is not well determined, but a preliminary analysis of its morphology suggests that the latest displacement is no older than about 12,000 years (Holocene). The basal part of the scarp, which probably was produced in a single event, is approximately 8 m high and, corrected for erosion, represents displacement of approximately 5-6 m."

"If the 40 km-long gap were to be filled by a single fault event having an average displacement of 5 m, an earthquake of approximately 7.4 ML could be generated."

Rates from Instrumental Data.

The results of Wallace and Pease provide an opportunity to compare recurrence rates determined from paleoseismic studies with those obtained from lists of instrumentally recorded earthquakes. Ryall (1977) gives a recurrence rate of

$$\log N = 4.85 - 0.784 M$$

based on 2,000 earthquakes recorded in the entire Nevada region from 1970 to

1974. For the period 1932 to 1969, Douglas and Ryall (1975) give a recurrence rate of

$\log N = 6.48 - 0.91 M$

for the western Great Basin. Assuming that the rupture zone for a typical large earthquake has an area of about 1,000 sq km and that such rupture zones comprise the entire 225,000 sq km in which Holocene faulting has occurred, these recurrence rates lead to re-rupture times of 7,000-10,000 years.

Foreshocks and Aftershocks.

With regard to foreshocks, a search of 13 northern Nevada newspapers for the historic period prior to 1917 (Appendix A) and instrumental recordings after that time (Appendix B) indicates that both the 1915 Pleasant Valley and 1954 Dixie Valley-Fairview Peak earthquakes were preceded by moderate seismicity for at least several decades prior to the main shock. A comparison of reported intensities for communities around the meizoseismal area of these events indicates that at least eight earthquakes with M = 4.0-4.5 occurred in the Pleasant Valley area during a 43-year period prior to 1915, and five shocks with M = 4.5-5.5 were located in the Dixie Valley area during an 82-year period prior to 1954. The level of activity for the two zones was similar: each had a return period of 5 to 6 years for events with M > 4. Taking into account differences in area of the rupture zone of a great earthquake and area of the region containing Quaternary faulting, this indicates a rate of foreshock activity several times higher than the average rate of occurrence of events with M > 4 in the Nevada region during the 1970's.

As observed in other regions, large shocks in the Great Basin may also be preceded by a period of quiescence following the general increase in moderate seismicity mentioned above. This is illustrated by Figures 3 and 4, which show the distribution, respectively, of individual earthquakes with M > 4 and of energy released in the Dixie Valley-Fairview Peak area before 1954. On Figure 3, earthquake activity in central Nevada (area bounded by latitude 38 to 40 deg N, longitude 117.5 to 119 deg W) is shown as a function of time and latitude, for the period 1930-1962. The figure is based on locations and magnitudes given by Slemmons et al (1965) for earthquakes with M > 4. On the left side of the plot, activity during the first half of the **1930's** represents mainshock/aftershock activity in the Cedar Mountains-Excelsion Mountains. Seismicity in the Fairview Peak area is represented by a few events in the late 1930's, a 12-year period of quiescence from 1940-1952, a burst of 11 events in late 1952 (largest M = 4.8) in the vicinity of the 1954 main shock, a 19-month gap, and intense activity before the main shock.

The same sort of increase in activity following a period of quiescence is shown by the plot of energy release on Figure 4. This figure also suggests that the increase in seismicity preceding the 1954 earthquakes was distributed over a sizeable region, and not simply concentrated in the immediate vicinity of the impending main shock. A similar observation, of a period of quiescence followed by intense foreshock activity distributed over a sizeable area, was made by Ryall and Ryall (1981) for the recent Mammoth Lakes earthquake sequence. For the 1915 Pleasant Valley shock, however, the pattern was somewhat different. As indicated by the newspaper accounts in Appendix A, there were reports of six or seven earthquakes from 1872 to 1900 that were probably located in Pleasant

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Figure 3. Distribution of earthquakes in central Nevada as a function of latitude and time. Plot includes events with latitude 38-40N, longitude 117.5-119W and $M \ge 4.0$ (data from Slemmons et al, 1965).



Figure 4. Energy release as a function of latitude and time for central Nevada. Plot includes same data as Figure 3, but period is 1945-1954. A(2.0) is the logarithm of the number of ML 2.0 earthquakes that would be equivalent to the total energy released by earthquakes within a 6-km by 75-day grid square. Numbers were smoothed with a triangular filter before log N(2.0) was calculated. Energy was determined from Richter's (1958) formula log E (ergs) = 11.8 + 1.5 M.

Valley, and no events in that area were reported in the period 1901-1914. However, the only foreshock activity following this 15-year period of quiescence appears to have consisted of two strong foreshocks within six hours of the main shock.

With regard to aftershocks, Figure 5 shows the rate of microearthquake activity for 1970-1975 in each of the zones that had major earthquakes in the western Great Basin during the historic period. The ordinate on this figure shows the average number of earthquakes located in each of the rupture zones per year, and the abscissa is the time in years after the main shock. The figure clearly indicates that the rate of earthquake occurrence in these zones is inversely proportional to time after the main shock, and that aftershock activity following a typical large earthquake in this region decays to a minimum level after about a century. This is almost an order of magnitude longer than the 15-year aftershock sequence Fedotov (1968, 1971) found for major earthquakes in the western Pacific (see below), and in general supports the long re-rupture times discussed above for faults in the Great Basin.

One final observation about aftershock activity in the Dixie Valley is illustrated by Figure 6, which shows energy released in that area as a function of latitude and time, for the period 1 July 1954-30 June 1981. Following the main shock, shown by the large spike in the lower left corner of the figure, the seismicity appears not only to decrease in terms of energy release, but also to migrate south as residual stress in the northern part of the zone is relieved.

Seismic Cycle in the Western Great Basin.

Fedotov (1968, 1971), studying long-term behavior of earthquakes in the western Pacific seismic belt, concluded that within, the rupture zone of a typical great earthquake in that region there is a "cycle" of activity that lasts for about 140 years, and that this cycle has three parts. During the first 15 years after a catastrophic earthquake, the level of activity -- expressed in equivalent number of small (M about 3) earthquakes needed to achieve the total energy release -- decreases by a factor of 1,000 to some minimum level. Following this aftershock period the activity stabilizes at the minimum level for more than a century; during any 5-year period there is high probability that the activity will be within a factor of two of this level. During the final 10-15 years of the cycle there is a foreshock sequence in which the activity gradually increases, reaching a level about 2.5 times the minimum just before another great earthquake occurs.

While tectonic processes in the western Pacific region are different than those in the Great Basin, Fedotov's seismic cycle, based as it is on a large number of major earthquakes, provides a model for comparison with earthquake occurrence in the Nevada region in general, and Dixie Valley in particular. From the preceding discussion, we draw the following conclusions:

• The "seismic cycle" in Nevada is of the order of thousands of years long. In support of this conclusion we note that no two of the five major historic earthquakes in the western Great Basin have occurred in the same rupture zone, that fault-scarp morphology studies in northern Nevada indicate re-rupture times of thousands of years, and that recurrence rates based on current seismicity are in agreement with long re-rupture times.



Figure 5. Average number N of earthquakes per year for 1970-1975 within the rupture zones of large (M > 7) historic earthquakes, as a function of time T in years after the main shock.



Figure 6. Energy release as a function of latitude and time for northern Dixie Valley, latitude 39.5-40.2N, for the period 1 July 1954 - 30 June 1981. A(2.0) is the logarithm of the number of ML 2.0 earthquakes that would be equivalent to the total energy released by earthquakes within a 2-km by 197-day grid square. Numbers were smoothed with a triangular filter before log N(2.0) was calculated.

- A typical large (M > 7) earthquake in the Great Basin is followed by an aftershock sequence lasting about a century, which gradually relieves residual stresses in the rupture zone. Seismicity in the rupture zone then stabilizes at some minimum level for a long period of time. Most of the western Great Basin is observed to have a background level of minor seismicity, and the distribution of small earthquakes is often difficult to identify with particular faults on the surface.
- Foreshock activity in the Great Basin appears to consist of a moderate increase in seismicity in the zone of an impending rupture, occurring over a period of at least several decades before the main shock. In the final "preparation stage," there may be a period of quiescence lasting for several years, followed by an intense foreshock sequence in the months preceding the main shock. In the case of the 1915 Pleasant Valley earthquake, the main shock was preceded by about 15 years of quiescence, which appears to have ended only hours before the large earthquake.
- The area of this study, Dixie Valley, was affected by a large shock in 1954 that generated an average of 2 m displacement along normal faults in a zone 40 km long and 5 km wide. The Dixie Valley shock occurred 4 minutes after the Fairview Peak earthquake, and filled part of a seismic gap between that event and the 1915 Pleasant Valley rupture zone. The magnitude given by the California Institute of Technology for this earthquake was 6.8, based on surface-wave amplitudes. However, Romney (1957) notes that body-wave amplitudes at teleseismic distances were greater for the Dixie Valley shock than for the M 7.1 event four minutes earlier, and attributes the different partitioning of energy between body- and surface-waves to the fact that the second event was deeper (40 km) than the first (25 km).
- Qualitative analysis of activity in the Dixie Valley area since 1954 indicates a general decrease in energy release appropriate for an aftershock sequence. There also appears to have been a migration of activity toward the south, presumably as residual stresses in the northern part of the valley have been relieved over the last 27 years.
- According to Wallace (1978) the area of most importance to this study -northern Dixie Valley between Dixie Hot Springs and the Sou Hills -- is a seismic gap that has the potential for an earthquake with maximum magnitude of about 7 1/2. This 40-km long zone between the 1915 Pleasant Valley and 1954 Dixie Valley rupture zones is marked by a precipitous segment of range front bounded by Holocene scarps up to 8 m high; Wallace concludes that the average interval between movements along this zone is about 6,000 years, and that the latest movement is at least several thousand years old.

GEOLOGIC/STRUCTURAL SETTING

Stations of the Dixie Valley seismic network (DVSN) were located in and around Dixie Valley, a typical Basin-Range valley about 100 km long and 20 km wide, extending in a NNE-SSW direction in central Nevada (Figure 7). The valley is bounded on the west by the Stillwater Range and on the east by the Clan Alpine Mountains. According to Willden and Speed (1974):

"The Stillwater Range is a north-trending horst bounded by high-angle faults of large vertical displacement. Normal faults of less displacement cut the block, and differential displacements on the blocks within the horst have tilted erosion surfaces in various directions. Some normal faults in the interior of the range extend to the flanks and continue as range-front faults (Page, 1965). Faulting and earthquakes have occurred at a number of places in historic time in the vicinity of the Stillwater Range. . . The range contains several units of deformed Mesozoic rocks separated by thrust faults, and is the center of a large complex of mafic igneous rocks, and a remarkable succession of volcanic and intrusive rocks of probable Cenozoic age."

"The Clan Alpine Mountains trend northward from their southernmost point at Westgate on US Highway 50 to their intersection on the north with the Augusta and New Pass Mountains. The range contains the highest point in Churchill County, Mount Grant, 9,966 feet. The present configuration of the Clan Alpine Mountains is a result of block faulting, much of which is The western margin of the range is underlain by Tertiary Holocene. volcanic rocks along much of its length, and these rocks are downthrown by normal faults relative to the mountain range proper lying to the east. The structure suggests that the west flank of the Clan Alpine Mountains may descend gradually by step faults of relatively small displacement toward the floor of Dixie Valley. The opposite side of Dixie Valley differs because Quaternary deposits are faulted directly against a block of predominantly Mesozoic rocks. . . The Clan Alpine Mountains expose three successions of rocks of Mesozoic age, [and the northern part of the range] contains the easternmost exposures of the [Jurassic] Humboldt gobbroic complex. • The Clan Alpine Mountains contain contrasting terranes of Tertiary volcanic rocks. The northern part of the range in the Shoshone Creek drainage is partly covered by accumulations of rhyolitic ash flows that may be as much as 2,000 feet thick."

According to Whitney (1980), photogeologic analysis indicates that the valley is a complex graben bounded by high-angle Basin and Range faults. The main fault system in the area is the Stillwater fault, which bounds the the Stillwater Range on the southeast and trends N36E from Dixie Meadows into Pleasant Valley. This zone is marked at the surface by very fresh scarps in alluvium and bedrock, and geomorphic features indicate that the zone is very active. Whitney states that the fault dips 50-60 deg SE in the northern part of the valley, and that motion on the fault is primarily dip-slip, with a minor right-lateral slip component. South of Dixie Meadows, which marked the northernmost extent of primary faulting in December 1954 (Slemmons, 1957), the fault trends almost N-S (Figure 7). As illustrated by Figure 8 (from Whitney, 1980), the Dixie Valley graben is asymmetric, with a number of NW-dipping step faults bounding the southeast side of the valley.



Figure 7. Generalized geologic and geophysical map of the west Humboldt, Stillwater and Clan Alpine Ranges, Nevada. Subsurface structures inferred from geophysical data (from Smith, 1967). Hachured area is not referred to in the present report. Solid triangles -- stations installed for this study (Table 1).



Figure 8. Three-dimensional model of the northern portion of Dixie Valley. Structural relationships among the various tectonic elements are depicted, with alluvium removed and the bedrock surface restored (from Whitney, 1980).

Wallace (1979, 1980) proposes a tectonic model for the Basin and Range that involves listric faulting in an upper, brittle crustal plate in response to regional extension concentrated in narrow zones of anelastic deformation or intrusion at depth. Diagrammatic cross sections illustrating this model for the northern Great Basin are shown in Figure 9. In Figure 9a, the model is constructed to incorporate (1) eastward tilting of the ranges (defined by the dip of basalt flows that cap most of the ranges); (2) first- and second-order relationship of ranges, with some blocks having rotated and slumped off other blocks; (3) Mohorovicic discontinuity at depth about 30 km; (4) maximum depth of seismicity -- corresponding to maximum depth of the brittle zone -- about 15-18 km; (5) narrow zones of extension at depth; (6) listric form of faults in the brittle part of the crust; and (7) zone of decoupling to accomodate the listric style of faulting. Figure 9b is more speculative, and includes (1) zones of extension possibly invaded by intrusive (A on figure); (2) vertical propagation of tension cracks through the listric faults (B on figure); (3) complexity of the glide blocks (C on figure) incorporating blocks of various size and some unrotated blocks; (4) sequence of events, including earlier westward tilting of the Humboldt Range (D on figure) and cross-cutting of older by younger faults; and (5) a zone of decoupling that is complex rather than simple, and includes glide planes biased toward westward sliding and toward the local zones of extension (E on figure). Wallace notes further that:

"Scarps that developed in Holocene time (last 12,000 + years) appear to be distributed in narrow, elongate belts trending N10-25E (see figure 10). The belts cross range blocks, for example, the belt produced in 1915 (see 1915, figure 10) crosses four range blocks, and the Western Cortez belt (see WC, figure 10) trends at high angle across the Cortez Range front fault trend. Within other belts such as the Shoshone Range (see SR, figure 10) belt an the Humboldt Range belt (see HR, figure 10), faults of diverse trends appear reactivated. This independence of the belts from obvious surficial structures suggests that the elongate belts of reactivation relate to deep structures, and that the effects are propagated upwards and become superimposed across upper crustal blocks. Regional extension appears to be normal to the long axes of these belts."



Figure 9. Possible interpretations of fault and tectonic data by Wallace (written communication). See text for explanation.



Figure 10. Figure by Wallace (written communication), showing distribution of Holocene scarps in narrow, elongate belts. Independence of these belts from surficial structures is evidenced by the belts crossing range blocks. Wallace suggests that the "belts of reactivation relate to deep structures, and that the effects are propagated upwards and become superimposed across upper crustal blocks. Regional extension appears to be normal to the long axes of these belts."

INSTRUMENTATION AND ANALYSIS

Dixie Valley Station Sites.

As shown on Figure 7, stations of the Dixie Valley Seismic Network (DVSN) were located in and around the valley; station data is listed in Table 1, and site descriptions are given below.

The 3-component HYX station was located approximately 150 feet beyond the adit to the abandoned Hoyt antimony mine, and station BYX was about 50 feet inside an abandoned mine near the Boyar Ranch. The remaining stations were buried about two feet deep in soil. The BOX station was buried on a granitic outcropping; HYX, DIX and CHX were on gabbro; CBX was on basalt; HLX, BYX and SOX were on rhyolite; GRX was at the upper end of an old alluvial fan; and FPN is on rhyolite on the NE side of Fairview Peak, south of Dixie Valley. In addition to the ten stations in and around Dixie Valley, two nearby stations of the Nevada Seismic Network were used in the analysis: Station BMN, 55 km NE of station SOX, is in a mine tunnel in Paleozoic metasediments, and KVN is situated at the bottom of a 700-foot deep mine shaft, 18 km SSE of FPK, on rhyolite.

As indicated by Table 1, seven of the DVSN stations began operation in early January, 1980. Stations at FPN and CBX were added in April. The mine station near Boyar Ranch was vandalized in August, 1980, and all components except the seismometer were stolen; the seismometer and new electronics were installed at CHX about a week later. Figure 11 shows the amount of up-time for all the stations, and indicates that with the exception of CBX, CHX and HLX the stations were in operation more than 95% of the 22.5 months of the experiment. Station CBX was down about a third of the time, HLX had about 14% down time, and CHX was inoperable about 8% of the time. In mid-November, 1981, six of the stations were removed, leaving DIX, FPN and HYX (together with BMN and KVN of the permanent network) continuing to monitor activity in the Dixie Valley-Fairview Peak area.

Instrumentation.

Seismometers. Eight of the stations listed in Table 1 were equipped with Teledyne-Geotech model S-13 seismometers. The S-13 weighs less than 25 pounds and has a 5-kilogram mass; it is designed for use in field operations where a small, light-weight, short-period, moving-coil type seismometer is desired. All of the seismometers had free period of 1.0 second, and 3,600-ohm output coil. The S-13 may be used in either vertical or horizontal configuration, and two of the seismometers were installed as horizontal-component instruments at station HYX. The seismometers used at FPN, KVN and BMN were Teledyne-Geotech models 1501 (vertical-component) and 1101 (horizontal) Benioff seismometers. These instruments have been standard for many networks, including the Worldwide Seismic Station Network (WWSSN); they have the advantage of high sensitivity to small ground motion, but the disadvantage of being cumbersome (about 450 pounds and too large for field use). The Benioff transducer is a balanced, variable-reluctance type, constructed so as to produce a voltage proportional to earth velocity. All of the Benioff instruments had period of 1.0 second; station FPN had only a vertical-component seismometer, and BMN and KVN were triaxial.

	TABLE 1.	STATION	DATA, DIALE	VALUEI OUI			÷
ĨD	Station Name	Latitude deg N	Longitude deg W	Elevation meters	VCO Hz	Start Date	Seis Model
BMN	Battle Mountain	40.4313	117.2217	1450	2040 (Z) 1020 (E) 2720 (N)	09/02/69	1051
BOX	Box Canyon	39.6153	118.1763	1283	2380	01/11/80	S13
BYX	Boyer Ranch	39.9542	117.9178	1128	2040	01/11/80	S13
CBX	Chocolate Butte	39.9741	118.1557	1439	2720	04/03/80	S13
CHX	Cottonwood Canyon	39.9859	117.8674	1234	2040	08/21/80	S13
DIX	Dixie Hot Springs	39.8022	118.0820	1143	- 1020	01/10/80	S13
FPN	Fairview Peak	39,2028	118.1550	2256	2720	04/01/80	1051
GRX	Grover Canyon	39.6355	117.9890	1390	1360	01/10/80	S13
HLX	Hole-in-the-Wall	39.9337	117.6632	1219	680	01/10/80	S13
НҮХ	Hoyt Mine	39,7728	117.7633	1661	2380 (Z) 2040 (E) 1700 (N)	01/10/80	S13
KVN	Kaiserville	39.0510	118.1000	1829	1360 (Z) 1700 (E) 680 (N)	01/13/72	1051
SOX	Sou Hot Springs	40.1017	117.71.67	7 1198	1020	01/11/80	S13

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STATION DATA, DIXIE VALLEY SEISMIC NETWORK TABLE 1.



Figure 11. Record of station operation. Station down-time is indicated by gaps in bar chart. Time is by weeks, from 16 December 1979 to 14 November 1981. The VCO, transmitter and batteries were stolen from station BYX on 5 August 1980, and the station was reinstalled at CHX on 21 August 1980. Station CBX was installed on 3 April 1980, but was not operational until 6 May when a relay was installed on Toulon Peak. Stations FPN, DIX and HYX are still operating in January 1982, as single-component stations.

<u>Amplifier/VCO's</u>. Stations with S-13 seismometers had Emtel model 6242 preamp/VCO's. These units amplify the seismometer output and modulate a voice-range carrier frequency to permit multiplexing of up to eight channels of information on a single data link. The 6242 is a small, all solid-state, low-power device capable of high amplification in a field environment. All of the preamps were set to -18dB attenuation, equivalent to a voltage gain factor of approximately 3,000; VCO center frequencies are given in Table 1. Stations FPN and CHX had older type Emtel model 6202 preamp-VCO's; EMN has Teledyne model 4300 phototube amplifiers and Emtel model 6207 VCO's; KVN has preamp/VCO's designed and constructed at the University of Nevada.

<u>Telemetry</u>. Signals from most of the stations were transmitted to the Reno facility on VHF radio links. Except for CHX the S-13 stations used Monitron model T15 transmitters and R15 receivers — small, rugged units that are capable of low-power operation. Signals from HLX, HYX, GRX and DIX were multiplexed on one link, and signals from SOX, BYX (or CHX) and BOX on another. All of these signals, together with FPN and KVN, were relayed from Fairview Peak, where a single solar panel was used to charge six lead-acid batteries to run two one-watt VHF radios. Signals from Fairview Peak were received at Slide Mountain, south of Reno, and then transmitted over telephone lines to the University. The CBX signal was telemetered to a relay on Toulon Peak, near Lovelock, from there to Virginia Peak, north of Reno, and from there to the University. Data from BMN were transmitted to Reno over telephone lines.

<u>Recording</u>. At, the Seismological Laboratory, the multiplexed seismic signals were recorded in direct mode on a Bell and Howell model VR-3700B, 1-inch tape transport. For analysis, individual earthquakes were played out from the tape onto a 16-channel chart recorder (Siemens Oscillomink), together with time marks from a chronometer and a WWVB radio-time signal (Figure 12).

<u>Calibration of the S-13 Stations</u>. Equipment at the S-13 stations was new when the network was installed. At that time the seismometers were bench-checked for free period, and the VCO's were checked for center frequency and deviation. Step-input and frequency response checks were made at only one site, the three-component station HYX. Since the results on all three components were identical, and since all the equipment was new, it was assumed that operating characteristics of the other stations would be identical to HYX, and as a result only step-input calibration was performed at those stations. Results of calibration are shown on Figure 13.

Stations HYX, HLX, GRX, SOX and DIX were found to be very nearly identical in sensitivity. Station BOX had only half the gain of the other stations, apparently because of a faulty or maladjusted amplifier. Station CHX, which had an older model amplifier after most of the original equipment was stolen from BYX, also had only half the gain of the other S-13 sites. Station CBX was not calibrated, but appears to be identical to HYX in operation, based on playbacks of teleseisms recorded by both stations.

Data Analysis.

Analysis of the DVSN data began in January, 1980, following installation of the network. Arrival times of crustal phases Pg, Pn and Sg were measured to an accuracy of about 0.02 second, together with direction of first motion, coda length, and for some of the smaller events, maximum amplitude. Calculation of



Figure 12. Example of playback for a small (ML 0.3) event in northern Dixie Valley (39.931N, 117.774, depth 7.2 km). Time marks are 10 seconds apart.



Figure 13. System magnification for 3-component station HYX. Pulse calibrations indicated that HYX, GRX, SOX and DIX had essentially identical magnification; stations BOX and CHX had one-half the gain indicated on this calibration curve.

hypocenter coordinates and magnitude were accomplished using the HYPO71 algorithm of Lee and Lahr (1975). For the period 1970-1975, most central Nevada earthquakes were located using a program NEVLOC, modified from an original algorithm described by Ryall and Jones (1964), although some of the better-recorded events were analyzed with the HYPO71 routine. For the period 1976-1979, most earthquakes were located using a program NEVLOC2, written by W. Peppin (described in the 1975-1979 and 1980 issues of the Bulletin of the Seismological Laboratory), but some events were located using HYPO71. For all of these calculations, the crust was assumed to consist of a single layer with P- and S-wave velocities of 6.0 and 3.5 km/sec, respectively, and thickness of 28 km. Results of analysis are tabulated in Appendix B, and discussed in the next section of this report. In Appendix B, events located using HYPO71 are those with quality A, B, C or D; events without quality listings were located using NEVLOC or NEVLOC2.

Magnitude.

Magnitude was determined from measurements of either maximum trace amplitude or coda duration. Most magnitudes for 1980 and 1981 were determined from coda length on playbacks similar to that on Figure 12. Magnitude for events located with NEVLOC were determined from measured trace amplitudes, either on Wood-Anderson recordings at Reno or on playbacks of signals recorded at BMN, TNP (Tonopah) or WCN (Washoe City). Magnitude for events located with NEVLOC2 for 1976-1979 were determined either from amplitude or coda measurements. In cases, constants in the magnitude equations as well as station correction factors were determined so that magnitudes for events in the Sierra Nevada- Nevada-Great Basin boundary zone would be consistent with those determined from the Reno Wood-Anderson recordings, or with University of magnitudes. Descriptions California, Berkeley, of the two magnitude calculations are given below.

<u>Amplitude-magnitude</u>. Empirical relations were developed by J. D. VanWormer between measured trace amplitude at stations Battle Mountain, Tonopah and Washoe City, in the form

$$ML = C1 + \log A + C2,$$
 (1)

where ML is "local magnitude" (see Appendix C), Cl is a correction for distance and C2 is a station correction factor. VanWormer found Cl to be

Cl	2 2	1.268	Ŧ	.027	D	D	less th	an 50) km	
		2.492	+	.005	D	D	= 50-40	0	,	
		3.613	+	.002	D	D	greater	tha	n 400	km.

In equation (1), A is the maximum, unclipped, peak-to-peak amplitude measured in mm on strip-chart recordings like that on Figure 12. To determine the correction factor C2 VanWormer compared magnitudes determined from Nevada network recordings with those of the University of California for 34 events with ML in the range 2.7-4.3. VanWormer's analysis was extended to the Dixie Valley network by comparing coda- and amplitude-magnitudes; station correction factors C2 are listed in Table 2a.

STATION	A. CORRECTION FACTOR F AMPLITUDE-MAGNITUDE	
	0.75 + 40	.09 ± .20
BMN	$-2.75 \pm .46$	
BOX	$-2.17 \pm .50$	$32 \pm .23$
BYX	$-2.46 \pm .46$	
CHX	$-2.75 \pm .58$	
DIX	$-2.56 \pm .44$.21 ± .28
FPN		.68 ± .19
GRX	$-2.52 \pm .47$	$.45 \pm .37$
HLX	$-2.41 \pm .34$	$.20 \pm .24$
		$.04 \pm .22$
SOX	$-2.73 \pm .26$.50 ± .21
HYX	$-2.62 \pm .50$	$.04 \pm .22$

TABLE 2. STATION CORRECTION FACTORS FOR MAGNITUDE DETERMINATION.

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<u>Coda-magnitude</u>. The coda-magnitude relation for short-period vertical instruments of the Nevada network is

$$ML = -1.2 + 2.65 \log T + 0.0013 D,$$
 (2)

and a set of

All All All All All

where T is coda duration in seconds, as defined by Lee and Lahr (1975), and D is epicentral distance in kilometers. The relation is based on a set of 12 events with Berkeley magnitude ML = 3.0-5.7, that occurred along the Sierra Nevada-Great Basin boundary zone. The standard deviation between BRK and UNR for these events is 0.09 ML. Duration data for nine Benioff stations and 21 other stations of the permanent network were analyzed separately. The coefficients of the coda duration and distance terms were found to be the same for the two sets of stations. The coda-magnitude calculation was extended to the Dixie Valley network by comparing magnitudes for 25 events at Mammoth Lakes with known magnitudes in the range ML = 3.1-4.9; station corrections used in this study are listed in Table 2b.

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 $\frac{\partial f}{\partial t} = \frac{\partial f}{\partial t} \frac{\partial f}{\partial t} + \frac{\partial f}{\partial t} \frac{\partial f}{\partial t} + \frac{\partial f}{\partial t} \frac{\partial f}{\partial t} + \frac{\partial f}{\partial t} + \frac{\partial f}{\partial t} + \frac{\partial f}{\partial t} \frac{\partial f$

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21 - 35 *5

INTERPRETATION AND DISCUSSION

Distribution of Earthquakes in Space and Time.

The spatial distribution of earthquakes in the Dixie Valley-Fairview Peak area is shown on several figures. Figure 14 shows the location of 1,128 earthquakes analyzed by the University of Nevada for the period 1 January 1970 to 30 June 1981, within the area 38.8-40.5°N, 117.3-118.7°W. Clusters in the lower part of the figure (latitude about 38.99N) are in the northern end of the The dense cluster in the center of the 1932 Cedar Mountains rupture zone. figure (latitude 39.0-39.5°N) is in the southern (Fairview Peak) part of the 1954 rupture zone, and activity in Dixie Valley (latitude 39.5-39.8°N) is in the northern part of the zone. Scattered epicenters in the upper right part of the figure are in the 1915 Pleasant Valley rupture zone, and the long cluster of events west of about 118.3°W is in the aftershock zone of two magnitude 6.8 shocks in July and August, 1954 (Rainbow Mountain, Stillwater earthquakes). In general, this figure indicates that aftershocks of the Dixie Valley-Fairview Peak earthquakes are distributed in a 90 km-long, 20 km-wide zone trending almost N-S. In the area of particular interest to this investigation -- between Dixie Meadows and Sou Hot Springs (DIX and SOX on the figure) only 11 events were detected for the entire 11.5-year period. Five of these -- with magnitudes in the range 0.1-1.6 -- were recorded by the DVSN in 1980 and 1981; the largest event in this area, in 1973, had ML 3.0.

Figures 15-18 show the distribution of events for four three-year periods -- 1970-72, 1973-75, 1976-78 and 1979-81 -- with numbers of events, respectively, 415, 392, 307 and 143. Although part of this decrease in annual number of events is an artifact of a fluctuating network detection threshold, there does appear to have been a real change in seismicity, starting in 1977 and continuing through 1981. This is illustrated by Figure 19, which shows the distribution of 447 events in the Dixie Valley-Fairview Peak zone as a function of time and latitude, and includes only earthquakes with ML greater than 2.0. On this figure the change in seismicity is particularly apparent in the southern part of the zone, where a change in density of the symbols corresponds to a change in the rate of activity: the average number of ML 2+ events in the southern part of the zone $(30.0-39.15^{\circ}N)$ was 29/year from 1970 to 1976, and only 6/year for the 4 1/2-year period 1/1/1977-6/30/1981.

A similar pattern of decreasing seismicity was reported by VanWormer and Ryall (1980) for the Sierra Nevada-Great Basin boundary zone in 1977 and 1978. This period of quiescence was followed by "a burst of sizeable earthquakes, aftershocks and swarms in all parts of the zone in late 1978 and 1979." As noted by Ryall and Ryall (1981), the increased activity in 1979 was most intense in the Mammoth Lakes area, culminating with several magnitude 6+ earthquakes on 25 Comparing several characteristics of the Mammoth Lakes and 27 May, 1980. sequence with behavior that various authors have tentatively identified as precursory to large earthquakes, Ryall and Ryall concluded that "precursory phenomena reported in the literature may, at least in some cases, have been the result of regional, rather than local stress changes. If so, such changes could create favorable conditions for large earthquakes to occur anywhere within a sizeable region, rather than in the immediate vicinity of some particular observation." The decrease in seismicity within the Fairview anomalous Peak-Dixie Valley aftershock zone was generally synchronous with that in the



Figure 14. Earthquakes in the Dixie Valley-Fairview Peak area, 1970-1981.





 $\mathcal{T}_{ij} = \mathcal{T}_{ij}$





33.



Figure 17. Earthquakes in the Dixie Valley-Fairview Peak area, 1976-1978. Area is same as Figure 14.

34.



Figure 18. Earthquakes in the Dixie Valley-Fairview Peak area, 1979-1981. Area is same as Figure 14.



Figure 19. Distribution of earthquakes in the 1954 Dixie Valley-Fairview Peak rupture zone as a function of time for 1970-1981. Only events with ML \geq 2.0 are shown; symbols are scaled as a function of magnitude.

Sierran frontal fault zone, supporting the idea that major changes in seismicity may occur on a regional scale.

Inspection of Figures 15-18 suggests that the Fairview Peak-Dixie Valley zone may be made up of a number of NW- and NE-trending fractures. This pattern was observed previously by Ryall and Malone (1971), who stated that

"The over-all trend of the Dixie Valley and Fairview Peak zones is about NO7°E, which agrees well with the strike of NO9°E given by Savage and Hastie (1969) for a dislocation model based on geodetic changes that accompanied the 1954 Fairview Peak earthquake. However, our results indicate that this zone is made up of a zigzag series of northwest- and fractures northeast-striking very much like the pattern of northwest-trending compression ridges and northeast-striking tensional cracks found by Larson (1957) in a study of minor features of the 1954 fault zone. The main Fairview Peak zone strikes about N12°W, and the sense of motion in this zone is right-lateral oblique slip, the east side moving southeast and down. At the southern terminus of the main zone is a southwest-trending line of fractures, and north of the main zone there appears to be an en-echelon series of northeast-striking faults."

Figure 20 shows all events from 1974 to 1981 for which the quality of the HYPO71 hypocenter determination was A or B (Lee and Lahr, 1975). Of the 42 events shown, 36 are in the 1954 rupture zone, and these appear to fall in three areas -- a N-S main zone about 60 km long, and terminal fractures at the north and south ends of the main zone that trend N30-40°E. An average trend for all the events is about N09°E, in agreement with Savage and Hastie's (1969) estimate of the trend of the 1954 dislocation surface.

One other point worth mentioning is that Figure 19 clearly indicates a general lack of activity in northern Dixie Valley, the area of most interest to geothermal exploration. On the figure, the 1954 rupture zone is south of the dashed line at about $39.77^{\circ}N$; 443 of the events shown on the figure are located south of the line, and only 4 are north of it. Thus, for the last decade the zone which Wallace (1978) identified as a seismic gap has been almost completely quiescent.

Focal Depth.

In a previous study, Ryall and Savage (1969) compared focal depths for a number of natural earthquake sequences in the western Great Basin, and concluded that average focal depth appears to correlate with the magnitude of the main shock: "the 1954 Fairview Peak earthquake, with magnitude about 7, had aftershock activity in the depth range 12-14 km along the main fault zone and 8-13 km in a terminal cross-fault; recent earthquake swarms with maximum magnitude 5.5-6.0 in southern Oregon and southeast Nevada had focal depths concentrated in the range 5-10 km, and the shallowest natural earthquake sequence observed so far in the Nevada region was the Truckee, California, series, with magnitude of the main shock about 5.3 and concentration of focal depths in the range 3-7 km."

In the present investigation, focal depths were compared for 217 earthquakes with location quality of C or better (Lee and Lahr, 1975) for the



Figure 20. Earthquakes with location quality A or B (Lee and Lahr, 1975), 1974-1981. Average trend of entire zone shown by heavy line, hypothetical division into main zone and terminal fractures by dashed lines. Area shown is same as top part of Figure 14.

period 1974-1981. The epicentral distribution of these events is shown on Figure 21, and the distribution of events as a function of depth is shown by the histogram on Figure 22. Analysis of 207 events with depth in the range 3-22 km gave a mean depth of 11.3 km and standard deviation of + 3.22 km. As Figure 22 indicates, the earthquakes are not evenly distributed but tend to cluster at depths of 10-12 km; this clustering is well approximated by a normal distribution. A histogram for 42 events with B-quality solutions showed less concentration in the 10-12 km range (shaded part of Figure 22), but the mean value of 10.2 + 2.9 km was not significantly different than that for the larger sample. For 34 events in the main Fairview Peak zone, 39.15-39.55°N, mean depth was 12.8 + 4.0 km, in agreement with the observation by Ryall and Savage that events in this part of the rupture zone were slightly deeper than at the ends.

On Figure 23, the location of events in different depth ranges is shown by outlines of different types. The figure supports the observation that seismicity in the 1954 rupture zone occurs in three areas -- a N-S central zone and terminal zones that trend NE-SW. Dip directions cannot be ascertained from the plot. The figure also indicates the orientation of two cross-sections shown on Figures 24 and 25.

On Figure 24, depths are shown for 131 events with quality of C or better in the southern part of the area studied, on a cross-section viewed in the direction N36E. The cluster on the right side of the figure is in the northern part of the 1932 Cedar Mountains earthquake; the dip of this zone is 58°SE and the largest focal depth is about 16.5 km. The cluster on the left side of the figure consists of events in the southern part of the 1954 rupture zone; mean depth in that area is 10.6 + 2.8 km, and most events are in the range 7-15 km. No dip can be seen for this part of the rupture zone.

Figure 25 is a cross-section viewed in the direction $NO8^{\circ}$ W, approximately the direction given by Savage and Hastie (1969) for the trend of the 1954 dislocation surface, and the figure contains only events for which the location quality was A or B. The figure would appear to rule out listric faulting in the main part of the 1954 rupture zone -- a dip of 59°E fits the points on the left (W) side of the figure, down to maximum depth of almost 17 km, and no flattening of the zone can be seen. On the right (E) side of the figure, two points suggest a possible dip of 50-60°W, and a cluster of points in the center of the figure could have a similar dip.

In the northernmost part of Dixie Valley, only five events were large enough to be analyzed, and these had depth of 7.2 km or less. A plot of these earthquakes viewed in the direction N4OE suggests flattening of the fault with depth (Figure 26), but much more data would be needed to substantiate such a conclusion.

Recurrence Rates.

Recurrence curves showing the cumulative number of earthquakes as a function of magnitude areauseful in comparing the seismicity of one area with another. Such a plot for the Dixie Valley area for 1974-1981 is shown on Figure 27, with magnitude restricted to ML 2.5 or greater; the figure also shows all western Great Basin earthquakes outside the Manmoth Lakes area for 1980, with magnitude ML 2.3 or greater. The straight-line portion of a recurrence curve









Figure 23. Map showing epicentral areas for different depth ranges: crosses -- 0-5 km; solid line -- 5-10 km; dashed line -- 10-15 km; dotted line -- >15 km.



Figure 24. Cross-section viewed in direction N36E for southern part of Fairview Peak and northern part of Cedar Mountains rupture zones. Symbols are not scaled to magnitude. Cluster on right dips about 58 deg SE. Plot includes 131 events with location quality of C or better.



Figure 25. Cross-section viewed in direction NO8E for main part of 1954 rupture zone. Plot shows 37 events with location quality of A or B. Dip of 59 deg SE for points on left of figure is in good agreement with dip found by Savage and Hastie (1969) for theoretical dislocation surface for the 1954 shocks.



Figure 26. Cross-section for events in Dixie Valley north of the 1954 rupture zone, viewed in the direction N40E. Only events with location quality of C or better are shown. There is a vague suggestion of listric faulting indicated by dotted line, but much more data would be needed to support such an interpretation.



Figure 27. Recurrence curves for the Dixie Valley-Fairview Peak area, 1970-1981 (dots), and for all non-Mammoth Lakes events in 1980 (squares). N is the cumulative number of events with ML equal to or greater than the specified value.

can be represented by the equation

 $\log N = a - bM$,

where N is the number of events with magnitude equal to or greater than M, a is an intercept that depends on the time period and area considered, and b is a constant that is thought to characterize seismicity in a given region (Gutenberg and Richter, 1954). B-values for the Dixie Valley and western Great Basin samples shown on the figure were determined by a maximum likelihood calculation (Aki, 1965) to be 1.29 and 0.60, respectively.

In general, relatively high b-values are found to be characteristic of aftershock sequences, earthquake swarms, and volcanic earthquakes (Mogi, 1966; Utsu, 1971; Hamilton et al, 1971; Wyss, 1973). A number of investigators (Eaton et al, 1970; Butovskays and Kuznetsova, 1971; Hamilton, 1972) found b to decrease with increasing focal depth, and Healy et al (1968) found it to increase with pore pressure for earthquakes associated with fluid injection. In the laboratory, Scholz (1968) observed different b-values for different rock types and found that b decreased with increasing stress in all cases. Explanations of the various sets of observations are in terms of an inverse dependence of b on applied stress (Scholz, 1968; Butovskaya and Kuznetsova, 1971; Wyss, 1973); moreover, the various authors concluded that higher b-values should be observed in regions (or at depths in the crust) where the number of inhomogeneities per unit volume is relatively large. In the Dixie Valley-Fairview Peak area, the relatively high average b-value would be expected in connection with ongoing aftershock activity following the large earthquakes there in 1954.

On Figure 28, the slope of the recurrence curve is shown as a function of time for the period 1970-1981. The b-values on this plot were determined for overlapping groups of 100 events each, with minimum magnitude in all cases taken as ML 2.0. The plot indicates that b increased from .80 in 1970 to 1.18 in 1972; in 1974 it began to gradually decrease, reaching a minimum value of 0.68 in 1978. Since that time fewer than 100 events with ML 2.0 or greater have occurred, so the value plotted in mid-1978 is the latest one we could calculate.

In a study of the 1980 Mammoth Lakes earthquake sequence, Ryall and Ryall (1981) observed that the b-value for earthquakes along the Sierra Nevada-Great Basin boundary zone gradually decreased from 1.00 to 0.61 over the period 1970-1979, and then rapidly increased to 0.99 before the occurrence of several ML 6+ events in mid-1980. This slow decrease followed by rapid increase agrees with a pattern that Scholz et al (1973) related to increasing effective differential stress, related in turn to dilatancy hardening preceding a large earthquake. If a similar process is occurring in Dixie Valley, the length of the "bay" -- period of low b-values -- of about 8 years would be appropriate for a major earthquake. Should this be the case, based on experience in 1954 and 1980 a significant increase in seismicity would be expected before the occurrence of a large event.



Figure 28. Running b-value for the Dixie Valley-Fairview Peak rupture zone, 1970-1981. Based on overlapping samples of 100 events each, with ML 2.0 or greater. Error bars indicate 95% confidence limits. Points are plotted at the center of the time period for each sample. Last point is in 1978 because too few events occurred after that time for accurate determinations of b.

Fault Plane Solutions.

In a previous study, Ryall and Malone (1971) compared epicenter distributions and composite fault-plane solutions for microearthquakes in the Fairview Peak area. They concluded that earthquake activity in west-central Nevada is directly related to regional extension, with the minimum principal-stress axis having an average trend of about $N60 \circ W$, and the intermediate and maximum principal stresses lying near a plane striking about N30°E and dipping 85°SE. The authors noted that this stress pattern is similar to that found for aftershocks of a nuclear explosion in southern Nevada (Hamilton and Healy, 1969), and concluded that similar stresses may be operating over a broad region in the western Great Basin. The one exception to this pattern was a fault-plane solution for earthquakes in Dixie Valley, where the stress axes were rotated about 50 degrees clockwise from corresponding axes in the Fairview Peak and Rainbow Mountain zones.

For the present study, fault-plane solutions were determined for eleven earthquakes with ML 3.0-4.5, in the area 39.0-40.0°N, 117.5-118.5°W. Depths of the events ranged from 3.4 to 14.7 km. The focal mechanisms (Figure 29) were based on first-motion patterns for clearly-recorded P arrivals -- primarily Pg but also including numerous Pn arrivals. Following Ryall and Malone (1971) the angle of incidence for Pn was assumed to be 50 degrees, appropriate for a single crustal layer with velocity 6.0 km/sec over a mantle with velocity 7.8 km/sec. Some of the observations used in the analysis were questionable, but omitting these points did not appreciably change the results.

The solutions on Figure 29 are of two main types. Five (labeled A to E on the figure) are consistent with primarily dip-slip, or normal faulting, on planes striking NE-SW and dipping NW or SE -- we shall refer to these as "Type I" mechanisms. The other six ("Type II"; F-K on the figure) range from strike-slip to oblique-slip, with a right-lateral slip component if the plane striking N to NE is taken as the fault plane. Four of the five Type I sources are SE of Fairview Peak (FPN), and the other (point E) is west of the Stillwater Range in the Rainbow Mountain fault zone. Four of the Type II sources are in the southern part of Dixie Valley, one is south of Fairview Peak, and one is in the north end of Smith Creek Valley, 50 km east of Dixie Valley. While the two types of mechanisms cannot be separated by area, they can be generally separated by focal depth: the Type I events had depths in the range 8.4-14.7 km and averaged 11.9 + 2.6 km; range and mean depth for the Type II events were 3.4-11.8 and 7.2 + 3.1 km, respectively. Descriptions of the two types of sources are given below. 13. 32. 10 mm - 1705 +

<u>Dip-slip mechanisms</u>. Solutions A-E all show two planes striking NE-SW, one dipping NW and the other SE. A decision as to which plane is the fault and which the auxiliary plane cannot be made from the figure: if the SE-dipping plane is arbitrarily selected as the fault, the parameters are those listed in Table 3. Note that the trend of the T-axes, or axes of minimum compressive stress for these events range from 44 to 75 degrees, in a NV-SE direction; the average trend is -61 + 11 degrees (azimuths NE, or clockwise being taken as positive). Average plunge of the T-axes is about 8 degrees. Of the five mechanisms, one (E) is based on few data and is a poor solution. Mechanisms A and B are not unique, and could also be interpreted in terms of a fault striking NW and dipping NE or SW; however, such a mechanism would be less probable than the ones shown on the basis of Basin and Range tectonics. The location of event



Figure 29. P-wave fault-plane solutions for selected events in the Dixie Valley area. Plots are lower-hemisphere, equal-angle projections, shaded sectors are compression. Orientation of T-axis indicated by heavy line on each mechanism.

50.

EVENT	"FAULT PLANE"		X-AXIS		P-A	P-AXIS		T-AXIS	
	STRIKE	DIP	TREND	PLUNGE	TREND	PLUNGE	TREND	PLUNGE	
Α	45	40SE	110	38	54	76	-58	05	
B	31	40SE	109	40	60	81	-66	05	
C	25	24SE	120	24	120	23	-60	20	
D	05	49SE	115	48	-150	80	105	03	
E	45	50SE	135	50	-45	85	146	05	
F	39	43NW	20	18	55	45	170	19	
G	45	53SE	-147	28	-95	49	-08	05	
H	24	74NW	-147	28	67	33	-30	10	
I	34	90					165	20	
J	32	90			75	14	-15	14	
K .	0	50W	-19	20	28	46	130	11	

TABLE 3. PARAMETERS OF FAULT-PLANE SOLUTIONS

Note: Strike and trend angles are taken to be positive in clockwise direction from north; all angles are in degrees.

D was obtained by fixing the focal depth, so the latter could not be used in determining average depth for the Type I events.

<u>Oblique-slip solutions</u>. For the six oblique-slip mechanisms (F-K), Table 3 lists the source parameters, assuming always that the N- or NE-striking plane is the fault. Note that the T-axes trend from 8 to 50 degrees in a NW-SE direction; the average trend is -21 + 16 degrees, and the average plunge is 13 degrees. Of the six solutions, I is the worst, with possible variations in the fault and auxiliary planes of +10 and +25 degrees, respectively; however, the extension direction for this solution can be varied by only +10 degrees. With the exception of solution K, all of the six oblique-slip mechanisms have a fault plane striking N24-45⁹E and dipping from 40 degrees to vertical; in three cases the W or NW block is downdropped, in one case the SE block is downdropped, and for two cases motion is right-lateral strike-slip.

Taken together, these observations can be explained by the known tectonic stress regime in the western Great Basin, together with a depth effect. Thus, T-axes for all of the eleven mechanisms are oriented NW-SE, in agreement with known regional extension of the province. The main difference between the two types of mechanism is that the maximum principal stress (P-axis) is almost horizontal for the strike-slip (Type II) sources and near-vertical for the dip-slip (Type I) events. This could be explained by depth differences, with weight of the overburden at larger depth leading to maximum stress in the vertical direction, and possibly some rotation of principal stress directions in the horizontal plane. Such an effect was suggested by McGarr (1980), who noted that experimental data indicate increasing maximum deviatoric stress with depth, and concluded that "a vertical profile of measurements from the surface downward might show a systematic rotation of the horizontal principal stress directions with the implication that stress trajectories measured at the surface may not be indicative of directions throughout much of the crustal section... The interesting aspect of the effect is that it is a systematic regional phenomenon, which is not related to any local source of stress inhomogeneity."

A ten-station seismic network was operated in and around the Dixie Valley area from January 1980 to November 1981; three of these stations are still in operation. Data from the Dixie Valley network were analyzed through 30 June 1981, and results of analysis were compared with analysis of somewhat larger events for the period 1970-1979. Conclusions and recommendations are the following:

1. Seismicity in the Dixie Valley-Fairview Peak area is almost exclusively confined to the rupture zone of large earthquakes that occurred there in 1954. During the entire 11.5-year period of observation only 11 events large enough to be analyzed were located in the northern part of Dixie Valley (north of Dixie Meadows). Five of these events -- with magnitude 0.1-1.6 -- were recorded by the Dixie Valley net in 1980 and 1981; the largest shock in this area was an ML 3.0 event in 1973. Thus, the area of most interest to geothermal exploration has been quiescent for at least a decade.

2. Earthquakes in northern Dixie Valley had focal depth of 7.2 km or less. This is quite shallow compared to mean depth of 11.3 km for the 1954 rupture zone to the south. A cross-section showing depth of five events in the area north of Dixie Meadows suggests that listric faulting may be occurring there, but much more data would need to be collected and analyzed to confirm this. Cross-sections for well-located events in the 1932 and 1954 rupture zones suggest that fault dips of 50-60 degrees persist to depths greater than 16 km, thus arguing against listric faulting in those areas.

3. Focal mechanisms for larger events in the area of interest are consistent with regional extension in the NW-SE direction, as found by a number of previous investigators. There appears to be a correlation between the ratio of dip-slip to strike-slip motion and focal depth, with deeper events having primarily normal slip on NW- or SE-dipping faults. This may be explained by a relatively larger increase in vertical than in horizontal stress with depth, due to the overburden (McGarr, 1980).

4. Perhaps unfortunately for this investigation, the level of seismicity in the Dixie Valley-Fairview Peak area was anomolously low during the period of operation of the Dixie Valley seismic network. In the southern part of the 1954 rupture zone, the average rate of occurrence of events with ML > 2 for the period after 1976 was almost five times less than that for the preceding seven years. This decrease in activity was generally synchronous with a period of quiescence reported by VanWormer and Ryall (1980) for the entire Sierra Nevada frontal fault zone. In the case of the Sierra Nevada, however, the quiescence ended with a burst of moderate earthquakes in all parts of the zone in late 1978 and 1979, culminating with several ML 6+ shocks in the Mammoth Lakes area in May, 1980 (Ryall and Ryall, 1981). In the Dixie Valley-Fairview Peak area, the period of quiescence has continued until the time of this writing.

5. A recurrence curve for events in the 1954 rupture zone for 1970-1981 indicates a b-value (related to the ratio of small-to-large earthquakes) that is high (1.3) for the western Great Basin. Such high b-values have been found to be characteristic of aftershock sequences, earthquake swarms and volcanic earthquakes (Mogi, 1966; Utsu, 1971; Hamilton et al, 1971; Wyss, 1973). However, a running b-value calculation indicates that b increased from 1970 to

1974, and has since decreased during a period of generally reduced seismicity. Ryall and Ryall (1981) observed a similar decrease of b in the southern Sierra Nevada from 1970 to 1979, followed by a rapid increase just preceding the large earthquakes at Mammoth Lakes in 1980.

6. Evidence related to the possibility of a large earthquake in northern Dixie Valley is ambiguous. Based on geologic observations, Wallace (1978; personal communication, 1981) considers the area between the 1915 and 1954 breaks to be a "seismic gap," with the potential for an earthquake of magnitude 7+ in the near future. In a way, the possibility of a large shock there is supported by the observation of very low seismicity in the area north of Dixie Meadows, and by a decrease in b-value starting in the mid-1970's and continuing to the present time.

On the other hand, some of the evidence argues against an impending large shock in this area. Earthquakes in northern Dixie Valley appear to be very shallow, and could possibly even be associated with listric faulting in a shallow crustal section overlying a zone in which deformation does not involve brittle fracturing. In a previous study, Richins (1974) noted that earthquakes in northwest Nevada, a region characterized by high heat flow and geothermal activity, tend to be shallower than those in the major earthquake zones of central Nevada. Richins concluded that the maximum magnitude of earthquakes in geothermally active regions may be only 5 3/4 - 6, as a result of the weakening of crustal rocks by fracturing in the vicinity of intrusive bodies, or by the effects of stress corrosion and leaching due to geothermal fluids. We note also that decreases in both seismicity and in b-value were observed in the 1954 rupture zone, and not in the northern Dixie Valley area. The possibility of another major earthquake at Fairview Peak or in the southern part of Dixie Valley, only 30 years after the last such event, would not be likely on either geologic or seismologic grounds (Wallace, 1977, 1978; Ryall, 1977).

7. Because of the ambiguity relative to the possibility of a future large earthquake in the area of primary interest to geothermal exploration, it is recommended that some level of monitoring be continued in the future. Probably the present level of effort, with two stations in northern Dixie Valley, will be sufficient for the time being. If activity increases in the future, if a rapid increase is observed in b-value for the Dixie Valley-Fairview Peak area, or if geothermal production is undertaken then this passive monitoring should be replaced by an expanded level of effort, and one involving a dense station network.

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APPENDIX A: FELT EARTHQUAKES IN THE DIXIE VALLEY REGION, 1840-1954.

The first seismograph with timing in California was installed at Berkeley in 1910. The first such instrument in Nevada began operation at the Mackay School of Mines in 1916. Prior to that time information on moderate to large earthquakes in the Nevada region is available only from earthquake catalogs (Holden, 1898; Townley and Allen, 1939; Slemmons et al, 1965) or from newspapers that were published at the time the earthquakes occurred. Since Nevada had a large number of newspapers in the late 1800's, and since these were fairly evenly distributed over most of the west-central part of the state, the newspapers offer a valuable resource in determining the extent of the felt area for a given earthquake, location of the epicenter, maximum intensity, etc.

For this investigation, eleven newspapers in northern Nevada were searched for information on all of the shocks reported in earthquake catalogs as having occurred in the region before 1917. Epicenters for these early shocks were estimated by comparing the intensities reported in the newspapers for various locations within the felt region. For example, an earthquake on 23 March 1872 was reported as "violent" in Austin, "slight" in Winnemucca, "frightened people" in Unionville, and "not felt" at Gold Hill and Eureka. By plotting the intensities corresponding to these effects on a map the earthquake was estimated to have occurred in northern Dixie Valley.

For many of the earthquakes, magnitude could be determined from the radius of the felt area, probably to within about half of one magnitude unit. The relationship between these two parameters was determined by Gutenberg and Richter (Richter, 1958), but it appears to be applicable to Nevada earthquakes as well. Again taking the 23 March 1872 shock as an example, the earthquake was felt to 120 km (Winnemucca) but not to 160 km (Eureka, Virginia City). Taking 150 km as the radius of the felt area, the graph on Figure A-1 gives ML 5.2 for this event; as a conservative estimate we have listed it as having ML 5 1/2.

Earthquakes in the region around Dixie Valley are listed below for the period from about 1845 to 1954. Sources of information are indicated by capital letters, according to the abbreviations following the earthquake listing. Times through 1915 are local times; time for 1916-1954 is GMT.

- About 1845(?). 39.5, 118.5, near Stillwater, intensity IX. S65: "RRR of 10-17-65 citing VCTE. Piutes report quake near Pyramid Lake 13 years ago. Great cracks opened from which water spouted 100 feet high. Large landslide on Slide Mountain." GHN, 12-30-69: At the time of the 1869 Wadsworth shock Captain Charley, a Washoe Piute aged about 30 years, reported that there had been a great earthquake when he was a little boy. The earthquake occurred while Indians from the Walker, Carson and Truckee Rivers were fishing at the Carson Sink. The shock knocked people down. River banks were shaken down in the vicinity of Stillwater, and the river T-A does not list any large earthquake in western changed its course. Nevada until 1860. Presumably, an event with the effects described above would have been felt at Sacramento or Auburn, and would have been reported in California newspapers had it occurred after 1850 as reported by S65.
- 1872 March 23, 1:41 pm. 39.8, 117.8, Dixie Valley, intensity VIII, M 5 1/2. RRR, 3-23-72: "Violent shock" in Austin, lasted about 5 seconds, plaster fell in courthouse. HR, 3-30-72: "Slight shock" in Winnemucca,

lasted only a few seconds. Also reported felt in Unionville, where people were frightened. ES, GHN: Not reported.

- 1872 May 6. 41.0, 117.5, near Winnemucca. S65 quotes RRR of 5-13-72: Several distinct shocks felt in Winnemucca in early part of week. HR, ES, GHN: Not reported. Probably Pleasant Valley.
- 1872 May 24. 41.0, 117.5, near Winnemucca. S65 quotes RRR of 5-25-72: Two slight shocks at Winnemucca. HR, ES, GHN: Not reported. Probably Pleasant Valley.
- 1873 November 7, 6:40 pm. 39.7, 118.1, Dixie Valley, intensity IV, M 5. RRR, 11-6-73 and 11-10-73; GHN, 11-6-73; HR, 11-14-73: First shock at 10 am on 11-5 felt plainly on Reese River but not in Austin. Shocks at 2:30 and 7:30 pm on 11-5 and 4:30 am on 11-6 were lightly felt in Austin and Virginia City. The last of these was the strongest. Largest event of the series was at 6:30 pm on 11-7. It was "plainly perceptible" in Austin, "causing windows to rattle and lamps to shake," and people ran outdoors. It was also reported felt in Unionville and Winnemucca.
- 1874 November 29. 40.7, 117.7, northern Pleasant Valley, intensity VI, M about 4 1/2. GHN, 12-1-74, quoting SS of 11-30: Shock occurred on Sunday morning, startling persons who were still asleep in Winnemucca, rattled crockery and cracked a partition in the Courthouse. Many frightened and ran outdoors. Also felt strongly at Fairbanks (10 mi "above" Winnemucca) and Battle Mountain. Not mentioned as felt in Virginia City. T-A: Two heavy shocks at Oreana. RRR: Not reported.
- 1882 October 12. 41.0, 117.5. SS, 10-17-82: Southeastern Humboldt County, along the Jersey Range. RRR, EDL, EI, VCTE, REG: Not reported.
- 1885 May 1, 9:30 pm. 40.9, 117.7, northern Pleasant Valley, intensity IV. T-A: IV at Winnemucca. EI, VCIE, REG, ES, EDL: Not reported.
- 1900 February 29, 1:35 pm. 40.9, 117.7, northern Pleasant Valley, intensity IV. T-A; SS, 2-28-1900; RRR, 3-3-1900; EI, 3-2-1900; ES, 3-3-1900: Shock felt in Winnemucca, startled people in Courthouse. Not reported felt elsewhere. TT, VCTE: Not reported.
- 1915 October 2, 10:55 pm. 40.3, 117.6, Pleasant Valley main shock, intensity X, M 7.6 (Richter, 1958). T-A; BKS; REG, 10-4 and 10-15-15; RRR, 10-9-15; SS, 10-5, 10-7, 10-9, 10-12-15; S65: Felt from Washington to the Mexican border and from the Pacific coast to Montana, Wyoming, Colorado, and Arizona -- an area of more than 500,000 sq mi. Damage greatest at Kennedy, where practically every building was destroyed. Faulting along the western face of the Sonoma Range had 6-12 feet of displacement for a distance of 20-25 miles. Damaged Southern Pacific water tanks at Battle Mountain, Golconda, Lovelock and Parran. In Winnemucca nearly every chimney was toppled, walls were cracked, and a few were thrown down. At Golconda a steel railroad bridge was warped. Stock was thrown from shelves in Elko. At Austin the main shock lasted 40-90 seconds, according to different stop watches; damage was inconsequential. At Tonopah and Goldfield there was "prolonged swinging motion" and clocks stopped. There were two strong foreshocks, at 4:41 pm and 5:50 pm; at Winnemucca 26 aftershocks were noted in the first nine hours after the main

shock. S65 lists 8 aftershocks during October, and T-A lists an additional 70 shocks during the next three months. BKS: Amplitude of 700 microns at Berkeley, more than 800 at Lick Observatory. This would correspond to M 8.0. SS, 10-5-15: "This was the first earthquake known here for 31 years, a slight shock having been felt in 1884. However, pioneers were informed by Indians, many years ago, that during the early 40's heavy shocks were of frequent occurrence for four or five years." This would appear to indicate a period of quiescence preceding the main shock, up to the two foreshocks on October 2.

- 1916 October 21 to December 27. 40.5,118.1, Unionville sequence. T-A lists five events felt at Unionville from October 21 to 31, nine shocks in November, and eleven from December 1-27. Sources given are George A. Bice, MWR, and Seismological Dispatches of Georgetown University. AEJ does not list any of these events as having been recorded at Reno, so they must have had M < 3.7.
- 1917 April 15, 19:01 GMT. 39.7, 118.1, Dixie Valley, M 4.8. T-A: intensity III at Fallon, Rochester and Lovelock. AEJ gives date as April 11, probably an error. RW readings give distance of 150 km from Reno, amplitude 12 mm. BKS: Recorded at Berkeley, Lick Observatory.
- 1917 June 1, 20:37 GMT. Pleasant Valley (?) T-A: Intensity III at Winnemucca. AEJ, BKS: Not reported.
- 1918 August 19, 10:53 GMT. Pleasant Valley (?) T-A: Intensity IV at Winnemucca. Rapid rocking, duration 3 seconds. BKS: Earthquake recorded at Berkeley at 11:56, which may be this event. AEJ: Not reported.
- 1930 September 16, 11:30 GMT. 40.4, 117.2, Buffalo Valley. USEQ: Felt at Red House, 7 mi south of Blossom Ranch, light. BKS, AEJ: Not reported.
- 1933 April 30, 16:17 GMT. 39.6, 118.1, Dixie Valley, M 4 1/2. AFJ: RW readings give distance of 140 km from Reno, amplitude 4.4 mm. BKS: Event recorded at Berkeley and Mount Hamilton. USEQ: Not reported.
- 1936 September 21, 07:32 GMT. 40.6, 118.4, west of Imlay, M about 4 1/2. AEJ: RW readings give distance of about 150 km from Reno, amplitude 2.2 mm. BKS: Recorded at Berkeley and Fresno at the same time, distance about 460 km. USEQ: Felt weakly at Winnemucca and Beowawe.
- 1936 September 22, 10:40 GMT. 40.5, 117.5, Pleasant Valley, M about 4 1/2. AEJ: RW readings give distance of 215 km from Reno, amplitude 3.3 mm. PAS: Readings give distances of 700 km from Pasadena, 380 km from Tinemaha.
- 1937 May 25, 05:35 GMT. 40.0, 117.8, Dixie Valley, M about 4 1/2. AEJ: RW readings give distance of 180 km from Reno, amplitude 2.7 mm. PAS: Not reported. BKS: Readings give distances of 440 km from Mount Hamilton, 390 km from Fresno. USEQ: Not reported.
- 1954 July 6, 11:13 GMT. 39.42, 118.53, east of Fallon, M 6.8. BKS gives location and magnitude. USEQ: Felt over an area of approximately 130,000 sq mi. Maximum intensity IX at main fracture zone on east edge of Rainbow Mountain. Maximum intensity VII elsewhere. At Fallon Naval Air Station

heavy steel lockers fell over, injuring several. Paved highways in Fallon and Stillwater areas settled, cracked and buckled in several places; one section south of Fallon settled 18 inches for a distance of 200 feet. Several old and poorly build structures in Fallon were considerably damaged. Length of faulting 11 miles, with scarps from a fraction of an inch to a foot or more. Extensive damage, estimated at \$200,000, to canal and drainage facilities of the Newlands Project, particularly in the Lone Tree and Stillwater areas. Coleman diversion dam failed. Geysers spouted from fields. In Fallon, considerable damage to wood, brick, masonry and concrete structures. Plaster, windows, walls, chimneys and ground cracked. Felt from Wendover to San Francisco, and Lakeview to Moapa.

- 1954 August 24, 05:51 GMT. 39.58, 118.45, east of Fallon, M 6.8. BKS gives location and magnitude. USEQ: Felt over an area of approximately 150,000 sq mi of Nevada, California, Oregon, Idaho and Utah. Maximum intensity of IX assigned to ground fracture area 15 miles east of Fallon. Intensity VIII at Fallon, Lovelock, Stillwater and Upsal Hogback. Estimated damage to Newlands Project facilities was \$91,000.
- 1954 December 16, 11:07 and 11:11 GMT. 39.32, 118.20, Fairview Peak and Dixie Valley, M 7.1 and 6.8. Location and magnitude by BKS. USEQ: Felt over an area of approximately 200,000 sq mi, from Weiser, Idaho to Los Angeles and from the coast to Salt Lake City. Maximum intensity of X assigned to the spectacular surface ruptures, which extended for 55 miles along the east side of Fairview Peak and the west side of Dixie Valley. Scarp heights up to 15 feet. In Dixie Valley all wells increased in flow, and water bubbled from the ground in spots. An adobe cellar, gasoline and water tanks and stone wall collapsed. Piano shoved across room, buffet up-ended, stove moved several feet, dressers toppled, etc. US Highway 50 cracked, with up to six feet of vertical movement. At Frenchman's Station, only negligible damage to buildings, but estimated loss of \$3,000 in liquor stock.

List of Abbreviations:

AEJ -- Jones, A. E., Reporting of Earthquakes at Reno, 1916-1951, Univ. Nev. Seismol. Lab., 1975.

BKS -- Univ. Calif., Bull. of the Seismographic Stations, 1911-.

EDL -- Eureka Daily Leader, 1878-1885.

- EFP Elko Free Press, 1883-.
- EI --- Elko Independent, 1869-.
- ES --- Eureka Sentinel, 1870-.
- GHN -- Gold Hill Daily News, 1863-1882.
- HR --- Humboldt Register (Unionville and Winnemucca), 1863-1876.

JCJ -- J. Claude Jones, Univ. of Nevada, diary.

LT --- Lovelock Tribune, 1898-1912.

MWR -- Monthly Weather Review, US Weather Service, 1915-.

PAS -- Calif. Inst. of Technology, Seismological Bulletin, 1936-.

REG -- Reno Evening Gazette, 1876-.

RRR -- Reese River Reveille (Austin), 1863-.

RW ---- Reno Wiechert, seismograph at Univ. of Nevada, 1916-1950.

S65 -- Slemmons, D. B., A. E. Jones and J. I. Gimlett, Catalog of Nevada Earthquakes, 1852-1960, Bull. Seism. Soc. Am., 55, 537-583, 1965.

SS ---- Silver State News (Winnemucca), 1875-1925.

T-A -- Townley, S. D. and M. W. Allen, Descriptive Catalog of Earthquakes of the Pacific Coast of the United States, 1769-1928, Bull. Seism. Soc. Am., 29, 1-297, 1939.

USCGS US Department of Commerce, Coast and Geodetic Survey.

USEQ - United States Earthquakes, USCGS, 1929-.

VCTE - Virginia City Territorial Enterprise, 1859-.

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APPENDIX R. LISTING OF ALL EARTHQUAKES IN THE DIXIE VALLEY AREA

The following list contains all earthquakes listed in available catalogs or located by the University of Nevada Seismological Laboratory, for the area bounded by latitude 38.8-40.5 deg N and longitude 117.3-118.7 deg W. For the period prior to 1970 locations and magnitudes are from Slemmons et al (1965) and from the University of California Bulletin of the Seismographic Stations (1911-1969). For 1970-1981 the list contains earthquake parameters based on analysis by the University of Nevada, using data from the Nevada seismic telemetry network. For 1980 and 1981 the Nevada network was supplemented by data from ten stations installed in and around Dixie Valley as part of this investigation.

Accuracy of epicenter locations in this list varies as a function of the density of station networks used in the analysis. For the pre-instrumental period prior to 1916, locations are based on newspaper reports and are therefore only approximate. From about 1930 to 1948 locations were based only on data from California seismic stations and are accurate to perhaps several tens of kilometers. From 1948 to 1969 locations were based primarily on California data but included readings from 1-3 stations in Nevada; epicenters for this period are probably accurate to within about 10 km. For the 1970's location accuracy for events in Dixie Valley was of the order of a few kilometers; for 1980 and 1981 locations should have been accurate to within a few hundred meters.

The format for this table is as indicated at the top of each page: From left to right the listing includes date (year, month, day), origin time (hours, minutes, seconds; local time prior to 1916 and GMT thereafter), latitude (deg N), longitude (deg W), standard deviation of the location (seconds), number of stations used in the analysis, magnitude, depth (kilometers), and quality. For most events prior to 1950 times are given only to the nearest minute. Standard deviations are available only for the period beginning in 1970, and focal depth was determined only for those periods when supplementary stations were operated in the Fairview Peak-Dixie Valley area, 1974-1981.

YEAR MO DAY	HRMINSEC	LAT	LON	RMS NSTA	MAG	DEPTH	Q
1868 11 17	1315	39.000	118.000				
	2141	40.000	117.500	1 A. A.	5.5		
1873 11 05	1700	40.000	118.000		5.5		
1873 11 06	1700	40.000	118.000				1. 2. 1. 1. 1. 1. 1. 1. 1. 1. 1. 1. 1. 1. 1. 1. 1. 1
1874 11 29		40.300	118.300		4.3		
10/4 11 22							
1893 08 30	18	39.100	118.100				
1903		39.500	118.100				
1914 04 25	1703	39.400	117.700				
1915 10 02	2341	40.500	117.500				
1915 10 03	0149	40.500	117.500		6.1		
1915 10 03	0653	40.500	117.500		7.8		
1915 10 03	0705	40.500	117.500				
1915 10 03	0718	40.500	117.500				
1915 10 03	0733	40.500	117.500				
1915 10 03	0 749	40.500	117.500				
	0045	40 500	117 500				
1915 10 03	0845	40.500	117.500 117.500			·	
1915 10 15	2022	40.500	117.500			• • •	÷.
1915 10 20	0234	40.500 40.500	117.500				
1915 10 23	0409	40.500	117.500			•	
1915 11 23	0333	40.500	117.300				
1915 12 18	09	40.500	117.500			:	
1916 10 21	0540	40.400	118.200				
1916 10 21	0545	40.400	118.200				
1916 10 21	1650	40.400	118.200				
1916 10 21	1910	40.400	118.200				
2720 20 2-							
1916 11 01	0430	40.400	118.200				
1916 11 02	1600	40.400					
1916 11 02	1630	40.400	118.200				
1916 11 02	1730	40.400	118.200				
1916 11 03	0520	40.400	118.200				
		40.400	118.200				
1916 11 03	1500	40.400	118.200				
1916 11 04	0215	40.400	118.200				
1916 11 04	2300	40.400	118.200				
1916 11 04	2350	40.400	118.200				
1916 11 13	1722	40.400	110.200				
1916 12 17	1445	40.400	118.200				
1916 12 19	05	40.400					
1916 12 24	1930	40.400					
1916 12 25	0140	40.400					
1916 12 25	0206	40.400					
1916 12 25	0219	40.400					
1916 12 25	1505	40.400					
1916 12 26	0200	40.400				`	
1916 12 26	0255	40.400				1.1	
1916 12 26		40.400	118.200)			

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YEAR	MO	DAY	HRMINSEC	LAT	LON	RMS	NSTA	MAG	DEPTH	Q
1916	12	27	0650	40.400	118.200		1			
1917			1900	40.000	118.000			5.1		
1918			0457	40.500	117.500			3.1		
					118.000			4.2		
1927			0010	39.200				4.2		
1931	06	06	0210	39.500	118.200					
1932	12	21	0530	38.800	118.000			;		
1932	12	21	061004	38.800	117.980			7.2		
1932			1036	38.800	118.000			4.9		
1932			2354	38.800	118.000					
1932			1241	38.800	118.000			5.0	• •	
1932	12	24	1241	30.000	110,000			5.0		
1932	12	25	0355	38.800	118.000			5.5		
1932			1835	38.800	118.000			4.5		
1932			0541	38.800	118.000			4.4		
1932			0308	38.800	118.000			4.6		
1932	12	29	0621	38.800	118.000			5.2		
1932	12	29	0638	38.800	118.000			5.0		
1932		· · · · ·	0646		118.000			5.0		
1932			0418	38.800	118.000	- (•	4.6		
1932			1604	38.800	118.000			4.6		
1 9 33	01	02	1545	38.800	118.000			4.5	•	
1933	01	02	1707	38.800	118.000			4.7		
1933			1036	38.800	117.900			3.9	/ * •	
1933			1306		117.800			5.1		
1933			1333		118.000			4.5		
								2		
1933	UT .	TT	1730	38.900	117.800			5.2		
1933	03	12	2045	38.800	117.600			5.0		
1933			161713	40.000	A CONTRACT OF			4.5		
1933			2233	39.000	117.800			4.0		
1933			2058	39.200	118.200	• 		4.6		
					117.600			5.5		
1933	TO	21	1059	38.900	117.000			5.5	4	÷ .
1934	09	27	0928	39.500	117.700		* 	4.0		
1936			0530	39.500	117.500			4.6		6 a - 11 a
1936			2244	39.100	117.900			4.5		
1936			1629	39.300	118.200			5.0		
					and a second			4.5		a de la composición d
1936	60	21	0732	40.300	117.400			4.0		
1936	09	22	1040	40.400	117.300	1		4.7		
1936			1504	39.300	117.500	÷)	4.1		
1939			2200	39.400	118.300	-	х. н 2	4.2		
1939			1027	39.800	117.700	$C = 1^{-1}$	·	4.2		× .
1939			054248	38.800	117.800			4.0		
1240	00	03	004240	30.000	TT1.000	• • •		4. U		
1944	01	30	101830	39.500	118.500			3.5		
1945			1858	39.000	118.000	м.,				
1946			2215	39.000	118.000					
1949			142516	38.800	118.600			3.7		n
										D
1949	12	28	115839	39.400	118,000			2.9		D

YEAR MO	DAY	HRMINSEC	LAT	LON	RMS NSTA MA	g depih	Q
1950 01	11	135136	39.100	117.400	2.	a	D
1950 01		230245	39, 200	117.700	3.		D
1950 01	•	220213	39.400	118.000	3.		D
1950 09		081246	39.500	117.500	4.		D
					***	5	Ų
1950 11	1/	034651	38.920	118.320	•		
1 95 1 09	12	141230	38.800	117.900	4.	1 .	D
1952 03		010042	38,800	117.400	4.		D
1952 11		090542	39.000	118.000	4.		D
1952 11		022952	39.100	117.700	4.		D
1952 11		043810	39.100	117.700	4.		D
1752 11	15	043010	39.100	11/./00		•	5
1952 11	18	040408	39.800	117.700	4.	6	D
1952 12	31	022050	38.870	118.130	3.	5	С
1953 06	01	031654	39.600	118.000	3.	3	D
1954 07		111320	39.420	118.530	6.		A
1954 07		111804	39.420	118.530	5.		A
1701 07	00	111007	077120	1101000		-	
1954 07	06	1127	39.420	118.530	4.	8 :	D
1954 07	06	1141	39.420	118.530	4.	5	D
1954 07		114900	39.420	118,530	5.		D
1954 07		125400	39.420	118.530	4.		D
1954 07		125630	39.420	118.530	4.		D
1701 01	00	120000	0,00,000			-	
1954 07	06	131511	39.420	118.530	5.	2	D
1954 07	06	133601	39.420	118.530	4.	5	D
1954 07	06	145515	39.420	118.530	4.	5	D
1954 07	06	155121	39.420	118.530	4.	4	D
1954 07		173825	39.420	118.530	4.		D
							_
1954 07		175737	39.420	118.530	4.		D
1954 07		175740	39.420	118.530	4.		D
1954 07		1 9 00	39.420	118.530	4.		Α
1954 07		220741	39.300	118.500	6.		A
1954 07	06	2253	39.300	118.500	4.	1	D
1954 07	<u> </u>	231158	39.300	118.500	4.	۰ · ۲	D
1954 07		235707	39.300	118.500	4.		
1954 07			39.300	118.500	4.		D D
		002258					
1954 07		012951	39.300	118.500	4.		D
1954 07	07	022245	39.300	118.500	4.	L	D
1 954 07	07	043358	39.300	118.500	4.	1	D
1954 07		061108	39.300	118.300	4.	6	D
1954 07		103132	39.300	118.300	4.		D
1954 07		105248	39.300	118.300	4.		D
1954 07		124053	39.300	118.500	4.		D
179 <u>-</u> 1 ()	~/					-	
1954 07	07	160224	39.300	118.500	4.		D
1954 07	07	180055	39.300	118.500	4.	3	D
1954 07	07	214747	39.300	118.500	4.	1	D
1954 07		235045	39.300	118.500	4.	3	D
1954 07		021356	39.420	118.530	4.		A
		= =					

YEAR	MO	DAY	HRMINSEC	LAT	LON	RMS NSTA	MAG	DEPTH	Q
1954	07	02	040819	39.300	118.500	•	4.5		Ъ
						. }			D
1954			065836		118.500		4.3		D
1954	-		125510	39.420			4.7		Α
1954	07	- 80	193157	39.420	118.530		5.3		Α
1954	07	09	085003	39.420	118.530	· • • · · ·	4.9	•	Α
1954	07	10	012220	39.300	118.500		4.6		D
1954	07	11	001802	39.300			4.0		D
1954	-		070400		118.500		4.6		D
1954			095812	39.300			4.6		
									D
1954	07	12	101706	39.300	118.500		4.5		D
1954	07	10	160525	39.300	110 500		F 1		~
							5.1		D
1954			015712	39.300		- •	3.9		D
1954			013134		118.500		3.9		D
1954	07	16	115520	39.300	118.500		4.0		D
1954	07	17	000216	39.300	118.500		4.2		D
	-					•			-
1954	07	17	015311	39.300	118.500		4.4		D
1954			014155	39.300			4.1		D
1954			121322	39.300			3.9		D
1954							8 2 1 1		
			152812	39.300	1		4.1		D
1954	07	20	175604	39.300	118.500		4.0	4	D
1954	07	22	063035	39.300	118.500		3.9		n
1954			204158						D
					118.500	1.8	3.6		D
1954			131628		118,500		3.9		D
1954			035638	39.300	118.500		3.6		D
1954	07	28	035803	39.300	118.500		3.7		D
							_	· . ·	
1954			231716	39.300	118.500		4.0		D
1954	07	30	020010	39.300	118,500		5.1		D
1954	07	31	1100	39.420	118.530				
1954	07	31	135435	39.300	118.500	1	4.3		D
1954			1515	39.420	118.530				2
101		3T	1313	JJ. 720	110.330		1.1		
1954	07	31	172414	39.300	118.500	a an	4.5		D
1954	07		173116	39.300	118.500		4.3		D
1954			101853	39.420	118.530		5.4		Ă
1954									
			212454	39.300	118.500		4.7	$\Delta Q = 2 \sqrt{2}$	D
1954	08	05	050308	39.420	118.530		4.7		A
1954	00	6	154116	39.300	118.500		4.2		D
1954									
			142834	39.300	118.500		4.0		D
1954			062310	39.300	118.500		4.0		D
1954	80	10	230051	39.300	118.500		4.0		D
1954	Ŏ8	10	230422	39.300	118.500	ne Regular (Maria)	4.0		D
1054	~~	10	00040	20. 200	110 500	an an Allandar An Allandar an Allandar		an a	_
1954			062942	39.300	118.500		4.1		D
1954			105008	39.300	118.500	이 같은 것을 못했다.	4.0		D
1954	80	23	1750	39.420	118.530				
1954	80	24	0445	39.420	118.530				
1954			055132	39.420	118.530		6.8		Α

•	YEAR	MO	DAY	HRMINSEC	LAT	LON	RMS NSTA MAG	DEPTH	Q
. •	1954	08	24	055746	39.500	118.500	5.2		D
	1954			061450	39.580	118.480	4.1		D
	1954			063231	39.580	118.480	4.4		D
	1954			063602	39.580	118.480	4.4		D
	1954			065410	39.580	118.480	4.3		D
	1904	08	24	005410	39.300	110.400	4.5) 	D
	1954	ng	24	144718	39.580	118.480	4.0	, ,	D
	1954			212053	39.580	118.480	4.4		D
	1954			021713	39.580	118.480	4.8		D
	1954			024959	39.580	118.480	4.0		D
	1954			121235	39.580	118.480	4.0		D
	1904	00	23	121255	39,000	110.400	4.0		D
	1954	08	25	222110	39.580	118.480	4.7	1 - 1 - 1 - 1 - 1 - 1 - 1 - 1 - 1 - 1 -	D
	1954			124415	39.580	118.480	4.2		D
				125615	39.580	118.480	4.6		D
	1954			070503	39.580	118.480	4.0		D
						118.480	4.0		D
	1954	08	28	045134	39.5 80	110.400	4.0	,	ע
	1954	08	29	030951	39.580	118.480	4.2	2	D
	1954				39.580	118.480	4.7		D
				035805	39.580	118.480	4.8	· · · ·	D
			_	0513	39.420	118.530	4.3		D
	1954			111558	39.580	118.480	4.]		D
•	1904	00	30	111330	39,300	110+400	4.1	. .	D
	1954	08	30	191157	39.580	118.480	4.]		D
	1954	08	30	195719	39.580	118.480	4.1	L A A	D
	1954		-		39.580	118.480	3.9		D
	1954			133407	39.580	118,480	4.0		D
	1954			221929	39.580	118.450	4.4	· · · · · · · · · · · · · · · · · · ·	D
	1/34	00	51	221 <i>727</i>	57.500	1101-100		•	2
	1954	08	31	222032	39.600	118.200	5.8	3	Α
	1954	08	31		39.420	118.530	3.9		
	1954			051846	39.600	118.200	5.5		D
	1954			112925	39.600	118.200	4.3		D
	1954			075317	39.600	118.200	3.9		D
	1774	09	02	073317	39.000	110.200			D
	1954	09	04	042432	39.600	118.200	4.()	Α
	1954	09	05	201559	39.600	118.200	4.4	Į,	D
	1954	09	08	071730	39.600	118.200	4.3	3	Α
	1954			092105	39.600	118.200	4.7		A
	1954			223138	39.600	118.200	4.4		D
		••		200200				-	_
	1954	09	14	161902	39.600	118.200	4.()	D
	1954	09	23	081219	39.600	118.200	3.8	3	D
	1954	10	80		39.600	118.200			
	1954			043212	39.500	118.500	4.3	3 (2012)	D
	1954			063040	39.500	118.000	4.0		D
								na La transforma	-
	1954			110711	39.280	118.120	7.3		A
	1954			111134	39.800	118.100	6.9		D
	1954			115036	39.500	118.000	5.0		D
-	1954	12	16	115730	39.500	118.000	5.(De la suite de la s	D
	1954			141657	39.500	118.000	5.8	3	B

YEAR	MO	DAY	HRMINSEC	LAT	LON	RMS	NSTA	MAG	DEPTH	Q
1954	12	16	14241	39.500	118.000		det e q	5.3		D
1954			150942	39.500			14. 17. j.	5.1		B
1954				39.500	118.000			- 1 - C - C - C - C - C - C - C - C - C	and the second	
					1			5.0		D
1954				39.500	118.000		•	4.7		D
1954	12	20	1100	39.300	118.200			1997 - 19		
1954	12	20	173647	40.000	118.000			5.0	• * * •	D
1955	01	01	121354	39.000	118,000			5.1		
1955	01	02	214336	39.000	118.000			3.7		
1955			220700	39.000	118.000	•		4.2		
1955			082040	39,000	118.000			4.2		
1955	01	06	083232	39.000	118.000	· ·		3.7		
1955										
			045610	39.000	118.000			3.9	• . •	
1955				39.000	118.000			4.1		
1955				39.000	118,000		1. 	4.2		
1955	01	08 .	084317	39.000	118.000		• •	3.6		• .
1955	01	08	180950	39.000	118.000			4.0		/
1955	01	80	223253	39.000	118,000			4.2		
1955	01	09	091050	39.000				5.0		C
1955			115840	39.000	118,000			4.2	1997 - 1997 - 1997 - 1997 - 1997 - 1997 - 1997 - 1997 - 1997 - 1997 - 1997 - 1997 - 1997 - 1997 - 1997 - 1997 -	•
1955			131554	39.900	118,400			4.1	14 juli	С
1733	UT.	10	101004	37.500	110,400			7.1		C
1955	01	11	102140	39.000	118.000			4.7		
1955	01	12	032125	39.000	118.000			4.0		
1955			110009	39.000	118.000			3.7		
1955				39.000	118.000			4.1		
1955			004550	39.000	118.000			3.9		
1933	U1	T. 4	004000	39.000	110.000			3.9		
1955	01	14	025704	39.000	118.000			3.9		
1955	01	14	122111	39.000	118.000	-		3.8		
1955			204702	39.000				4.2		
1955			181756	39.500	118.020			4.2		B
1955			014854		118.000			3.9		, D
1933	ÚT.	19 . ()	014034	39.000	110.000			J. 3		
1955	01	19	015348	39.000	118.000			4.1		
1955	01	19	021010	39.350	118.250			4.6		B
1955	01	19	0216	39.000	118.000			4.3		•
1955			032921	39.000	118.000			4.4		
1955			0329	39.000	118.000			4.5		
1955	01	22	193419	39.000	118.000			4.1		
									e e e L'Anna de	
1955			044834	39.000	118.000			3.8		
1955			132155	39.000	118.000			3.9		11
1955			153732	39.000	118.000			3.9		F
1955	01	25	232646	39.000	118,000			4.7		
1955	01	26	094021	39.000	118.000			4.1		•
1955	01	27	090422	39,100	118.070			3.8		B
1955			153805	39.800	118.000			4.2		C
1955			1730	39.350	118.250					_
1955			0635	39.200	118.500		-	1997 - 1997 - 1997 - 1997 - 1997 - 1997 - 1997 - 1997 - 1997 - 1997 - 1997 - 1997 - 1997 - 1997 - 1997 - 1997 - 1997 - 1997 -		С
1.00	<u> </u>			JJ 1 200					2 C	Č

YEAR	MO	DAY	HRMINSEC	LAT	LON	RMS NSTA 1	MAG	DEPTH	Q
1955	02	01	153116	39.180	118.130		4.0		C
1955			183031	39.200	118.530		3.8		B
1955			023043	39.050	118.050		3.6	1 - 1 2 - 1	C
1955			111122	39.100	118.130		3.7		B
1955			194408	39.030	118.180		3.7		В
1900	02	10	194408	39.030	110.100		J•7		
1955			161232	39.450	118.100		4.7		В
1955			1129	39.000	118.000		4.6	n an	_
1955			094024	38.900	118.180		3.7		B
1955			004640	39.100	118.050		4.1	· · · ·	B
1955	02	17	013620	39.170	118.020		3.8		B
1955	02	19	234920	39.300	117.800		4.1		С
1955	02	19	235007	39.300	117.800		4.8		C
1955	02	20	193159	39.030	118.030		3.8		В
1955	02	22	054017	39.050	118.170		3.8		C
1955			141116	39.700	118.100		3.6	· · ·	D
1955	03	08	200517	39.200	118.550		4.5		C
1955			232804	39.650	118.000		4.2		B
1955			142316	39.300	118.100		4.5		D
1955			084023	39.560	118.050		4.6		Ċ
1955			210827	39.160	118.120		4.0		В
1955	03	11	182347	3 9.4 00	118.250		4.7		C
1955			031540	39.550	118.000		4.4		B
1955		1	040616	39.500	118.200		4.4		B
			195439	39.120	118.170		3.8		Ċ
1955							4.2		В
1955	03	26	035128	39.330	118.000		4.2		D
1955	03	30	092428	39.100	118.150		4.3		C
1955			022251	39.450	118.000		4.2		С
1955			013602	39.500	118.000		4.1		С
1955			010737	39.000	118.180		3.9		C
1955			115008	39.300	118.000		3.8		D
1955	∩4	12	113257	39.550	118.100		4.4		В
1955		03	181701	39.200	118.000		4.0		Ĉ
1955		08	103833	38.930	118.000		4.5		В
1955			044406	39.200	118.200		4.0		D
					118.000		4.5		č
1955	05	30	212826	39.400	110.000		4.5		
1955	06	06	092010	39.150	118.150		4.3		C
1955		08	122211	38.880	118.170		4.5	· · ·	
1955		19	192000	38.970	118.250		5.2		B
1955			174821	39.180	118.130		4.3		B
1955			123713	39.450	118,100		4.1		С
1955	07	05	070703	39.300	118.500		3.1		C
1955			1015	39.300	118,500			ta di ta	
1955			052421	39.250	118.050		4.2	en e	C
1955			052421	39.800	118.000		3.9		C
				39.450	118.000		4.4		C
1955	09	10	193534	37.430	110.000		717		<u> </u>

YEAR	MO	DAY	HRMINSEC	LAT	LON	RMS NSTA	MAG	DEPTH	Q
1955	09	18	194402	39.42 0	118.000		3.9		с
1955			221151	39.600			4.1		c
1955			054051	39.200		· · · · · · · · · · · · · · · · · · ·	4.5		c
1955			071542	39.300	118.050				
1955			071542 081246				4.1		C
1900	10	23	001240	39.500	117.500		4.5		D
1955			061517	39.500	118.050		4.6		В
1955			202534	39.420	118.080		5.5		B
1955			204055	39.420	118.090		4.4		D
1955	11	23	042408	39.870	118.050		4.1		С
1955	11	25	163001	39.400	118.050		4.3		С
1955	11	25	182649	39.400	118.000		4.2		С
1955	12	01	102457	39.400			4.3		Ĉ
1955			153116	39.180			4.0		č
1955			004640	39.120			4.1		В
1955			181634	39.560	118.500			· · ·	Ċ
	12	12	101034	39.000	110.000		4.0		C
1955	12	22	120510	39.000	118.500		4.8		в
1955	12	22	120654	38.980	118.700	•	4.6		ē
1955			135104	39.000			4.5		B
1956			181756	39.500	118.030		4.2		B
1956			225530	39.330	118.050		4.2		B
			220000	07.000	110.000		702		D
1956			072619	39.030	118.070		4.6		В
1956	03	10	140756	39.320	118.470		4.2		В
1956	04	09	044226	39.180	118.120		4.3	- 	В
1956	04	23	150300	39.500	118.000		3.0		D
1956	04	29	073622	39.500	118.500		3.0		D
1956	05	<u></u>	065719	20 520	110 050				_
				39.530	118.050		3.0		C
1956			065719	39.530	118.050		3.0		С
1956			235107	39.450	118.270		3.8		С
1956			191936	39.400	118.020		3.9	الأربي بالعري	B
1956	06	23	104509	39.170	118.150		3.5		B
1956			021708	39.500	118.100		3.2	e e Al an	D
1956	07	04	043536	39.330	118.500	an an an Arran an Ar Arran an Arran an Arr	3.7		B
1956	07	21	100918	39.470	118.090		3.2		С
1956	07	26	095317	39.550	118.450		5.1		B
1956	08		010212	39.080	118,050		3.4		B
1956	08	10	134521	39.330	118.030		3.8		В
1956			091656	38.900	118.300		3.8		
1956			185542	39.530	118.450		3.5		D
1956			045308	39.120	118.450	(a) A set of the se			C
1956			141917				3.8		C
1900.			エオエフエノ	39.500	118.100		4.0		B
1956			110255	39.500	118.050		4.6		В
1956 I			202753	39.600	117.900		4.1		D
1956 1	L2 (03 🛛	144808	39.370	118.090		3.0		Ĉ
1956	L2 ()5	180515	39.130	118.070		4.1		B
1956			054723	39.200	118.120		3.9		č
						1. Sec. 1. Sec. 1.		1	<u>,</u>

YEAR MO DAY	HRMINSEC	LA.	LON	RMS NSTA MAG	DEPTH	Q
1956 12 07	042258	39.650	118.000	3.9		С
1956 12 07		39.500	118.000	3.6		D
1957 01 10	033754		118.070	3.7		C
	133552		118.090	3.8		C
1957 01 12				4.0		В
1957 01 13	193533	39.500	118.080	4.0		Ð
1957 03 27	132212	39.320	118.530	3.6		B
1957 04 25	064546	39.200	118,000	4.2		D
1957 04 26	180307		118.050	4.0		C
1957 04 27	053927	39.400	118.500	4.2		C
1957 04 29	140615	39.200	118.100	3.9	ан на сталия. Сталия сталия сталия Сталия сталия	D
1957 05 30	104819	39.600	118.200	3.8		D
1957 06 10	175922	39.500	118.600	3.0		С
1957 06 11	165755	39.200	118.300	4.2		D
1957 07 22	172401	39.560	118.400			В
1957 09 19	105013	39.300		3.7		D
1937 09 19	102012	39.300	110.100			
1957 10 17	101409	39.280	118.430	4.6		В
1957 10 25		39.300	118.200	3.9		D
1957 11 01		39.600	118.500	3.8		D
1958 01 18		39.100	118.100	4.4		C
1958 02 08		39.800	118,600	3.5		D
< c						
1958 02 16		39.400	118.600	3.4		D
1958 02 23	-022501	39.680	117.860	3.8		С
1958 03 01	112941	40.100	118.400	3.3		D
1958 03 05		39.000	118.200	3.8	}	D
1958 05 28		39.400	118.100	4.1	•	D
1958 06 01	164024	39.400	118.000	3.7	, ,	D
1958 06 01	165025	39.420		3.9		C
1958 06 11	060605		118.020			č
1958 06 29	083712	39.550				č
1958 06 29	074954	39.600	118.000	3.6		D
1958 07 04	074954	39.000	110.000	3.0		D
1958 07 08	125510	39.300	118.500	4.7		D
1958 08 10	202304	39.300	118.100	3.6		$^{\circ}$ D
1958 09 08	073111	39.000	118.200			D
1958 09 16	232637	39.600		3.5	5	D
1958 09 22	102042	39.300	118.100	3.5	5	D
1958 09 22	102540	39.300	118.100	3.7	7	D
1958 10 17	063153	39.200	118.100			D
1958 11 09	214005	39.300				D
1958 11 09	222326	39.500	118.100			D
1959 02 10		39.500				č
1939 02 10	T00343	JJ.J2U	110,100			Ť
1959 02 10	182536	39.500	118.100	3.		D
1959 02 11	024641	39.520	118.100			
1959 03 22	104728	39.300	117.900		Э се се се	D
1959 03 23	071020	39.600	118.070		3	В
1959 03 23	114923	39.550	118.100			C
1737 03 23		021000			1	-

YEAR	MO	DAY	HRMINSEC	LAT	LON	RMS NSTA	a mag	DEPTH	Q
1959	03	26	055541.0	39.700	118.100		3.7	en transforma Secondaria	D
1959	03	27	113033.0	39.600	118.100		3.9		D
1959	03	27	115948.0	39.650	118.100		4.2		C
1959					118.000		4.0		°C
1959	04	03	043222.0	39.600	118.100		3.1		D
1959	04	04	061122.0	39.500	118.100		3.7		D
1959							3.3		D
1959			-		118.100		4.3		C
			071104.0		118.080	n de se	3.4		C
1959	05	03	173809.0	39.000	118.020	ter a te	3.5	· · ·	С
1959			175140.0		118.000	in the second	4.8		С
1959			124859.0			${\cal L}_{\rm eff} = - {\cal L}_{\rm eff} + {\cal L}_{\rm e$	3.7	1 A.	D
1959			211429.0				3.3		D
1959			145647.0				3.2		D
1959	07	19	071210.0	39,400	118.000	•	3.8		D
1959			221644.0	39.500	118.000		4.4		С
1959			051320.0		118.000		4.0		D
1959			085248.0		117.900		3.2		D
1959			063518.0		118.100		3.1	. •	D
1959	12	24	112424.0	39.500	117.900		4.2		D
1960			213437.0	1 1 1 A 1 A 1 A 1 A 1 A 1 A 1 A 1 A 1 A	118.000		3.4		D
1960			044245.0		118.300		3.5		D
1960			042844.0				3.5		D
1960			065855.0				3.7		D
1 96 0	07	30	081310.0	40.000	117.670		3.6		C
1960	80	10	154351.0	39.600	117.900		4.0		D
			085545.0		118.050		3.5		C
1960	11	29	170833.0	39.900	117.800	en e			
19 61	07				118.600		5.0		
1961	08	04	165604.0	39.300	117.400		4.8		D
1961	08	04	165609.0	39.000	117.700		4.3		ч. Ч
1962	01	31	040730.0	39.500	117.500	an e de la	4.3		D
1962	07	20	090208.0	39.500	118.300		5.2		i Statistica
1962	12	16	110646.0	39.233	118.300		3.7		
1963	02	23	082239.7	39.523	117.957		3.7	e jak Nationalise	
					117.838		3.5		
1963	11	29	084856.0	39.560	117.930		3.2	•	e y
					118,800		3.5		•
1964					118.700		3.5		*
1964	03	22	181745.8	39.100	118.700		3.5		
1964	03				118.817		3.0	• • •	
1964					118.700	ter de la companya de	3.5		с.
1964			184320.4	38.927	118.727		3.7		
					118.700	ka shekara	3.2		
1964	07	02	132907.3	39.100	118.100		4.0		•
X									

YEAR	MO	DAY	HRMINSEC	LAT	LON	RMS NS	TA	MAG	DEPTH	Q
1964	07	04	072251.6	39.300	118.100			3.6		
1964	10	07	073707.1	39.200	118.800			3.4		
1965				40.292				3.9		
1965					118.650			3.2		
1965			155705.5	38,945				3.7		
1900	~	00	100.00.0	JU: 5 3 5	110.700			5.7		
1965			131422.1	38.900	117.700			4.6		
1965			065135.0	39.900	118,100			3.2		
1965	06	02	204704.0	38.900				3.8	1997 - 1998 1997 - 1998	· .
1965	06	25	001856.1	39.100	118.100			4.4		
1965	07	14	082643.0	39.600	117.800			4.3		
1965	11	01	171015.0	39.600	118.500			4.3		
1966			162022.0	39.300				3.2		
1966			072015.0	39,300				3.5		
1967			181545.0	39.300				3.7		
								3.5		
1 9 67	02	25	020032.8	39.300	117.600			3.0		
1967	03	05	213044.0	39.200	117.700			3.7		1.7
			171906.0	39.300				3.3		
			154508.0	38.953				3.1		
			111500.9	39.597				3.3		
1967			034434.5	39.135				3.1	•	
1907	11	07	034434+3	59.155	110.095			J.1		
1968	03	01	083314.9	39.650	118.433			3.4		
1968	03	27	194727.0	39.800	118.300			3.3		
1968	05	29	114107.1	39.067	118.050			4.9		
1968				39.700				3.3		
1968			173754.0	39.500				3.6		
1968	00	20	205043.0	39.000	118.000			3.3		
					117.750			3.4		
1968			094630.0							
1969			032455.0	39.600				3.2		
1969			105130.0	39.000				3.2		
1969	04	15	103000.5	38.850	117.950			3.2		
1 9 69	06	19	070233.0	40.067	118.617			3.3	4	
1970	1	7	1544 .5	39.344	118.130	0.11	5	2.4		
1 97 0	1	8	210027.5	39.315	118.458	0.09	7	2.2		
1970			1627 7.9	39.112	118.069	0.09	6	2.2		
1970		14	024526.9	39.028	118.170	0.13	7	2.3		
1070	1	14	052025 5	39.018	110 165	0 16	7	2 0		
1970		14	053925.5		118.165	0.16	7	3.0		
1970		15	105443.0	39.479	118.424	0.16	6	2.4		
1970		17	083916.8	39.391	118.090	0.08	7	3.3		
1970		20	060126.8	39.390	118.093	0.07	6	3.0		
1 97 0	1	21	134554.4	39.023	118.128	0.30	7	3.1		
197 0	1	21	171329.4	39.020	118.116	0.09	7	3.1		
1970			093243.0	39.072	118.154	0.10	6	2.5		
1970			061448.2	39.393	118.089	0.09	7	2.2		
1970			052557.0	39.246	118.517	0.07	5	2.3		
1970			114822.6	39.298	118.111	0.13	7	2.5		
1770	2	2			~_~****		-			·

YEAR	MO	DAY	HRMINSEC	LAT	LON	RMS NSTA	MAG	DEPIH	Q
1970 1970 1970 1970 1970	2 2	13 13 19 2 6	060939.7 083234.6 144018.0 192734.7 115551.8			0.10 6 0.01 5 0.10 6	2.2 2.6 2.1 2.8		
1970 1970 1970 1970 1970	3 3 3	7 7 11 11 13	194544.9 194551.2 1252 2.6 1302 7.7 015936.6	39.135 39.068	118.073 118.075 118.064 118.056 118.474	0.11 6 0.03 5 0.09 6	2.8 3.0 2.3 2.6 2.6		
1970 1970 1970 1970 1970	3 3 3	20 23 23 24 25	0152 9.9 151840.4 215540.1 051442.7 2053 1.6	39.578	118.094 118.075	0.06 6 0.10 6	2.5 2.2 4.2 3.4 2.8		
1970 1970 1970 1970 1970	3 3 3	27 27 28 31 10	013038.4 232046.9 085351.9 010225.4 125756.0	39.685 39.302	117.978 118.496 118.089		2.0 2.4 2.5		
1970 1970 1970 1970 1970	4 5	3	225258.0 214822.2 165713.2 214937.0 095335.3	39.459 39.042	118.402 118.087 117.747		2.7 2.6 2.7		
1970 1970 1970 1970 1970		8 25 31 6 8	135240.1 074122.2 132458.3 182817.4 170139.7	39.138 39.077 39.031	118.064 118.125 118.060	0.28 6			
1970 1970 1970 1970 1970	6 6	11 14 14 21 26	093158.8 074410.3 130038.1 170216.2 071628.2	39.712 39.061 38.808 39.722 39.320	118.077 118.124 117.966 118.010 118.435	0.09 5 0.67 7 0.12 4 0.28 5 0.27 8	1.8 1.7		
1970 1970 1970 1970 1970	7 7 7	30 3 8 18 27	091151.4 0104 1.7 070043.2 092934.6 084358.7	39.075 39.661 40.314 39.397 38.995	118.163 118.053 118.480 117.772 118.234	0.06 6 0.06 5 0.42 7 0.28 7 0.13 6	1.6 2.7 2.4 3.1		
1970 1970 1970 1970 1970	7 7 7 8 8	27 30 31 1 4	2329 1.0 075750.6 1009 .1 082445.2 1923 2.6	39.067 39.592 39.523 39.356 39.316	118.166 118.084 118.050 118.447 118.142	0.02 4 0.21 7 0.33 4 0.13 5 0.35 6	2.3 2.6 2.5 2.2 2.3		

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YEAR MO DAY	HRMINSEC	LAT	LON	RMS NST	a mag	DEPTH	Q
19708131970822197082719708271970831	092629.6 000544.7 092056.9 143611.2 212427.0	39.279	118.143	0.24 5 1.21 6 0.46 8 0.07 8 0.45 9	1.9 2.0 2.2		
1970 8 31 1970 9 4 1970 9 14 1970 9 19 1970 9 19	213428.5 052451.6 085752.5 230023.3 2308 1.9	38.874 39.091 39.064 39.909	118.037 118.052 118.103	0.14 5	2.1 2.1 3.5		
19709211970924197092419709291970930	2335 7.5 123452.0 181037.3 115215.4 1815 7.8	39.520 39.013 39.113	118.082 118.378 118.130 118.071 118.081	0.07 9 0.14 6 0.08 7 0.10 7 0.13 7	2.8 2.2		
19701011970101197010519701081970109	165042.2 165240.6 161724.0 071729.4 035926.8	39.713 40.175	118.015 118.061	0.76 4 0.17 7	2.2		
1970 10 12 1970 10 14 1970 10 15 1970 10 17 1970 10 18	093552.6 192331.7 162719.0 103052.8 044248.0	40.298 39.020 39.027 39.032 39.251		0.27 6 0.11 6 0.11 7 0.15 8 0.18 6	2.3 2.1 2.4		
1970 10 20 1970 10 20 1970 10 22 1970 10 22 1970 10 24	061928.2 185132.1 043820.8 071816.4 035747.7	39.010	117.968 118.131 118.189 118.126 118.172	0.09 7	2.4 2.8 2.7		
1970 10 26 1970 10 29 1970 10 29 1970 10 29 1970 10 29 1970 10 29	041829.9 024610.3 1531 3.5 153726.5 153931.9		118.076	0.16 5 0.08 7	2.4 2.7		
1970 10 30 1970 10 30 1970 11 6 1970 11 10 1970 11 13	000148.6 222626.9 030529.6 101429.5 171032.0	39.160	118.080 118.119		2.2		
19701118197011211970112919701211970124	112757.1 172330.0 015430.3 205029.2 023655.3	39.109 39.175 38.917 39.719 38.819	118.061 118.062 117.852 118.086 117.968	0.16 7 0.38 8 0.20 4 0.14 5 0.08 5	3.4 1.8 2.2		

YEAR	MO	DAY	HRMINSEC	LAT	LON	RIS NSTA	MAG	DEPTH	Q
1 97 0	12	14	102027.2	39.336	118.114	0.14 8	2.2		
			070447.3						
			132847.6						
			003053.8			0.19 11			
1971			095427.2			0.08 7			
			· · · · · · · · · · · · · · · · · · ·						
1971	1	7.	200818.0	39.361	118.337	0.32 8	2.5		
1971	1	17	1211 2.1	39.339	118.458	0.18 6	2.4		
1971	1	20	001859.0	39.358	118.461	0.13 5	2.0		
1971	1	26	002544.9	38.864	117.899	0.52 4	2.2		
1971	1	28	044035.9	40.293	117.836	0.16 5	1.4	2	
1971	1	28	140732.3	40.201	117.809	0.30 7	2.9		
1971		29	065731.6						
1971			024216.3					,	
1971		31	102652.9		118.084			•	
1971			164642.8		117.813				
	. —					0100 <u>0</u> 0		e in the second	
1971	1	31	203326.6	40.138	118.540	0.64 6	2.0		
1971			154554.3						
1971	2		230153.8						
1971		6							
1971	2	14	005946.9		118.492				
	-				-	353			
1971		15	115326.8						
1971			000246.6		118.693				
1971			104836.9					•	
1971			0847 5.1			0.17 7			
1971	3	9,	140151.5	39.046	118.179	0.07 7	2.2		
1971	3	13	061121.7	39,103	118,054	0.18 8	2.3		
1971			144552.7		118.066				
1971			1703 2.0						
1971			170241.9	· · · · · · · · · · · · · · · · · · ·	118,156				
1971			170330.4						
			$- \nabla f = - (1 + 1) \int_{0}^{\infty} dt = \int$						
1971			094550.9						
1971			170430.9						
1971			0731 6.4						
1971			025635.3						
1971	4	7	192440.6	39.046	118.191	0.13 7	2.6		
1971	4	15	054916.3	39.134	118.074	0.12 7	2.4		
1971			030327.6				1.6		
1971			125018.6					n a shekar in Birthean a	
1971			080044.2		118.126				
1971			002035.6		118.139				
107-	•	•							
1971			204619.8		118.429				
1971			205250.6		118.430		2.3	•	
1971	5				118.448		2.8		
1971	5		185234.4		118.142		1.9		
1971	5	2	185917.0	39.018	118.131	0.05 7	2.4		

YEAR MO DAY	HRMINSEC	LAT	LON	RMS NST	'A MAG	DEPTH Q
1971531971551971551971561971510	170620.2 0308 2.6 093331.0 0539 3.0 044116.2	39.289 39.132 39.486 39.022 39.079	118.111 118.080 118.453 118.132 118.111	0.18 7 0.04 6 0.07 4 0.11 6 0.06 6	1.7	
19715101971514197151419715151971517	105925.1 0818 4.3 1453 1.4 183752.1 153318.3	39.496 39.037 39.029 40.445 39.505	118.446 118.087 118.127 117.603 118.461	0.19 4 0.06 4 0.10 5 0.49 4 0.12 6	1.8 4.0 2.4	
19715181971518197152119715211971523	191924.8 192921.9 165237.2 201610.8 084259.7	39.977 39.525 39.591 39.193 39.601	117.708 117.693 118.082 118.059 118.022	0.30 6 0.10 6 0.10 5 0.14 8 0.02 4	3.0 2.4 2.2	
19715271971613197161819716271971629	101825.9 103152.0 021911.4 025330.0 075030.6	39.584 39.019 39.116 39.032 39.105	118.078 118.134 118.021 118.158 118.077	0.45 6 0.09 8 0.48 6 0.09 5 0.08 7	3 2.4 5 2.2 5 2.1	
1971 6 29 1971 7 2 1971 7 2 1971 7 2 1971 7 3 1971 7 5	144321.7 070845.7 180738.4 085137.0 0415 7.6	39.166 38.809 39.652 39.075 38.984	118.086 117.955 118.283 118.086 118.169	0.07 8 0.01 4 0.18 6 0.13 5 0.08 5	1.9 1.9 2.0	
19717519717619717819717101971710	0626 9.5 0317 4.8 133610.6 072145.0 074657.0	39.703 38.983 39.708 39.196 39.105	118.370 118.177 118.350 118.169 118.081	0.23 7 0.16 5 0.23 6 0.26 5 0.04 4	2.3 5 2.5	
19717101971710197171019717101971710	150750.8 150816.5 150840.3 152944.7 153228.4	39.692 39.689 39.681 39.682 39.680	118.034 118.041 118.062 118.028 118.062	0.16 7	2.1 2.7	
19717111971711197171119717131971713	064941.8 065422.0 075441.0 0935 .1 132443.3	39.681 39.683 39.690 39.050 39.683	118.040 118.045 118.048 117.364 118.061	0.21 7 0.18 8 0.11 5 0.09 4 0.13 5	2.8 2.0 2.3	
19717201971720197172219717221971722	172812.6 172940.8 110149.5 153328.1 1604 8.8	39.700 39.712 39.233 39.172 39.171	118.045 117.987 118.502 118.079 118.082	0.70 4 0.89 4 0.11 5 0.06 5 0.01 5	2.1	

YEAR	MO	DAY	HRMINSEC	LAT	LON	RMS NSTA	MAG	DEPTH	Q
1971 1971 1971 1971	7 7	23 30 30 20	051232.9 123034.6 140757.7 073930.8	39.104 39.295 39.536 39.710	118.060 118.497 118.469 118.081	0.15 5 0.16 6	2.3 2.3 2.4 1.6		
1971		21	104938.8	38.864	117.765		1.6		
1971 1971 1971 1971 1971	8	31 6	134533.8 214841.9 122038.0 223322.4 072655.1	39.250 39.409 39.116 39.347 39.055	118.095 118.064 118.049 118.106 118.179	0.17 8 0.23 6 0.02 6 0.04 4 0.16 5	2.5 2.0 1.9		
1971 1971 1971 1971 1971 1971	9 9 10	14 19 22 2 2	003652.9 203626.7 110647.6 045748.1 050647.1	39.740 38.874 39.051 39.072 39.076	117.984 118.520 118.095 118.117 118.196	0.33 6 0.12 4 0.01 4 0.27 9 0.06 5	1.9	• • •	
1971 1971 1971 1971 1971	10 10 10	2 5 6 8	085445.6 085235.2 085354.4 1731 2.6 092923.6	39.065 39.233 39.241 39.116 39.093		0.06 8 0.04 5 0.15 5 0.78 5 0.09 5	2.0 1.4		
1971 1971 1971 1971 1971 1971	10 10 10	9 17	1613 8.4 142623.6 1839 9.9 0344 3.6 143428.5	39.050 39.118 38.997 38.991 39.056		0.18 5 0.29 6 0.53 6 0.13 5 0.10 7	2.4 1.8 1.7		
1971 1971 1971 1971 1971	10 11 11			39.088 39.087 39.494 39.354 39.355	118.065 118.076 118.427 118.088 118.093	0.18 5 0.10 6 0.26 7 0.11 9 0.11 9	1.6 2.5		
1971 1971 1971 1971 1971	11 11 11	5 5 5 12	053458.5 120458.7 123622.1 124419.9 102736.3	39.341 40.417 40.342 40.352 39.563		0.24 6 0.75 4 0.22 4 0.19 4 0.16 5	1.5 2.7 1.2 1.6 2.0		
1971 1971 1971 1971 1971 1971	11 12 12		081813.6 032038.2	39.392 39.311 39.235 39.664 38.890	118.068 118.490 118.112 118.083 117.865	0.21 5 0.27 4 0.23 7			· · ·
1971 1971 1971 1971 1971	12 12 12	10 11	184331.0 014859.6 024527.7 103518.7 1453 7.6	39.089 39.117 39.093 39.000 40.481	118.103	0.08 5 0.20 5 0.10 7 0.19 5 0.12 5	1.7 2.3 2.6 2.2 2.8		

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,	YEAR	MO	DAY	HRMINSEC	LAT	LON	RMS NSTA	MAG	DEPTH	Q
	1971 1971 1971 1971 1971 1972	12 12 12	19 28 29	210524.7 041929.7 163949.0 204137.6 201028.4	39.413 40.499	118.057 118.053 117.653 118.120 117.845	0.08 6	2.5 2.2 2.3 2.3 2.2		
	1972 1972 1972 1972 1972 1972	1 1 1 1 1	7 7 8 8	0158 1.8 093149.0 111317.9 023638.4 053149.7	38.992 38.989 39.005	118.081 118.363 118.369 118.373 118.304	0.25 8 0.19 7 0.14 7 0.18 6 0.65 7	1.4 1.9 2.3 1.8 2.1		
	1972 1972 1972 1972 1972 1972	1 1 1	19 19 19 20 20	045348.2 045552.8 073912.1 0126 4.0 0909 9.2	39.450	117.958 118.027 118.047 118.338 118.088	0.07 7 0.28 6 0.14 7	2.3 1.5		
	1972 1972 1972 1972 1972 1972	1 1 1	21 23 25 26 31	002447.9 191329.4 133317.3 140921.4 182312.9	39.093 39.200	117.866 118.111 118.051 118.070 118.112	0.32 7 0.08 6	2.2 1.4 1.5		
	1972 1972 1972 1972 1972 1972	2 2 2 2 2 2	9	040714.1 063318.7 200841.6 112034.9 144549.3		118.147 118.453 117.837 118.401 118.088	0.46 7	1.3 2.2 1.8 1.8 2.2		
	1972 1972 1972 1972 1972 1972	2 2 2	11 12 12 14 20	055932.3 061121.5 061142.1 115322.9 131422.1	39.098 39.109 39.748	118.063 118.055 118.084 118.161 118.086	0.26 8 0.10 6 0.50 8	2.5 4.0 1.6	· · · · ·	
	1972 1972 1972 1972 1972 1972	2 2 2		131554.8 1549 2.6 182726.9 023658.6 054728.2	39.134 40.179	117.770	0.24 6 0.20 6	1.5 4.3 1.7 2.4		
	1972 1972 1972 1972 1972 1972		29 29 2 2 2 6	224054.3 2303 7.9 162954.7 203940.0 131531.1	39.686 39.186	118.416 118.442 118.066 118.074 118.116	0.24 9 0.33 8 0.12 7 0.14 8 0.58 10	2.5 2.2 1.5 1.8 2.6		
	1972 1972 1972 1972 1972	3 3 3 3 3 3	6 6 7 7 7	180710.8 203551.8 045024.5 1110 7.1 210940.5	39.126 39.233	118.087 118.084 118.101 118.071 118.044	0.33 6 0.16 6 0.18 9	1.0 1.4 2.1 1.7 2.2		

	YEAR	MO	DAY	HRMINSEC	LAT	LON	RMS NS	TA	MAG	DEPTH	Q
	1972	3	8	183023.9	39.712	118.234	0.17	5	1.1		
	1972			190723.7							
	1972			162736.2							
	1972			030326.3							
	1972	3	11	063327.1	39.860	117.447	0.19	5	2.0		
	1972	- 3	12	043056.2	39,054	118,151	0.42	9	2.0		
			17	184021.8							
	1972										
				205145.9							
									1.4		
	1972	3	20	160132.1	39.090	118.056	0.07	4			
	1972	3	22	213836.1	38.994	118.159	0.29	9			
	1972	4	1	001726.2	38.912	117.892	1.02	7	2.3		
	1972			033914.4		118.076			1.8		
	1972			102361.1		118.083			1.9		
	1972	4		002554.7		117.933		7	1.9		
	19/2		5	002334.1	20,910	117.933	0.90	/	1.9		
	1972	4	6	003758.8	39.247	118.116	0.48	6	1.9		
	1972	4	6	062722.9							
	1972			0638 4.0							
	1972			102756.4							
	1972	4	6	1132 9.5	40.335	118.499	0.16	6	1.7		
	1972	4	7	2341 6.9	39.121	118.083		6	2.1		
	1972			061950.4	39.592	118.086		7	1.9		
	1972			201646.3		118.098			2.2		
	1972		21	200726.8					1.9		
	1972		23	002430.0		118.016	0.08		1.6		
	19/2	4	23	002430.0	39.913	110.010	0.00*	4	1.0		
	1972	4	24	2120 7.9	39.924	118.411	0.54	6	1.5		
	1972				39,920	118.395	0.38		1.4		
	1972	5	2	184947.7					1.4		
	1972	5		190959.2					1.3		
	1972			085939.1							
	19/2	5	5	003333.1	39.000	110.204	0.13	5	1.5		
	1972	5	5	162254.0	39.228	118.030	0.26	8	2.3	ж. До	
	1972		8	232029.9		117.861				•• * •	
	1972		11	040628.3		118.052			1.6	••	
÷	1972		11	185924.5					1.9		
	1972		12	110230.4		118.153				·	
	19/2	5	12	110230.4	39.575	110,133	U.22 I	T	1.9		
	1972	5	13	083031.9		118.107		8	1.8		
	1972	5	13	230515.2	40.275	118.499	0.27	5	1.8		
	1972		14	1323 6.4		118.308	the second s	~	2.0		
	1972		17	011831.8		118.534	0.12				
	1972		17	040422.4						· · · · ·	
	1712	5	+1	V7V722•¶	JJ•HII	110.090	0.13	<i>.</i>	£•£		
	1972	-5	17	095339.8	39.582	118.124	0.30 1	2	2.4		
	1972	5	18	113412.3	39.228	118.086	0.54	7	2.0		
	1972		22	200730.4	38.915			5	2.0		
	1972		24	1930 3.6		118.071		7	2.1		
	1972		25	083650.3		118.073			2.2		
	1916	5	23		37.107	110.073	0.17	0	£ • £		

YEAR	MO	DAY	HRMINSEC	LAT	LON	RMS NSTA	MAG	DEPTH	Q
1972	5	25	230840.7	39.105	118.041	0.22 5	2.1		
1972		27	160636.6	39.152	117.993	0.22 7	2.3		
1972		30	032614.9	38.807	117.951	0.56 7	2.5		
1972		30	1910 1.2	38.898	117.868	0.54 7	2.3		
1972	5	31	164438.0	39.254	117.894	1.20 5	2.2		
1972	6	1	091842.9	40.157	118.302	0.49 7	2.6		
1972	6	1	232041.3	38.911	117.874	0.24 5	2.2		
1972	б	2	114613.5	39.151	117.999	0.33 10	4.1		
1972	6	6	172418.0		118.066	0.23 9	2.1		
1972	6	6	195210.2	39.477	118.040	0.04 6			
1 97 2		16	231913.2	38.968	118.080	0.21 5	2.2		
1972	6	21	083942.8	40.116	118.293	0.70 6	2.3		
1972	6	21	085336.1	40.107	118.272	0.32 6	2.2		
1972	6	21	2349 2.0	39.453	118.118	0.34 6	1.9		
1972	6	22	035346.4	39.691	118.016	0.33 8	2.6		
									271
1972	6	22	051746.2	39.734	118.030	0.35 4	2.1	1.00	
1972	6	22	070859.0	39.690	118.009	0.26 6	2.1		1
1972	6	29	230437.8	39.715	118.064	0.19 9	2.3		
1972	7	4	010347.6	38.995	117.419	0.87 6			
1972	7	4	065535.0	39.090	118.024	0.27 12	1.5		
1972	7	4	232130.6	38.994	118.205	0.25 12	2.2		
1972	7	5	014913.2	40.179	118.461	0.63 7	1.9		
1972	7	6	232252.8	38.918	117.893	0.70 9	2.8		
1972	7	8	181113.3	39.134	118.031	0.17 7	2.8		
1972	7	9	2326 5.9	39.189	118.108	0.44 9	1.7		
	_								
1972		12	0123 5.8	39.283	117.691	0.31 14	3.3		
1972		12	021645.0	39.283	117.690	0.11 8	1.8		
1972		12	151728.8	39.287	117.689	0.20 11	2.5		
1972		18	131334.6	39.159	118.183	0.61 7	2.4		
1972	7	18	195956.9	39.327	118.120	0.16 9	2.8		
1 9 72	7	18	200136.3	39.313	118.091	0.20 7	1.9		
1972	7	23	223719.6	39.496	118.467	0.28 12	2.8		
1972	7	29	101331.4	39.133	118.108	1.68 14	2.5		
1972	7	29	141420.3	39.007	117.580	0.16 6	2.2		
1972	7	29	1 95256. 0	39.446	118.609	0.22 4	1.9		
1972	8	4	185538.1	39.305	118.069	0.16 6			
1972	8	4	2320 4.4		117.917		2.3		
1972	8		084644.7		118.088	0.46 8	1.6		
1972	8	8	181012.8		118.068	0.53 4			
							2.1		
1972	8	10	043548.9	39.722	118.045	0.31 7	2.7		
1972	8	13	090742.8	39.014	117.624	0.59 13	2.2		. 1
1972	8	19	033214.0	39.085	118.053	0.19 9	2.7		
1972		21	0755 3.3	39.577	118.149	0.15 6	2.0		
1972		28	122535.6		117.570		1.8		
1972		28	123416.7		117.564		1.6		
	0								

	YEAR	MO	DAY	HRMINSEC	LAT	LON	RMS NSTA	MAG	DEPTH	Q
	1972	9	2	075840.1	38,974	118.384	0.20 12	1.5		
	1972	9		0759 1.4						
	1972	ģ		004138.3				2.8		
	1972	ģ	-	011146.9						
	1972	ģ		190342.8						
	1912		0	190342.0		11/.029	0.43 /	Z•1		
	1972			181415.4						
	1972			115132.1						
	1972			235737.5			0.54 8			
	1972			234245.5						
	1972	9	18	101966.9	40,444	117.371	1.29 8	1.4		
	1 97 2			115352.9						
	1972		27	053051.3						
	1972	9	29	235225.8						
	1972	9	30	230832.9	39.394	118.145	0.43 11	1.9		
	1972	10	1	1124 1.1	40.291	117.734	0.39 12	2.3		
	1972	10	1	1133 4.8	40.392	117.861	0.47 6	1.8		
				024017.2						
				202819.8						
				194511.8						
				075043.2						
	1972	11	12	2045 6.7	30 084	118 077	0 44 10	् २ ६		
	1972			0106 7.9						
	1972			102159.1						
	1972			120026.1						
	1972			2353 9.4			0.19 10			
	1912	ΤT	22	2555 9.4	39.111	110.034	0.20 /	_ Z• 1		
	1972	11	23	032453.9	39.105	118.067	0.34 10	2.2		
	1972	11	23	130143.4	39.090	118.035	0.32 7	1.8	· .	
	1972	11	23	151017.1	39.111	118.033	0.16 7	2.2		
	1972	11	25	1453 8.9	39.108	118.049	0.29 9	2.1		
	1972	11	26	223562.3	39.082	118.051	0.49 8			
	1972	11	29	200114.5	38.914	117.938	0.88 6		n ing ing ing ing ing ing ing ing ing in	r i V V
÷	1972		1	081637.2			0.37 13			
• •	1972			003719.6			0.22 7			
	1972		6	010116.6						
	1972		8	0715 6.0				2.8		
	1. 1. A.							1		
	1972			091826.2						
	1972			043514.0						{ · ·
	1972			094125.2						1
	1972			104228.5						
	1972	12	13	003330.7	38.896	117.841	0.57 5			·
	1972	12	28	062815.5	39.294	118.527	0.13 6			
	1973	1	1	003831.7				2.4		
	1973	1	6	002835.3	38.901	117.828	0.44 7	1.9	*	
	1973	1	9	171120.4	39.344					
	1973	1	13	010230.4						
	1									

YEAR	MO	DAY	HRMINSEC	LAT	LON	RMS NSTA	MAG	DEPTH Q
1973	1	13	2002 3.5	39.100	118.073	0.13 8	2.1	14. State
1973	1	15	160912.2	39.200	118.016	0.43 6	1.5	
1973			0416 9.9	39.340	118.063	0.22 9	2.0	
1973		25		39.086		0.18 6	1.8	
1973		27	112119.1	39.052	118.169	0.31 12	2.7	
1975	-	21	11211701	57.052	110.107	0.31 12	2	·····
1 9 73		28	2255 5.9	39.166	118.032	0.23 7		• •
1 97 3		29		39.893	117.828	0.20 7	2.6	
1 97 3	1	31	002021.9	38.898	117.874	0.55 7	2.3	
1 97 3	2	4	043040.3	39.305	118.111	0.19 9	2.8	
1973	2	10	194030.4	39.137	118.081	0.12 7	1.8	
197 3	2	10	210748.6	39.404	118.096	0.23 10	2.8	
1973			212835.1	39.406		0.21 10	3.0	
1973		10	2308 2.1	39.154		0.17 8	2.0	
1973		11	000248.0	39.405		0.24 8	2.6	
1973	2	15	043931.8	39.020	118.453	0.11 9	2.1	
1973	2	15	112917.3	40.389	117.578	0.18 6	2.4	
1973	2	19	011138.6	39.360	118.001	0.19 8	2.1	
1973		19	100552.9		118.055	0.27 7	2.0	
1973		19	104050.8	39.144		0.11 7	2.0	
1973		22	0232 6.4	39.475	118.425	0.27 8	2.0	s - Arres
1072	n	24	002425.7	38.917	117.899	0.82 6	2.1	
1973								
1973		10	064920.0	39.016	118.214	0.17 5	1.6	
1973		10	101241.6	39.277	118.084	0.15 8	1.8	
1973		11	222151.4	39.215	118.017	0.36 7	1.8	
1 9 73	3	12	171412.1	39.031	118.158	0.33 8	1.8	
1973	3	16	145928.0	39.183	118.053	0.34 8	3.4	
1973	3	18	075924.2	39.242	118.029	0.27 6	1.8	
1973	3	18	201231.8	39.146	117.985	0.34 8	2.2	
1973			003115.3	38.913	117.895	0.81 6	2.3	
1973	-		151430.2		118.207			
	-	~ ~	051406 5	00 00 F		0 4C 7		
1973		24	051436.7		118.137	0.46 7		· · · · · · · · · · · · · · · · · · ·
1973			201840.3		118.093		2.6	
1973					118.062	0.20 9	2.6	
1973	3	29	212849.1	39.123		0.15 6	2.5	
1973	4	1	002344.3	38.898	117.827	0.46 6	2.4	
1973	4	2	054838.6	39.016	117.738	0.26 8	2.3	
1973			0757 9.3		118.165		1.5	
1973			065514.8	39.315	118.138	1.08 8	1.9	
1973			205829.9			0.27 8		
			1 A A A A A A A A A A A A A A A A A A A		118.186	0.27 8		
1973	4	8	0529 1.0	39.603	110.190	U.ZI /	1.6	
1973							1.7	
1973		12	010237.2		117.936	0.26 6		-22 -32
1973	4	14	123534.3	39.078	118.109	0.33 8	2.1	
1973	4	17	002522.5	38.911	117.883	0.62 7	1.9	
1973		21	002237.0	38.900	117.850			
	-		· · · · · · · · · · · · · · · · · · ·					

YEAR	MO	DAY	HRMINSEC	LAT	LON	RMS N	ista mag	DEPTH	Q
1 97 3		21	122021.3	40.445			6 2.8		
1973		22	031028.7			0.33	7 1.8		÷
1973	4	22	114658.7	39.195	118.027	0.38	6 2.5	i	
1973	4	22	114816.4	39.196	118.037	0.20	5 2.0) · · · ·	
1973		25	002345.2	38.869		0.70	6		
	-	-7		00.007	11/0/01	0.70	. •		
1973	4	27	0357 1.2	39.633	117.915	0.26	6		
1973	5	1	2324 3.2						
					117.881		6 2.1		
1973	5	5	200316.4				8 2.1		
1973	5	7	014618.6		118.057				
1973	5	, 7	083139.9	39.340	118.141	0.16	9 2.5		
	_	_		· · · · · · · · · · · · · · · · · · ·		1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1			
1973	. 5		152546.8		118.058		10 2.2		
1973		11	114656.6	39.323	117.570	0.29	6 2.2		
1973	5	13	180151.6	39.307	118.502	0.33	8 2.0		~
1973	5	16	120912.6	39.095					
1973		18	072057.9		118.051		6 1.8		1
	Ţ			031200	1101001	0.00	U 1.U		
1973	5	21	2321 2.5	38.894	117.822	0.47	6 2.3		
1973		22	150316.4		118.132		7 2.6		
1973		26	1957 8.4	39.023					
1973							6		
		27	140531.1	38.857					
1973	5	27	172836.2	38.870	118.115	0.20	9 2.1		
1973	5	28	090824.7	39.136	118.092	0.31	7.0		
1973	6	7	072829.3	39.064			9 2.6		
1973	6	8	033210.2		118.123				
1973		9					4 . T		
	6		0603 4.7	39.040			8 1.8		
1973	6	17	164418.5	39.497	118.425	0.32	5 1.7		•
1072	~	10	1010 1 0	20.002	110 064				
1973		18	1319 1.3		118.064	0.29	8 2.5		
1973		19	2305 9.5	39.040			7 2.2		
1973		20	232521.3	38.898					
1973		24	203659.0	39.450	118.093	0.16	6 2.0		
1973	6	26	204228.2	39.680	118.432	0.22	5 1.6		
1973	6	29	052530.3	39.423	118.413	0.28	5 2.1		
1973	6	29	234250.6	39.316	118.111	0.39	7 2.7	and the second	
1973		30	012812.1	39.323	118.117		5 2.2		÷ .
1973		30	175413.7		118.217	0.10	5		
1973		30	233019.7	38.893	117.819		7 2.4		
	. .		20001717	50.075	11/0013	0.52	1 2.17		
1973	7	1	102853.3	39.088	118.152	0.27	6		
1973	7		0652 3.2	39.833		0.23			- Geb
1973	7		084619.6	39.582	118.094		5 2.0		
1973	7	5	221738.3	39.194	118.088		· · · ·		ر. مرجع د
1973	7								
12/3	1	. 0	015354.8	40.086	118.074	0.42	6 2.7		
1973	7	6	174923.6	38.935	117.915	0.89	7 2.1		
1973	7		051014.1	39.745	118.032	0.24	8 2.6		
1973			232518.7	38.892	117.816	0.49	5 1.9		
1973			020455.4						
				39.263	118.122	0.09	4 2.4		
1973	1	15	094624.9	39.040	118.110	0.30	5 1.9		•

YEAR MO DAY	HRMINSEC	LAT	LON	RMS NSTA	MAG	DEPTH Q
19737151973718197372519737291973731	232210.5 232228.9 231818.3 024610.2 232426.6	39.029 39.209 39.118 39.065 38.890	118.071 118.000 118.075 118.136 117.835	0.20 8 0.36 8 0.22 5 0.29 7 0.46 6	1.8 2.2 2.1 2.0 2.6	
1973 8 2 1973 8 3 1973 8 5 1973 8 6 1973 8 9	172129.6 070234.5 024539.9 035138.0 010344.7	39.211 39.325 39.041 39.213 39.066	118.032 118.083 118.146 118.112 118.034	0.25 6 0.09 7 0.12 8 0.02 5 0.26 5	2.0 2.0 2.0 2.1 1.8	
19738101973810197381119738121973816	215329.2 232417.2 065518.7 095950.0 020254.4	39.377 38.900 39.036 39.289 39.204	118.128 117.843 118.188 118.091 118.028	0.09 5 0.45 5 0.29 7 0.10 7 0.16 6	1.8 2.3 1.9 1.9 2.0	
1973 8 19 1973 8 27 1973 8 31 1973 9 8 1973 9 11	022628.8 062857.5 2325 7.2 0247 5.1 2323 2.4	38.887 39.051 38.898 39.430 38.907	117.718 118.115 117.855 118.490 117.869	0.40 7 0.18 7 0.50 5 0.60 6 0.63 6	2.2 2.5 2.6 2.1 2.2	
19739121973917197391719739191973919	2144 9.3 2011 7.4 2101 8.6 0646 1.6 103015.6	39.033 39.536 39.011 39.416 39.388	118.157 118.536 118.142 118.581 118.481	0.41 7 0.12 5 0.12 6 0.58 6 0.22 4	2.0 2.0 1.8 2.2 1.7	
1973 9 22 1973 9 25 1973 9 27 1973 10 1 1973 10 4	062752.3 2320 2.6 203712.4 121720.9 055133.5	39.093 38.895 39.132 39.240 39.154	118.098 117.860 118.070 118.022 118.047	0.32 9 0.67 6 0.14 5 0.58 5 0.21 9	2.4 2.4 2.1 1.5	
1973 10 6 1973 10 7 1973 10 7 1973 10 1 1973 10 11 1973 10 12	232350.7 092858.8 171452.8 1853 4.8 1646 2.5	38.896 39.579 39.600 39.020 39.505	117.824 118.033 118.030 118.472 118.413	0.12 9	2.1 2.3 2.5 3.1	
1973 10 13 1973 10 13 1973 10 16 1973 10 16 1973 10 20	020343.5 224310.7 001412.5 231620.5 140736.2	39.584 39.616 39.740 38.893 38.873	118.062 118.012 117.860 117.835 118.187	0.27 6	2.4 2.2 2.2 2.4 2.4	
1973 10 26 1973 10 26 1973 10 30 1973 10 31 1973 11 1	043119.9 191159.1 043714.4 1330 4.1 145520.7	39.495 38.891 39.248 39.230 39.233	118.289 117.826 117.358 117.389 118.110	0.18 7 0.48 7 0.14 6 0.17 10 0.06 5	2.5 2.6	

YEAR	MO	DAY	HRMINSEC	LAT	LON	RMS NSTA	MAG	DEPTH	Q
1973 1973 1973 1973 1973	11 11 11	3	060824.0 090458.0 061459.2 202736.6 164542.5	39.496 39.059 38.810	117.618 118.106	0.36 6 0.29 6 0.17 5	2.2		
1973 1973 1973 1973 1973	11 11 11	12 13 14	153438.4 121133.2 002459.1 221329.7 132949.4	39.353 38.898 39.042	118.094	0.19 7	1.8 2.2 2.4		
1973 1973 1973 1973 1973 1973	11 11 11	17 17 21	203633.1 071541.5 1244 9.4 123425.6 002328.7	39.310 39.607 39.675	118.530	0.25 7 0.22 9 0.05 5 0.44 6	2.0 2.5		
1973 1973 1973 1973 1973	11 11 12	28 30 2	101751.5 002212.2 221232.1 1450 .4 230421.2	38.897 39.163 39.066	117.852 118.042 118.166	0 .29 6	2.4	•	
1973 1973 1973 1973 1973 1973	12 12 12	10 11 17	010823.4 231611.0 222551.0 064216.4 1726 8.7	39.694 39.100	117.999 118.046	0.08 7	2.4 2.1 3.0		
1973 1973 1973 1973 1973 1974	12 12 12	28 28 29	001431.4 002247.4 153525.8 041558.5 062231.7	38.913 39.469 39.024	117.806 118.053	0.41 6 0.37 7 0.15 7 0.24 9	1.9 2.4 2.3 3.0		
1974 1974 1974 1974 1974 1974	1 1	18	152311.7 232736.4 033045.0 232237.4 1958 1.0	38.901 39.373 38.925	117.811 117.905 117.769	0.27 6 0.26 8	2.0 2.6 2.1 2.5		•
1974 1974 1974 1974 1974	2 2 2	30 7 9 10 11	114740.5 061328.6 1555 1.8 0528 4.8 184522.4	39.076 39.019 39.057	117.782 118.102 118.087 118.177 117.970	0.24 5 0.31 8 0.37 8 0.23 8 0.32 5	2.2 2.2 2.6 2.6 2.0		
1974 1974 1974 1974 1974	2 2 2	11 12 13 17 20	184640.4 082020.6 1451 5.9 055314.1 035724.1	39.001	117.971 118.116 118.481 118.028 118.063	0.32 6 0.02 4 0.36 8 0.10 6 0.32 8	2.1 2.1 1.7 2.4		

YEAR	MO	DAY	HRMINSEC	LAT	LON	RMS NSTA	MAG	DEPTH	Q
1974	2	20	041338.2	39.585	118.054	0.24 8	2.5		
1974	2	20	1940 .8	39.591	118.025	0.26 6	2.2		
1974		23	160124.7	39.134	118.069	0.27 4			
1974		25	202038.5	39.397	118.099	0.27 8	2.4		
1974		18	190340.4	38.919	117.812	0.31 6	2.2		
1974	3	22	954 3.0	38.874	118.061	0.20 4	1.6	28.12	С
1974		22	135551.8	39.352	117.891	0.51 6		2.75	D
1974		22	135557.0	39.057	118.119	0.11 9	2.3		
1974	3	23	011319.9	39.607	118.052	0.21 7	2.0		
1974	3	23	173239.7	39.190	117.576	0.41 7		0.34	D
1974	2	26	254 2.1	39.648	118.139	0.03 6	3.4	5.62	n
1974		26 26	254 2.1	39.048	118.139	0.03 8	2.8	5.02	D
1974		26 26	232119.6	38.897	117.843	0.23 8	2.8		D
1974		26 27	232119.0	38.831	117.926		2.0	2.45	D
1974		30	111357.5	39.005	117.928	0.13 10 0.18 7	2.3	2.40	D
19/4	3	30	111357.5	39.005	110,221	0.18 /	2.3	-	
1974	4	2	232524.9	38.905	117.827	0.46 7	2.2		
1974		5	093219.4	39.086	118.052	0.34 6	_ • _		
1974	-	5	2136 2.7	39.112	118,000	0.09 6			
1974		7	174321.6	39.320	118.078	0.15 8	2.5	1.38	D
1974		15	233332.9	39.062	118.453	0.78 7	2.8		-
	-								
1974	4	23	07Ì533.9	38.952	118.178	0.17 9	2.4		
1974	5	2	194958.1	39.112	118.038	0.21 5	2.6		
1974	5	3	042040.3	39.054	118.171	0.40 11	3.2		
1974	5	3	083746.1	39.054	118.074	0.27 11			
1974	5	5	221014.4	39.081	118.176	0.24 10	2.1	1.46	D
1974	E	12	231853.8	39.358	118.130	0.16 7	2.4		
1974		13	122617.7	39.156	117.999		2.4	4 01	-
		13	192152.0	38.916		0.31 14		4.01	D
1974					117.842	0.28 7	2.4	0.40	-
1974		15 15	61132.9 61136.4	39.057	118.093	0.09 6	2.2	0.40	D
1974	2	12	61136.4	38.859	118.147	0.03 5	1.9	•	D
1974	5	15	082938.3	39.245	118.168	0.16 11	2.2		
1974	5	15	131046.2	39.268	118.057	0.19 6	2.8	0.34	D
1974	5	15	131049.5	39.078	118.145	0.34 6	2.4		
1974	5	16	010638.3		118.424	0.10 8	2.5		
1974		17	14 915.9		118.134	0.10 7	2.4	1.87	D
	_							•	
1974		19	0014 3.6	39.314	118.110	0.22 5			
1974		23	1148 8.7	40.290	118.551	0.26 9	2.8		
1974		23	232335.7	38.895	117.827	0.48 6	2.7		
1974		26	174539.3		117.981	0.16 6	2.4	9.02	D
1974	5	26	174544.0	39.134	118.045	0.25 10	1.8		
1974	6	5	042049.6	39.243	117.387	0.48 7	2.9		
1974		5	163223.8	38.882	117.898	0.39 6	2.7		
1974		15	134853.6	39.455	117.986	1. 1. 1. 1. 1.	2.1		
1974		21	41134.7		117.623	2.72 9		55.70	D
1974		22	232543.3	38.949	117.835	0.78 7	2.5	55170	D
17/4	o	24	232343.3	30. 747	TT1.000	0.70 /			

YEAR	MO	DAY	HRMINSEC	LAT	LON	RMS NSTA	MAG	DEPTH Q	
1974	6	25	235632.9	39.628	118.460				
1974	6	27	1811 .8	39.507	118.022	0.22 6	2.3		
1974		29	205334.3		118.021	0.45 7	2.0	n an an Anna an Anna an Anna Anna Anna	
1974		1	105029.9			0.32 7	1 9	1°	
1974			143241.5		118.498		2 1		
						0.39 /	2.1		
1974			195736.4						
1974		19							
1974		20			118.113				
1974		22			117.955				
1974	7	25	232720.5	38.825	118.001	1.07 6	2.5		
1974	8	1	020025.3	39.026	118.235	0.09 .9	2.9		
1974	8	1 .	101555.1	39.107	118.052	0.19 7	1.6		
1974	8	.9	162717.1	38.826		0.28 11	2.6	2.48 D	
1974		11							
1974	-	14	193135.1			0.37 10			
		1 A.		· · · •, ·	117.000			National Action of the second se	
1974	8	25	041551.6		118.373	0.676			
1974		26	103332.8	39.105	118.059	0.23 14	2.2	5.12 D	
1974	8	26	183331.0	39.079	118.140	0.18 13	2.6	5.20 D	
1974	.9	2	233217.6		118.127	0.25 16	2.4	5.20 D 2.23 D	
1974	9	13	101628.2	39.374	117.932	0.09 6	2.5	2.48 D	
1974	9	13	101633.8	39.077	118.030	0.04 6	2.1		
1974	9	13	142114.8	39.158	118.058	0.16 8	2.3	2.92 D	
1974	9	14	01812.4		117.958			1.27 D	
1974			173159.6		118.064				
1974		17	1732 .2		118.115	0.21 6		0.01 2	
	9	26	5 626.6		118.032	0.17 16	2.6	1.62 D	
1974		27	34053.8		118.160	0.10 8	2.1	8.59 D	
1974			175227.3		117.885	0.07 7	2.4	1.85 D 2.19 D	
1974			44034.2		117.841	0.13 11	2.3		
1974	10	11	141954.0	39.074	118.201	0.17 7	2.1	5.02 D	
			133640.6					56.29 D	
1974	10	24	133415.2	39.542	118.419	0.28 6	3.0		
1974	10	26	81840.6 5 619.3	38.862	118.072	3.12 10	2.0	0.30 D	
1974	10	27	5 619.3	39.119	118.033	0.15 10	2.2	1.53 D	
1974	10	31	224048.3	38,990	118,238	1.01 7	2.1	28.00 D	
1974	11	1-	15514.9	39.189	118.140	0.29 11	2.2	0.50 D	
1974	11	2	14647.6	38.802	117.969	0.28 6	1.9	3.31 D	
1974	11	6	95527.7	39.300	118.162	0.40 10		5.50 D	
1974	11	8	133624.4	39.715	118.260	0.10 8			
1974		12	201614.2	38.917	117.845	0.40 6	2.7		
1974	11	15	165421.5	39.616	118.075	0.25 9	2.6		
1974	11		83149.6			0.18 12		0.50 D	
1974			002425.4		117,905	0.24 5			
1974			194329.0			0.38 8			
1974			093327.0			0.16 7			
x 2 / 7	**	20	093327.0	32.331	110.104	0.10 1	2.2		

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YEAR	MO	DAY	HRMINSEC	LAT	LON	RMS NSTA	MAG	DEPTH	Q
1974 1974			101024.2 105743.8	39.605 39.596	118.071 117.981	0.36 9 0.31 9	2.2 1.9	•	
1974	11	29	003253.1	39.600	118.084	0.36 10	3.1		
1974		8	103316.5	39.074	118.168	0.18 10	2.4	16.78	в
1974			153013.0		118.113	0.19 11		16.24	C
19/4	12	8	123013.0	39.075	118,113	0.19 11	3.0	10.24	C
1974			132850.6	39.157	118.124	0.21 5	2.6		
1974			074442.2	39.248	118.446	0.10 5			
1974			2104 5.7	39.506	118.472	0.17 10	2.8		
1974			141549.4	38 .94 7	117.679	1.01 7	2.0		
1974	12	17	1731 4.7	39.234	118.021	0.36 5	2.5		
1974	12	17	1731 7.4	39.043	118.214	0.15 9	2.5	2.50	с
1974	12	17	201540.1	39.033	118.213	0.02 7	2.0		
1974	12	22	0024 2.6	39.290	118.077	0.74 5	2.6		
1974			1432 8.4	39.346	118.168	0.38 9	2.6	0.14	D
1974			62848.7	39.074	118.128	0.05 8	2.6	9.35	B
1974			171618.6	39.148	118.135	0.04 8	3.2	11.42	В
1975	1	2	090144.8	39.048	118.153	0.13 7	2.0		
1975	1	3	150820.3	39.172	117.659	0.24 10	2.4		
1975	1	3	194218.8	38.993	118,182	0.23 7	2.0		
1975	1	4	053812.4	3 9.7 30	117.991	0.27 9	2.7		
1975	1	4	164630.5	39.045	118.147	0.29 11	2.2	4.01	D
1975	1	5	04754.8	39.019	118.189	0.26 27	3.0	8.66	в
1975	1	9	031926.4	39.396	118.451	0.23 6	2.1		
1975	1	9	0749 5.4	39.044	118.056	0.25 6	2.2		
1975	ī		094725.0	40.260	118.475	0.28 8	2.2		
1975		16	165249.1	39.084	118.065	0.22 15	2.3	0.74	D
1975	1	18	213522.3	39.022	118.160	0.24 19	2.5	9.85	В
1975	1	20	1240 9.6	39.036	118.130	0.05 7	0.4	7.81	С
1975	1	21	459 4.9	39.017	118.154	0.26 13	1.9	10.19	С
1975	1	23	205319.1	39.025	118.117	0.20 12	2.1	6.72	С
1975	1	24	4 614.9	39.035	118.174	0.16 11	1.8	7.23	с
1975		29	135452.4	39.627	118.106	0.27 7	2.5		-
1975		30	005637.1	39.522	118.472	0.09 9	1.9		
1975		30	2315 4.6	39.172	117.401	0.33 8	2.1		
1975	2	4	7 515.2	39.031	118.169	0.20 16	2.3	14 12	~
1975	2	4	/ 515.2	39.031	110.109	0.20 10	2.5	14.13	C
1975	2	10	062639.4	39.349	118.140	0.16 8	1.9		
1975	2	10	114729.1	39.761	118.396	0.30 7			
1975		15	45733.0	39.376	118.479	0.09 5	1.9	22.74	D
1975		15	142737.5	39.084	118,108	0.12 10	2.1	9.41	č
1975		22	0 248.4	39.066	118,107	0.15 12	2.0	7.80	č
1975	2	22	140424.4	40,009	118.244	0.18 8	2.4		
1975		26	81011.1	39.381	118,111	0.12 14	2.3	12.82	С
		26 26	95432.8	39.017	118.661	0.12 14			C .
1975							2.8	12.95	
1975		27	094439.6	39.078	118.050	0.43 7	1.5	17 50	~
1975	3	2	34119.0	38.806	117.944	0.19 11	1.8	17.53	D

YEAR	MO	DAY	HRMINSEC	LAT	LON	RUS NSTA	MAG	DEPTH	Q
1975	3	4	153254.5	39.032	118.178	0.05 7	1.0		
1975	3	9	023421.3	39.083	118.021	0.20 5	1.0		
1975	3	9	035818.2		118.409		3.5		
1975	3	9	1049 2.9				1.0		
1975		10	130434.8			0.16 9	1.6		
1975	3	10	130434.0	39.004	110.213	0.10 9	1.0		
1975		14	032827.9				2.4		_
1975		15	15 125.1		118.128		2.2	0.09	D
1975		19	225548.5				2.5		
1975		19	225750.3	39.358	118.088	0.31 8			
1975	3	20	025611.8	39.250	118.205	0.36 7	2.1		
1975	3	21	103747.5	39.381	118.166	0.37 8	2.1		
1975	3	24	15 231.2	39.030	118.197	0.19 14	2.6	10.59	В
1975	3	28	235958.7	38.994	118.202	0.26 8	2.1		
1975		30	215531.9		118.091		1.5		
1975		13	83056.0	39.064	118.057	0.17 9	2.6	11.81	С
		15		57.004	110.057	0.17 9	2.0	11.01	C
1975	4	13	10 024.1	39.067	118.073	0.21 11	3.2	11.23	С
1975	4	16	916 5.9	39.045	118.105	0.40 12	2.9	4.63	С
1975	4	18	151218.6	39.306	118.148	0.15 7	1.8		
1975		19	065749.6		118.574		2.9		
1975			14524.8	39.032	118.173	0.22 10	2.3	4.02	С
									C
1975		22	71219.9	38.989	118.193		2.2	7.86	В
1975			73328.7		118.111	0.12 8	2.3	5.61	С
1975	4	27	211149.1	39.061	118.154	0.15 11	1.9	13.63	С
1975			204443.4		118.479		3.1		-
1975		21	24736.8				2.3	10.13	С
1 - A		2						10010	Č
1975	6	1	0239 1.4		118.235		2.0		
` 1975	6	3	190359.3		118.229	0.41 7	2.6		
1975	6	4	105641.8	39.254	118.549	0.33 7	1.4		
1975	6	4	173932.0	39.327	117.599	0.21 6	2.2	•	
1975	6	5	191232.5	38,913	117.834	0.38 7	2.4		
1975	6	7	103228.9	39.226	118.544	0.37 8	2.3		
1975		14	12 253.9	39.255		0.14 9	2.2	15.33	D
1975		15	2056 8.3		118.184		2.5		č
1975			51522.7			0.18 11		9.02	D
1975		24	0504 3.2				2.3	9.02	D
1710	Ŭ				110.097	0.31 0	2.3		
1975	- 6	24	724 4.6	39.083	118.038	0.12 9	2.4	16.67	D
1975	6	27	95042.0	39.386	118.073	0.10 12			С
1975		28	221212.1	1. 1. 4	118.146		2.3		C
1975		1	175311.0			0.20 14			Č ·
1975	7		113056.2			0.15 12	2.2		č
: 1 ¹¹									C
1975		10	2 126.9	39.253	118.041	0.16 12	2.2	14.49	С
1975	7	13	011635.0	39.386	117.632	0.00	4.0		
1975		13	013658.2	39.384	117.651	0.00	4.2		
1975		21	1739 2.0	39.037	118.165	0.15 15	2.5	10.64	С
1975		23	151226.7		118.126		2.9	0.92	D
1713	1	20	-J1220.1	57.502	110.120	U. 12 1J	2.7	0.72	Д

YEAR MO DAY	HRMINSEC	LAT	LON	RMS NSTA	MAG	DEPTH	Q
1975 7 23	151512.8	39.402	118.084	0.07 11	2.6	9.35	С
1975 7 29	162429.5			0.26 6	3.1		
1975 7 29	163234.4		118.040	0.28 9	2.7		
1975 8 15	052039.3			0.11 5	2	- A.	
				0.13 23	2.9	1.62	D
1975 8 29	231746.2	39.034	110.112	0.13 23	2.9	1.02	U
1 975 8 30	003443.7	39.402	118.051	0,36 6	3.9		. C.
1975 9 6	15 8 3.1	39.213	118.062	0.12 13	2.5	15.01	C
1975 9 7	620 3.9		118.098	0.20 14		13.15	C
1975 9 10	1247 4.7		118.015	0.11 11			Č
1975 9 10	2023 5.8		118.112	0.08 9			č
1973 9 10	2023 3.0	32.041	110,112	0.00 5	2.7		Ť
1975 9 10	2032 5.8	39.040	118.107	0.00	2.3	5	1.0
1975 9 12	092554.6		117.855		2.5		
1975 9 22	085139.1	39.640		0.22 6			
1975 9 22	141740.8			0.28 5			
	112059.2		118.218	0.07 10		11.30	C
1975 9 25	112059.2	20.330	110.210	0.07 10	2.5	11.50	C
1975 9 25	1225 2.3	39.104	118.144	0.08 12	2.6	12.54	C
1975 10 4	111920.5	39.816	117.862	0.28 5			
1975 10 6	235110.1			0.07 8	2.4	11.39	C
1975 10 22	020123.0			0.26 7		ch.	
1975 10 22	155937.3	39.393	118.108	0.24 7	2.8	(+)	
1773 10 22	133337.3	07.070	1101100	0.2.			
1975 11 9	171658.6	38.853	118.012	0.06 10	2.5	13.42	С
1975 11 13	13 937.1	39.152	118.086	0.06 11	3.0	15.82	С
1975 11 22	346 7.1	39.022	118.187	0.05 11	2.3	11.23	С
1975 12 2	05156.2	39.104		0.07 14	2.8	9.46	С
1975 12 5	1153 2.9		118.444	0.20 6	2.4	Ľ.	
1975 12 5	224029.7	39.105	118.070	0.10 12		11.09	С
1976 1 5	134649.3		118.447	0.19 11			D
1976 1 21	52330.3	39.416	117.713	0.33 11	2.5	9.00	
1976 1 31	4 746.7	39.376	118.453	0.09 13	3.3	11.80	C
1976 2 14	221023.2	39.135	118.065	0.15 9	1.8	10.50	1 - 1 - 1 - 1 - 1 - 1 - 1 - 1 - 1 - 1 -
					~ ~	0.75	-
1976 2 15	18 611.5		118.044				C
1976 2 20	184531.9						
1976 2 28	233039.5			0.05 9		12.07	C
1976 3 31	133330.9					e de la companya de l	1.1
1976 3 31	134513.9	39.290	118.465	0.28 10	3.1	5.60	ie in the
1076 4 0	125627 2	20 520	117.664	0.35 9	A 1	10.40	24 -
1976 4 2							
1976 4 2				0.44 10			
1976 4 3			118.486		1.4		
1976 4 21							C
1976 4 22	214433.9	39.405	118.078	0.45 10	.2,0	6.40	
1976 5 4	434 0.3	39.022	118.193	0.09 8	2.7	13.59	3 C 27
1976 5 5				0.05 6			
1976 5 6				0.00 5		2.77	С
			118.073	0.14 10			
1976 5 16	427 7.9	39.269	118.144	0.13 13	2.1	1.71	C

YEAR MO DAY	HRMINSEC	LAT	LON	RMS NSTA	MAG	DEPTH	Q
1976 5 21	12216.6	39,169	118.161	0.24 14	2.7	9.94	С
1976 5 21	115521.5			0.13 11		11.23	č
1976 5 21				0.11 7		9.24	D
1976 5 23				0.37 6		2.24	D
1976 5 23				0.18 15			~
19/0 2 23	204250.8	39.114	118.078	0.18 15	2.T	9.77	С
1976 6 1	213640.1	20,006	110 045	0.31 9	0 1		
	213040.1						
1976 6 7						10.00	_
1976 6 9	7 916.5		118.056			13.06	D
1976 6 14				0.45 16		11.20	
1976 6 22	223907.3	39.324	118.502	0.74 10	1.7	9.80	
1076 7 6	140000 0	20 160	110 100	0 01 0	0.0	0.40	~
1976 7 6	142838.3	39.158	118.120				D
1976 7 14			117.714		3.0	2.23	D
1976 8 6	245 3.9		118.070			5.87	D
1976 8 31	6 858.6		118.104				D
1976 9 4	180303.2	39.389	117.576	0.34 9	2.0	19.50	
1976 9 4	180353.5	39.383		0.34 13		14.00	
1976 9 7	2 740.2			0.84 14		10.23	D
1976 [•] 9 12	231930.6	39.115		- 0.13 6		8.83	C
1976 10 2	84055.3	39.110	118.068	0.21 11	3.9	10.46	С
1976 10 6	233146.9	39.091	118.093	0.12 8	2.6	8.42	С
1076 10 7	111046 0		110 000	0.07.10	• •	11 50	-
1976 10 7	111846.8			0.07 10		11.78	С
1976 10 26				0.48 13	2.1	12.10	
1976 10 28	205646.4			0.67 8	2.1	5.50	
1976 11 4	01138.8	38.865		0.22 16	2.8	17.12	C
1976 11 11	45909.3	39.435	117.765	0.32 12	2.2	18.20	
1976 11 14				0.21 12		17.40	
1976 11 17				21.59 16		12,80	
1976 11 18		39.358		0.09 8		18.10	
1976 11 27	13019.0		118.363	0.36 9		18.60	
1976 12 3	13 017.4	39.112	118.063	0.23 7	2.8	11.04	D
1000							
1976 12 3				0.16 7		9.89	D
1976 12 5	140301.4						
1976 12 27	83141.6		118.050	0.13 8		7.36	C
1976 12 28	61314.1	39.579	118,287	0.24 6	3.0	1.62	D
1976 12 31	726 8.0	39.723	118.342	1.82 4	3.2	1.95	D
1077	19	00 000					-
	17-2 0.0						D
1977 1 14	1331 0.2						С
1977 1 17	224856.4			0.25 9		5.40	
1977 1 28	95251.6	39.087	118.601	0.47 13	2.1		
1977 1 28	115030.2	39.103	118.584	0.45 12	1.5		
1077 1	104005 5		110	o 44			
1977 1 28	184305.7		118.585	0.44 13			
1977 1 29	11236.0		118.567	0.34 14		21.10	
1977 1 29	25401.6	39.109	118.583	0.15 8		13.60	
1977 2 8	194931.1		118.160		2.3		
1977 2 9	21224.8	39.401	118.443	0.06 9	0.7	11.00	

YEAF	MO	DAY	HRMINSEC	LAT	LON	RMS NSTA	MAG	DEPTH	Q
1977	2	9	72409.7	39.149	118.080	0.24 13	3.5	16.40	14
1977		9	83633.4	39.179	118.054	0.12 9	1.5	11.77	С
1977		18	101946.2	39.354	118.112	0.35 9	115	15.00	<u> </u>
1977		17	14215.1	39.213	118.091	0.14 8	1.6	11.83	C
1977		23	52427.2	39.128	118.158	0.12 8	1.6	6.39	D
19/1	3	23	52427.2	39.120	110,130	0.12 0	1.0	0.39	D
1977	3	25	152444.8	39.592	118.047	0.35 19	2.2	11.70	
1977	' 3	26	214640.6	39.033	118.405	0.06 8	2.1	15.07	C
1977		4	12533.0	39.019	118.374	0.18 11	2.3	11.52	С
1977		11	93244.9	39.450	118.146	0.30 11	1.9	20.20	
1977		12	90011.1	39.749	118.375	0.29 12	3.0		
13/1	-	12	JOOTI • 1	571145	1101010	0.25 12	0.0		
1977	4	15	10 717.2	39.171	118.108	0.19 10	1.7	13.69	С
1977		15	225559.4	39.202	118.143	0.13 10	1.7	10.35	С
1977		28	64448.1	40.192	118.297	0.43 12	2.9		
1977			1531 2.9	39.088	118.191	0.14 11		10.32	C
1977		10	6 524.7	39.095	118.073	0.18 15	2.2	7.95	č
1977	5	10	0 524.7	39.093	110.075	0.10 13	4.2		С. н
1977	5	10	85336.7	38.893	117.943	1.29 9	1.2	5.50	D.
1977	5	21	1 233.7	39.108	118.053	0.12 13	1.7	13.52	C
1977		28	50731.0	39.603	118.423	0.28 12	2.0		
1977		28	172717.7	39.280	118.021	0.46 11	2.0		
1977		25	170208.3	39.103	118.035	0.29 7	1.0	• ¹¹ •	
1711	Ŭ	2.5	170200.0	37.103	110,000	0.25	~••		
1977	7	22	215331.3	39.144	118.109	0.20 11	2.5	3.26	D
1977	⁷ 8	2	51456.6	38.859	118.288	0.12 11	1.4	14.47	В
1977			80622.2	39.332	118.453	0.18 6		;	
1977			112041.5	38.873	118.003	0.16 10	3.0	14.68	С
1977			212824.9	39.087	118.157	0.16 12	2.2	11.94	č
-2777	-	•	2202.00						-
1977	8	10	111412.6	39.310	117.429	0.28 8	1.8	13.50	
1977	/ 8	13	91052.6	39.066	118.105	0.09 7	0.9	11.46	С
1977	⁷ 8	17	60806.3	40.284	117.421	0.29 5	2.1		
1977		18	43326.9	39.123	118.071	0.10 7	0.7	13.07	С
1977		27	84304.9	39.276	118.529	0.31 11	2.2	10.50	-
-2777	-		0.00117	0012.0	1101010	0102 22	2.2		
1977	10	1	026 8.4	39.212	118.065	0.05 7	1.1	12.05	С
1977	10	7	81331.0	39.230	118.088	0.10 7	1.7	0.23	D
1977	10	10	84723.3	39.578	118.075	0.31 13	2.2		
1977	10	13	143435.8	39.470	118.505	0.36 12	1.7	15.80	
1977	/ 10	15	154745.8	39.144	118.129	0.13 9	1.7	13.20	С
1977			205110.3	38.962		0.45 11			
1977	/ 10	18	2212.9	39.392	118.093	0.22 15	2.7	9.50	
1977			235 0.8	38.860	11 7.9 81		1.9	13.94	
1977	/ 10	25	1010 4.6	38.854	118.004	0.13 8	1.2	10.49	C
1977	10	27	014 9.5	38.869	117.982	0.13 10	2.6	12.74	С
1075	, · · · ·	. 07	120000 1	20.064	117 070	0.00 5	2 0	15 00	
1977			132922.1	38.864		0.00 5			C
1977			18 616.6	38.866	117.986	0.15 10		13.77	C
1977			04035.8	38.867	117.984			13.71	C
1977			83021.4					13.79	C
1977	11	2	11 011.0	39.049	118,196	0.11 10	2.0	12.97	С

	YEAR	MO	DAY	HRMINSEC	LAT	LON	RMS NSTA	MAG	DEPTH	Q
	1977	11	8	105235.4	39.111	118.116	0.00 4	1 A	8.68	С
	1977		9	11 134.3	38.986	118.192	0.13 8		8.24	C
	1977			104752.7	38.861	118.001	0.13 0	1.7		c
	1977			85422.5						
						118.109				~
	1977	12	10	111652.5	38.859	117.993	0.09 7	L./	13./3	C
	1977			2152 8.6		118.128			11.40	D
	1977			34311.7	38.899	117.984	- · · ·		12.41	С
	1977			35258.0		117.989			10.62	С
				13430.5		118.076				
	1978	1	8	113101.3	39.659	118.255	0.36 11	3.3		
	1978		10	131336.1	38.909	117.988			13.30	С
	1978	1	12	152127.0	38.911	118.002	0.31 14	2.3	5.43	D
	1978	1	13	32910.8	39,382	117.577	0.29 9	2.2		
	1978	1,	13	33937.7	39.388	117.588	0.01 5	4.5	5.12	С
	1978	្រា	13	34608.9	39.379	117.562	0.32 10	2.3	ł	
	1978	1	13	40431.7	39.375	117.571	0.14 8	2.1	,	
	1978	1	13	42449.6	39.380	117.551	0.10 6			
	1978					117.576	0.17 9			
	1978			53011.2	39.377	117.584		,		
	1978		13	63322.9	39.367	117.577	0.33 6			
	1978		13			117.571			11.90	-
	1978		21	19 014.1	38.891	118.011	0.07 6			
	1978		24	224526.2	39.101	118.087	0.11 10		7.87	С
	1978	1	28	163115.5		118.002			9.76	С
	1978	. 1	28	181340.7	38.987	118.189	0.31 12	2.0		
	1978		30	191436.9		117.994	0.22 10			
	1978	2	15	92531.9	39.552	118.442	0.30 10			
	1978		18	155848.8	39.672	118.280	0.28 14	2.8	12.40	
	1978	2		215141.8		117.989	0.18 11		13.44	С
	1978	3	5	224619.1	38.867	117.993	0.00 4	4.4	13.07	С
	1978	3	5	225320.8	38.811	118.115	0.05 5	2.0	6.82	С
ţ	1978	3	5	23 123.0	38.830	118.064	0.05 4	2.6	8.87	С
	1978	3	6	22929.4	38.804	118.128	0.04 5	2.2	5.24	С
	1978	3	6	110605.6	39.380	117.556	0.26 7	2.0		
	1978	3	8	223016.8	38.856	117.993	0.19 9	1.8	10.10	
	1070	~	~	000010 5	20.050	117 000	o 11 o	1.0		_
	1978		.8	223616.5	38.856	117.993	0.11 9	1.9	12.55	С
	1978	3	9	83548.7	38.867	117.974	0.28 13	2.5	14.91	C
	1978		10	5 415.8	38.872	117.982	0.15 13	2.0	14.32	B
	1978		11	84245.5	38.862	117.990	0.19 11	2.0	14.47	С
	1978	3	12	224850.8	38.868	117.968	0.31 9	2.3	9.50	
	1978		15	6 954.7	38.855	117.995	0.12 7	1.9	11.82	C
	1978		19	12322.8	39.401	117.725	0.20 8		•	
	1978		07	210430.5	39.111	118.061	0.21 18	2.2	11.00	
	1978	4	24	14 621.5	38.863	117,969	0.03 5	1.5	16.48	C
	1978	5	9	33937.5	39.728	118.510	0.37 12	2.0		
					*					

YEAR MO DAY	HRMINSEC	LAT	LON	RMS NSTA	MAG	DEPTH Q
1978 5 13	6 529.5	39.386	117.993	0.20 10	2.7	9.65 Ď
1978 5 14	60529.7	39.386	118.019	0.23 16	2.7	
1978 5 15	113440.1	39.649	118.054	0.30 9	2.3	
1978 5 23	19 452.4	38.869	117.989	0.04 7	2.1	10.55 C
		39.744	118.062	0.40 10	2.3	10.55 C
1978 6 13	193517.1	39.744	110.002	0.40 10	2.3	
1978 6 29	144239.5	39.392	117.583	0.32 9		
1978 7 2	1411 2.4	38.857	117.977	0.03 6	0.9	1 4.81 B
1978 7 22	412 2.1	39.116	118.070	0.14 7	1.8	10 .57 D
1978 8 5	23319.1	3 9. 070	118.099	0.25 11	2.0	9.40 C
1978 8 6	102919.8	39.044	118.113	0.15 8	0.7	11.13 C
1978 8 10	544 4.0	39.238	118.064	0.11 8	1.4	3.43 C
1978 8 12	103028.2	39.412	117.705	0.21 11	1.8	18.30
1978 8 12	115751.9	39.428	118.119	0.30 11	2.9	5.00
1978 8 13	124945.2	39.376	118.087	0.09 8	2.0	5.64 D
1978 8 13	143932.1	39.159	118.067	0.16 9	2.1	12.16 C
1978 8 14	123515.1	39.201	118,148	0.01 5	1.0	2.57 C
1978 8 15	184852.1	39.168	118.066	0.22 9	1.9	9.78 D
1978 9 5	222850.5	39.041	118.169	0.12 10		5.43 C
1978 9 5	223152.0	39.032	118,149	0.56 10	2.3	
1978 9 6	1 531.2	39.028	118,123	0.14 7	1.5	10.45 D
1978 9 8	153240.7	39.126	118.070	0.04 9	1.2	13.10 C
1978 9 8	153930.0	39.120	118.057	0.08 6	0.9	12.32 C
	225046.6		117.687	0.20 7	2.3	0.61 D
1978 9 12		38.894				
1978 9 26	22658.9	39.369	117.525	0.24 16	3.2	10.90
1978 10 14	1 252.4	39.115	118.117	0.17 9	2.0	10:08 C
1978 10 24	1945 3.4	39.075	118.075	0.13 10	2.3	10.20 C
1978 11 7	235410.3	38.881	117.862	0.21 9	2.5	0.46 D
1978 12 1	19 652.0	38 .9 09	117.798	0.11 6	2.0	0.42 D
1978 12 15	19 1 9.2	38.882	117.859	0.27 9	2.4	1.38 D
1979 1 3	154638.0	38.9 00	117.549	0.24 12	2.8	15.17 C
1979 1 10	185712.7	39.363	118,086	0.21 12	2.8	8.67 C
1979 1 23	0 039.3	39.651	118.464	0.47 14	3.0	10.54 D
	93547.3		118.172	0.13 7	2.3	5.48 C
1979 3 6			117.997	0.14 9	2.7	10 .49 C
1979 3 15	16 041.0		118.139			8.72 D
1979 3 18	64523.3	39.109	118.094	0.13 7	•	8.94 D
1979 3 18	65913.5		118.086			7.74 C
1979 3 20	234842.4		117.871			0.58 D
1979 3 25			118.051			7.52 D
1979 4 7	43449.5		118.093			10.29 C
1979 5 01	114602.8	39.304	118.139	0.34 10	2.8	1. a
1979 5 01			118.139		2.2	
			118.090			6.23 D
1979 5 17	141234.1					
1979 5 22	05220.0		118.220			
1979 5 25	17 9 2.1	39.000	118,210	0.14 14	2.1	7.83 B

YEAR	MO	Day	HRMINSEC	LAT	LON	rms nsta	MAG	DEPTH	Q
1979	6	11	93254.3	39.103	118.093	0.18 9	1.6	12.07	С
1979		30	105852.2	39.328	118.058	0.21 12	2.8	10.29	č
1979		21	43149.9	39.603	118.095		2.5	13.66	č
1979		21	45431.3	39.597	118.080	0.28 8	1.6	15.03	D
1979		26	45717.9	39.595	118.113	0.26 7		11.47	D
1212		20	-3727.53	37.373	110.110	0.20 /	2•2	11.441	D
1979	7	26	51527.1	39.589	118,111	0.29 7	2.9	12.36	D
1979		26	85822.8	39.593	118.098	0.28 7		12.30	D
1979		26	92326.1	39.601		0.28 9	3.1	3.38	D
1979		26	103857.9	39.592	118.074	0.01 6	4.2	8.52	ĉ
1979		26	104811.2	39.618		0.43 8	2.6	11.88	Ď
		20	20102212	071010	1101100	00.000		11100	2
1979	7	26	1052 5.6	39.601	118.114	0.15 8	3.2	8.67	С
1979		1	811 5.7		118.196	0.01 6	2.1	11.31	C
1979		02	035847.6	40.193	117.777	0.36 8	1.9		•
1979		24	41027.5	38.890	118.003	0.15 7	2.4	10.98	С
1979		16	52053.9	40.310	117.626	1.75 10	2.2		-
1979	. 9	29	53751.0	39.354	118.112	0.09 9	2.7	2.90	C
1979	10	10	53729.4	39.086	118.084	0.09 6		9.36	С
1979			162311.0	39.101	118.111	0.08 7	2.4	4.45	C
1979			142025.8	39.120	118.605	0.08 10	1.8	15.10	C
1979			93557.3	39.111	118.591	0.15 12	2.2	18.15	В
		,							
1979	11	1	17 727.3	39.013	118.241	0.00 4	1.5	9.58	С
1979	11	3	15 4 2.4	39.634	118.392	0.14 8	2.8	13.09	D
1979	11	6	52543.2	39.084	118.334	0.14 11	3.0	15.27	С
1979	11	8	94332.0	39.930	117.799	0.63 6	2.3		
1980	2 1	18	213241.4	40.133	117.898	0.02 5	0.6	7.44	С
								1	
1980			143927.6	39.316	117.781	0.08 8		20.25	С
1980		24	85747.6	39.011	118.078	0.07 7	0.9	11.89	С
1980			53343.6	39.694	118.483	0.12 6	1.0	0.52	D
1980		25	111459.2	39.131	118.113	0.04 7	2.7	16.61	С
19 80	1	26	05239.3	39.571	118.179	0.07 7	0.2	12.99	C
1000					110 100				_
1980		26	05259.5	39.576		0.04 6		12.69	C
1980		27	55913.8	39.570	118.064	0.04 8	1.2	13.40	C
1980		29	55532.5	39.306	118.079	0.10 6	1.0	18.53	C
1980		29	231340.0	39.092	118.182	0.00 4	1.4	13.16	C
1980	2	4	4 023.7	39.344	118.091	0.03 7	1.6	11.40	C
1980	2	4	2027 2.3	39.061	118.060	0.07 8	1.0	14.83	с
1980		5	14117.3	39.151	118.036	0.06 6	0.4	19.19	c
1980	2		45150.2	39.131	118.030	0.07 9	2.5	18.10	C
1980	2	5		39.151	118.078		1.3		C
	2	9			118.132				
1980	· 2	7	71323.3	39.251	110.132	0.02 5	0.9	17.97	С
1980	2	10	45633.3	39.624	118.044	0.08 10	1.4	11.77	в
1980		10	20 146.1	39.310	117.321	0.12 7	1.5	8.64	D
1980		11	115253.0	39.676	118.011	0.00 4	0.1	15.83	č
1980		12	215820.7	39.709	118.023	0.14 12	2.8		B
1980		12	22 834.4	39.707		0.09 6	0.5	11.43	B
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YEAR	MO	DAY	HRMINSEC	LAT	LON	RMS N	STA	MAG	DEPIH	Q	
1980	2	14	102549.3	39.732	118.067	0.05	6	0.9	10.60	B	
1980		15	9 1 8.2	39.141	118.197	0.10	8	1.9	5.99	D	
1980		16	131239.0	39.542	118.202	0.02	6	0.7	5.46	C	
1980		17	111557.7	40.124	117.897	0.06	8	1.0	8.29	C	
1980		27	6 852.4	39.566	118.101	0.06	5	0.4	12.98	D	
1900	2	21	0 032.4	39.000	110.101	0.00	5	0.4	12, 90		
1 9 80	2	27	62636.5	39.319	118.104	0.04	6	1.0	20.21	C	
1980		27	195225.5	39.044	118.109	0.01	5	1.4	10.66	С	
1980		27	211941.4	39.590	118.111	0.06	6	0.4	12.79	С	
1980	2	28	123035.9	39.723	118.077	0.08	6	0.3	9.47	C	
1980	3	1	1612 4.7	39.513	118.055	0.00	4	0.7	10.15	C	
1900		-	1012 407	07.010	1101000	0.00	•	•••	20120	•	
1980	3	5	83621.8	39.051	118.103	0.03	7		14.16	С	
1980	3	13	22 358.1	39.706	118.024	0.04	7	1.3	10.27	В	
1980		15	31617.5	39.704	118.069	0.09	7	0.8	11.54	B	
1980		15	548 5.8	39.716	118.080	0.06	7	0.6		B	
1980		17	84231.7	39.309	118.107	0.01	6	••••	15.48	Ĉ	
2200	Ŭ	- /	0120207			••••	•			-	
1980	3	26	1154 1.7	39.380	118.134	0.10	7	1.0	16.76	C	
1980	3	26	175057.4	40.296	118.324	0.10	6	0.8	6.39	D	
1980	3	29	1955 3.6	39.228	118,121	0.09	7	1.0	22.65	С	
1980		30	11710.9	39.348	118.076	0.13	7	1.4	15.15	C	
1980		31	12316.1	40.166	118.494	0.18	6		0.98	D	
1980	4	8	1345 5.9	39.736	118.001	0.08	8	0.7	7.62	В	
1980	4	15	54452.6	39.729	118.072	0.03	6	0.8	9.08	В	
1980	4	15	152454.0	39.732	117.979	0.10	6	0.6	10.27	в	
1980	4	17	213336.0	39.940	117.749	0.20	6	0.3	0.08	С	
1 9 80	4	19	155547.9	40.262	117.768	0.11	6	0.8	7.62	D	
1000		20	10641 0	20, 061	117 072	0 12	7	0.2	0 17	C	
1980	4	29	13641.3	39.961	117.873	0.13			0.17		
1980	5	5	214624.3	40.146	118.318	0.22	6	0.7	3.02	D	
1980	5	6	224840.4	39.197	118.122	0.13	7	1.2	0.96	D	
1980	5	12	192713.4	39.311	118.142	0.14	8	1.0	11.48	С	
1 98 0	5	15	524 3.2	39.571	118.214	0.03	7	1.0	9.68	С	
1980	5	19	132710.1	39.857	118.121	0.11	7	0.3	8.40	С	
1980		19	18 6 0.3	39.150	118.071		12	3.2	8.42	č	
1980		22	223645.5	39.716	118.076	0.08	8	0.4	7.94	В	
1980		23	19 456.0	39.557	117.390		10	2.4	4.28	D	
		25 25		39.141	118.059	0.03	-6	1.4	0.03	Ċ	
1980	0	20	32846.2	37.141	110.009	0.04	0	1.4	0.03	C	
1980	6	30	104732.1	39.592	118.096	0.05	7	0.9	13.48	С	
1980	7	1	202752.0	39.617	118.115	0.22	9	1.5	10.34	В	
1980	7	2	92355.5	39.753	117.852	0.06	6	0.2	5.23	С	
1980	7	4	4 1 1.7	39.751	117.851	0.10	7	0.3	5.50	С	
1980	7	5	581.8	39.908	117.854	0.22	7	0.1	4.98	C	
	-	• -				o	~		~ ~ ~		
1980		11	05856.5	39.405	118.534	0.08	8	1.5	0.73	D	
1980		13	04539.3	39.931	117.774	0.04	7	0.3	7.20	B	
1 9 80	7		22 350.9	39.713	117.837	0.12	5	0.2	8.86	D	
1980	7		124038.0	39.892	117.676	0.02	6	0.6	10.52	С	
1 9 80	-7	28	132229.5	39.187	118.086	0.13	8	3.1		. •	

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	1000	~	~	014407 1		110 050	• • • •	-		A T A	~
	1980		2	214427.1	39.033	118.059	0.14	7		9.73	С
	1980		13	94111.5	39.462	118.496	0.08	8.	1.8	12.23	C
	1980	8	21	23 258.5	39.686	118.326	0.19	7	2.4	0.77	D
	1980	9	1	145430.6	39.287	118.096	0.04	6	2.4	7.76	C
	1980	9	3	152619.7	39.037	118.202	0.10	6	2.1	12.86	C
2	_,,,,	-	•	20202707	051001	2201202	0120	Ŭ	2	12100	Ŭ
	1980	٩	6	15 932.3	39.706	118.021	0.05	7	0.3	8.59	B
	1980					118.021					
		-			39.702		0.05	9	0.6	14.43	B
	1980		18	64032.3	40.413	117.754	0.14	9	1.5	2.91	D
	1980		26	6 750.2	39.471	118.133	0.06	9		11.84	Β.
	198 0	9	26	203827.0	39.668	118.094	0.04	7	1.2	10.29	В
	1980	9	27	1155 8.9	39.677	118.093	0.07	8	1.2	6.53	В
	1980	9		212145.4	39.672	118.097	0.06	7	1.5	10.15	B
	1980		27	2123 8.4	39.670	118.101	0.05	6	0.3	12.36	B
	1980		27	214110.5	39.672	118.097	0.04	8		10.24	В
	1980	9	27 🛛	214611.5	39.675	118.096	0.04	7	1.5	8.84	B
	1980	10	04	032042.4	39.038	118.123	0.14	7	1.9	12.00	
	1980	10	8	102523.7	39.640	117.532	0.19	5	0.7	0.65	D
	1980		8	1210 5.6	39.720	118.078		6	0.3	8.54	B
	1980			170200.0	39.071	118.162	0.16 1		1.9	11.00	2
					40.194	117.766					D
	1900	10	24	64054.8	40.194	11/./00	0.22	8	1.5	2.01	D
	1980			165537.9	39.032	118.184		6	1.6	12.00	
	1980	11	04	004303.9	39.2 10	118.106	0.18	7	1.0	12.00	
	1980	11	7	42333.5	39.766	118.167	0.02	5	1.5	10.69	С
	1980			2 536.0	39.769	118.178	0.14	5	1.3	9.66	D
	1980			185006.7	39.229	118.125	0.18	9	2.7		-
	1700	**	20	105000.7	574227	1101125	0110	2	2		
	1000		07	7 411 0	20 705	110 500	0.10		1.4	0.10	5
	1980			7 411.8		118.506	0.12	7		-	D
	1981			73658.1			0.05	6		8.82	В
	1981	1	24	211629.0	39.355	118.130	0.06	8	1.4	15.04	В
	1981	2	. 1	1748 4.4	40.257	117.704	0.17 1	LO	1.1	0.65	D
	1981	2	3	61753.4	39.685	118.046	0.05	9	1.4	11.23	B
	1981	3	18	75626.6	39.511	117.494	0.13	7	1.1	17.18	D
	1981		16	72125.6	39.592	118.092	the second se	9	2.8	11.46	B
			· · ·	131555.3		118.093			1.5	10.38	
	1981		17		39.588			8			B
	1981		25	4 950.9				7	1.6	6.90	B
	1981	8	1	21140.1	39.684	118.067	0.10	6	1.0	3.54	C
						an a					
	1981	8	22	184214.4	39.729	118.073	0.08	6	0.8	7.54	B
	1981		31	221644.6	39.563		0.03	5		14.32	C
	1981	9		165042.3		118.132	0.08	6	1.9	1.46	č
	1981			43925.1		the second s	0.06	5		0.44	D
		×.				118.105			1.5		
	1981	τÛ	12	13910.5	39.417	110.113	0.04	7	T.2	15.09	B
				1							_
	1981			141731.0		117.743		9	2.0	15.92	C
	1981	10	22	33154.1	39.312	118.128	0.07	8	1.8	9.99	C
	1981			22 552.0	39.496	118.075	0.02	6	1.3	7.80	C
	1981			14 025.0	39.419	117.748		9	1.8	16.15	C
	1981			173858.9	39.709	118.050	0.08	7	1.5	7.57	B
	T 20T	τU	50	T/202013	57.109	110.000		•	T • D	1.51	

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APPENDIX C. NOTES ON MAGNITUDE, RECURRENCE CURVES AND FOCAL MECHANISMS

In this section, terms that may be unclear to non-seismologists are explained in some detail. These terms include magnitude, recurrence curves and focal mechanisms.

<u>Magnitude</u>. For the period prior to 1969, magnitudes were taken from Slemmons et al (1965). As indicated by the authors, those magnitudes were taken from a number of sources, including the University of California Bulletin of the Seismographic Stations, earthquake lists from the California Institute of Technology, and amplitudes measured on recordings of the Wiechert seismograph at Reno, Nevada. For earthquakes prior to 1916, magnitudes were estimated from the extent of the felt area or from the maximum intensity at the epicenter. For the period 1970-1975, magnitudes were determined primarily from amplitudes measured on a Wood-Anderson seismograph at Reno, although University of California or US Coast and Geodetic Survey magnitudes were used for larger events. For 1976-1981 most of the magnitudes were determined from measurements of amplitude or coda duration on recordings of the Nevada seismic network.

In this report, the symbol M is used primarily in discussing earthquakes for the period prior to 1970, and ML is used to designate "local magnitude" for the period when magnitude could be determined from instruments in the Nevada region. In general, however, M or ML are both referenced to Richter's (1958) local magnitude scale up to magnitudes of about 6-6.5; for larger events magnitude M usually refers to teleseismic surface-wave magnitude, determined from the amplitude of surface waves with period of about 20 seconds.

<u>Recurrence</u> <u>Curves</u>. The formula most widely used for representing the frequency of occurrence of earthquakes as a function of magnitude is the following (Richter, 1958):

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$\log N = a - bM,$

where N = the number of events with magnitude equal to or greater than M, and a and b are constants. This equation has been applied to hundreds of earthquake sequences throughout the world, with magnitudes ranging from less than zero to more than 8. The constant b, or "b-value", represents the slope of the recurrence curve, and has been found to vary from about 0.5 to 1.5, with an average value of about 0.9. The b-value is related to the proportion of small earthquakes to large earthquakes, and appears to be relatively constant for a given earthquake sequence over a wide range of magnitude.

Relatively high b-values have been found to accompany swarmlike sequences of microfracturing in laboratory experiments on rock samples (Mogi, 1966; Scholz, 1968). Mogi showed that b increases with the degree of heterogeneity of the rock both in composition and in crack density, while Scholz concluded that b depended strongly on the state of stress in the specimen and only to a lesser extent on its physical properties. Scholz showed that b is inversely proportional to the effective stress, and that high b-values are observed for low-to-moderate stress. In a later study, Wyss (1973) concluded that b is directly proportional to the local pore pressure; the Denver earthquakes that were triggered by injection of waste fluids showed a strong correlation between b-value and pressure in the disposal well (Healy et al., 1968). High b-values have also been associated with aftershock sequences and low b-values with foreshock sequences (ie, Ryall and Kyall, 1981).

Values of b previously determined for this region include 0.79 for the Ventura-Winnemucca zone (Ryall et al, 1966), 0.81 for an earthquake sequence near Reno (Ryall et al, 1968), 0.79 for the 1968 Adel, Oregon swarm (Ryall and Savage, 1969), 0.83 for the 1965-1966 Caliente, Nevada sequence (Ryall and Savage, 1969), and 0.81 for the northern Nevada region (Ryall and Douglas, 1974). Richins' (1974) value of 1.00 for the 1973 Denio, Nevada swarm was higher than average values for the region as a whole, and was interpreted in terms of structural heterogeniety and high pore pressure.

Focal Mechanisms. The direction of fault movement in an earthquake can be inferred from various types of analysis of the seismic waves recorded by stations around the epicenter, at local, regional and teleseismic distances. The method used in this report is the "fault-plane" method of Byerly, which stems from the observation that the sense of first motion of P-waves (compression or dilatation) can be divided into a quadrant pattern, where the quadrants, say on an imaginary sphere around the focus of the earthquake ("focal sphere"), are separated by two nodal planes. One of these is the fault plane, and the other is a plane perpendicular to the fault plane, called the "auxiliary plane" (Figure Cl). The fault-plane method has an essential ambiguity in that it cannot distinguish between the fault and auxiliary planes, but this distinction can often be obtained from known geologic or structural parameters. In general, the fault plane may have any orientation, so observations made on the earth's surface must be projected onto the focal sphere by accounting for reflection and refraction of the waves as well as geometrical spreading, and then the three-dimensional pattern must be resolved using a stereographic projection, such as the equal-area projection used in this report.

An example is shown on Figure C2. The two lines represent the intersection, on a lower-hemisphere projection (ie, an imaginary hemisphere below the focus of the earthquake), of planes passing through the center of the sphere (the focus). The two planes are orthogonal, and separate quadrants in which incident P-waves appear as either compressions or dilatations. One plane has an almost E-W strike and dips steeply to the north; the other strikes NE-SW and dips SE. The sense of motion is either left-oblique on the E-W plane (ie, northern block moving W and down), or right-oblique on the NE plane (southeast block moving to the SW and down). The fault-plane solutions on Figure 29 of this report have the compression quadrants shaded, instead of showing individual observations by dots or circles, as on Figure C2. On Figure 29, solution E would be intrepreted as pure normal faulting, either with the northwest block moving down on a NE-SW striking, NW-dipping fault, or with the southeast block moving down on NE-SW striking, SE-dipping fault. Solutions J and I would be almost pure strike-slip motion, either right-lateral on a NE-SW striking vertical plane, or left-lateral on a NW-SE striking plane, dipping steeply to the NE for solution I and to the SW for J.



Figure Cl. Pair of force-couples at right angles, equivalent to maximum and minimum compressive stress (respectively called the P- and T-axes) at right angles to each other and oblique to the force couples. The dashed lines are the fault and auxiliary planes, which are indistinguishable on the basis of P-wave first motions. For this model, compressions would be observed in quadrants with outward-pointing arrows and dilatations in quadrants with inward-pointing arrows.



Figure C2. Example of a fault-plane solution, using an equal-area stereographic projection. Dots -- compressions; circles -- dilatations. Lines represent the intersection of the fault and auxiliary planes with the surface of a hemisphere below the earthquake focus.

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