DOE/ET/27010-2

MC COY AREA, NEVADA Geothermal Reservoir Assessment Case History Northern Basin and Range

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H. D. PILKINGTON August 1981

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TABLE OF CONTENTS

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AB	STRACT	·	•			•		•		•	•		•							·	1
IN'	TRODUCTI	ON		•		•	•	•	•	•			•		•		•		•		2
EXF	PLORATIO	N H	IST	ORY		•		•			•		•	•	•	•	•	•			4
	Geolog	ica	1 5	tud	ies	•	•	•		•		•	•	•		•	•	•	5 .	÷	4
	Geoche	mica	al	Stud	lies	5.		۲		•			•	•			•	2.0		•	6
	Geophy	sica	al :	Stud	lies	; .	÷			•				•		•		•		()•/	10
	Explor	atio	on (Dril	lir	ng				•						•				•	13

TABLE OF ILLUSTRATIONS

Figure	1.	Location map for the McCoy geothermal prospect	3
Figure	2.	Geologic map of the McCoy geothermal prospect, Nevada (after Adams, Moore and Struhsacker, 1980)	5
Figure	3	Geologic cross-sections through well 14-7 (top) and well 66-8 (bottom) after Pilkington, 1980.	7
Figure	4	Location map of drill holes used on geochemical study of drill cuttings.	9
Figure	5	Contour map of mercury in drill cuttings from 120-160 foot interval at McCoy (Pilkington, 1980)	. 11
Figure	6	Gravity profile with automatic interpretation for de (top) compared with geologic cross-section (bottom)	nsity 12
Figure	7	Resistivity at 5km depth from ID magnetotelluric inversion (Te mode) and location of survey lines and stations	14
Figure	8	MT section (Te mode, ID inversion) along line C-C' compared with geologic section (after Lange, 1980)	15
Figure	9	Survey location map of the McCoy prospect, after Wilt et al, 1980.	16
Figure	10	MT section (Te mode ID inversion) along line A-A" compared with EM data, after Lange, 1980	17

Table I

Chemical analyses of McCoy Mine well and Well 66-8 waters

Page

Page

8

ABSTRACT

The McCoy geothermal prospect is located at the junction of the Augusta Mountains, Clan Alpine Mountains and the New Pass Range. The prospect was discovered in 1977 and in 1978 was made a part of the Geothermal Reservoir Assessment Case Study Program of the Department of Energy under contract DE-AC 08-79 ET 27010.

Geothermal exploration done during 1980 included geological and geochemical studies. Geophysical work included interpretation of gravity data, a tensor MT survey and an EM-60 survey by the Lawrence Berkley Laboratory. Two intermediate depth exploration wells were completed in 1980.

A shallow low-temperature geothermal reservoir was encountered in the Triassic rocks. The analysis of the exploration continued in an attempt to determine the source area for the thermal fluids.

- 1 -

INTRODUCTION

The McCoy geothermal prospect was discovered in 1977 during reconnaissance coverage of Nevada. The prospect was identified by thermal gradient measurements of existing holes and hydrogeochemical analysis of well water from the McCoy Mine water well. The prospect is located approximately 72 kilometers northwest of Austin, Nevada (Figure 1) and can be reached by means of a graded road which leads from U. S. Highway 50.

The McCoy prospect is located at the confluence of the Augusta and Clan Alpine Mountains and the New Pass Range. The prospect straddles the Churchill and Lander County borders.

In 1978 AMAX submitted a proposal in response to the Department of Energy's RFP No. ET-78-R-08-003, Geothermal Reservoir Assessment Case Study, Northern Basin and Range and was awarded a contract providing partial funding for exploration at the property. Detailed results of the work funded through the DOE will be published by the DOE through the University of Utah Research Institute (UURI) under DOE contract DE-AC 08-79 ET 27010.

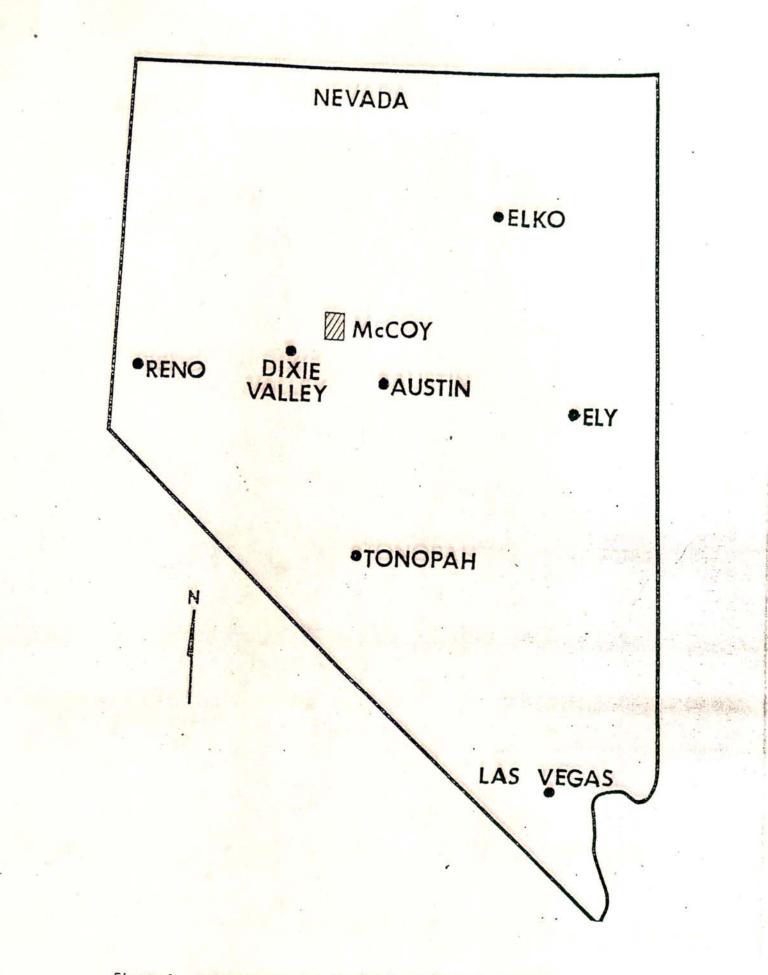


Figure 1. Location map for the McCoy geothermal project.

EXPLORATION HISTORY

The geothermal exploration partially funded under DOE contract DE ACO8-79 ET 27010 was an integrated approach including geological, geochemical and geophysical studies as well as exploration drilling. For the purpose of this report the exploration will be discussed under exploration methods rather than a chronological description.

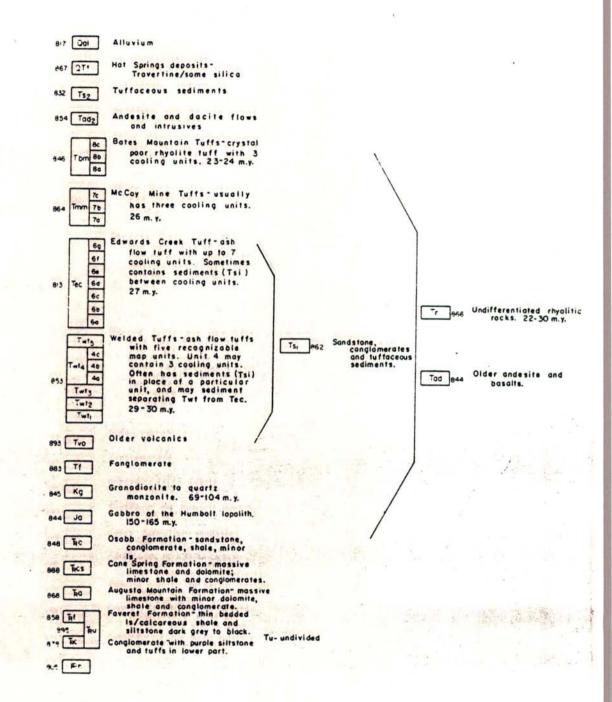
GEOLOGICAL STUDIES

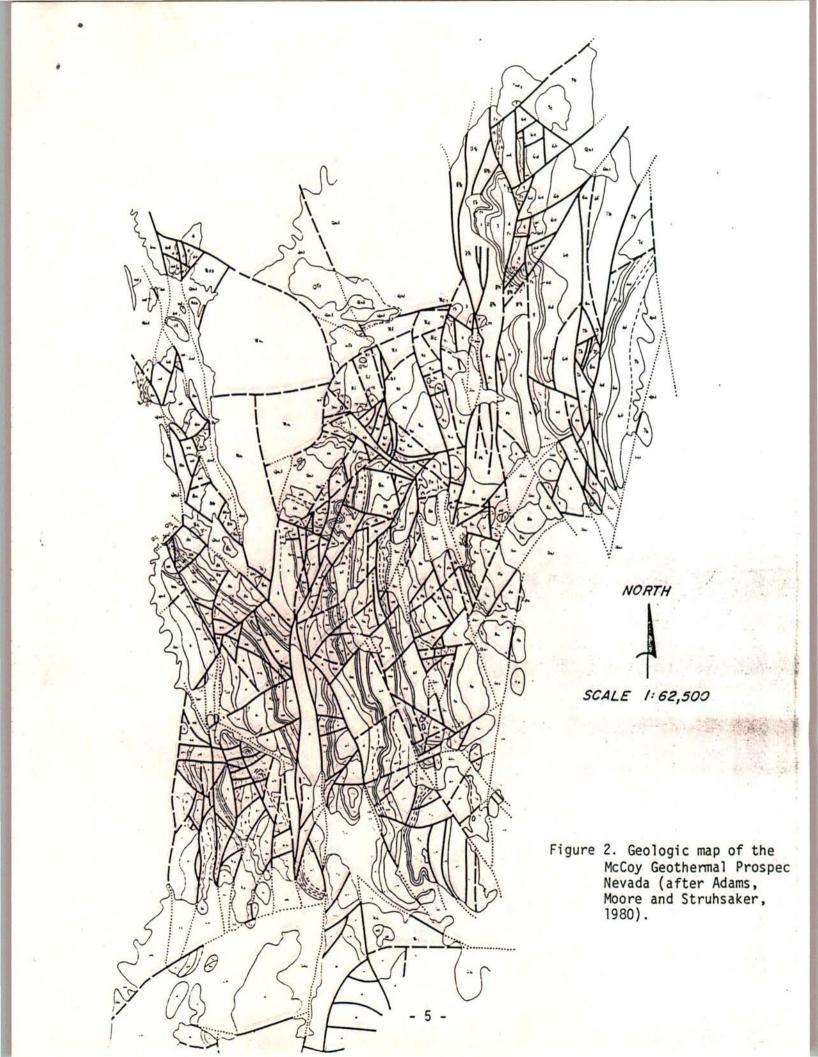
The McCoy geothermal prospect is located at the junction of the Augusta Mountains, the Clan Alpine Mountains and the New Pass range. The area is underlain by Tertiary volcanics and associated sediments, Triassic sediments and Permo-Pennsylvanian eugeosynclinal sediments as shown on the county maps (Stewart, J. H. and McKee, E. H., 1977 and Wilden, R. and Speed, R. L., 1974). The county geologic maps are at a scale of 1:250,000 and, therefore, do not show much detail.

In the late fall of 1979 Joe Moore of the University of Utah Research Institute began a geologic mapping program under DOE contract DE-AC 07 - 80 ID 12079. Joe Moore and Eric Struhsacker established the mapping units in the Tertiary volcanics and then Mike Adams of UURI did most of the field mapping in 1980. A detailed geologic map at a scale of 1:24,000 has been completed (Figure 2).

The oldest rocks mapped in the McCoy area are the cherts, volcanics, siltstones, sandstones and minor limestones of the Havallah sequence of Permo-Pennsylvanian age. The rocks were deformed, uplifted and deeply eroded prior to the deposition of the basal Triassic sediments which

- 4 -





consist of conglomerates, siltstones, sandstones and minor tuff. The basal member is overlain by several hundred feet of dominantly carbonate sediments of the Favret Formation, the Augusta Mountain Formation and the Cone Springs. Formation. The uppermost Triassic unit is the detrital sediments of the Osobb Formation.

The Tertiary volcanic rocks in the McCoy area (Fig. 2) range in age from about 36 m.y. to 15 m.y. (McKee and Stewart, 1971). Sedimentary rocks are found at various horizons within the volcanic units. A considerable thickness of sediments overlie the volcanic rocks to the northwest of the McCoy Mine area. West of the McCoy Mine over 2km² of fossil Quaternary travertine is exposed. The travertine lies unconformably upon the eroded Triassic rocks. The travertine dips gently westward and is only ślightly dissected.

Tectonically the area has had a long and complex history. The detailed mapping permitted the construction of detailed geologic cross section (Fig. 3). Reasonable approximations of the fault offsets could be made using the displacement of cooling units on Mike Adams geologic map.

Geochemical Studies

The hydrogeochemical analysis of a water sample from the McCoy Mine water well was one of the manifestations which attracted AMAX to the McCoy area. During the exploration drilling in 1980 water was encountered in well 66-8. Table I compares the chemical analyses from McCoy Mine well and exploration well 66-8.

- 6 -



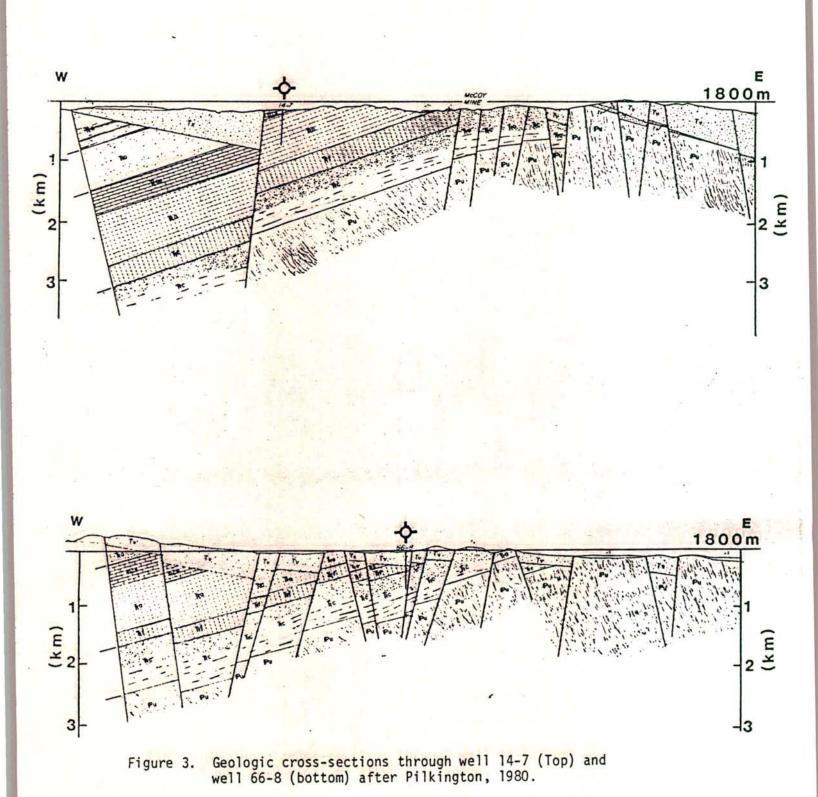


Table I. Chemical analyses of McCoy Mine Well and Well 66-8

	W 10981 McCoy H. W. Sec.9T20NR40E	W13453 Well 66-8, 1630' NWSE 8,T22NR40E	W13454 Well 66-8, 2050' NWSE 8,T22NR20E	W13456 Well 66-8,2410' NWSE 8, T22NR40E
Temp ^o C	39	+100	2 · · · ·	-
Flow (gpm)	-	- 25	-	-
pH	7.05	9.4	9.1	9.0
C1	22.0	38.0	31.0	31.0
F	4.4	5.6	3.0	4.1
S04	54.0	100.0	100.0	80.0
HCO3	611.6	144.0	142.0	204.0
C03	0.0	72.0	24.0	20.0
Si02	44.0	120.0	75.0	62.0
Na	260.0	160.0	98.0	110.0
K	15.0	21.0	14.0	14.0
Ca	43.0	6.6	9.6	6.0
Mg	9.0	2.6	16.0	18.0
Li	0.3	0.7	0.4	0.5
В	1.3	-		-
NH3	0.74	1		-
TDŠ	1065.3	670.0	513.0	550.0
Tq SiO ₂	98	148	120	112
Te Si02	66	122	94	83
T Na-K	174	242	250	239
T Na-K-Ca(1/3	3) 153	206	197	197
T Na-K-Ca(4/3				
T Na-K-Ca/				
Mg corr.	75	95	-	-

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0864-11 0864-46 864-470 0 864-48 0864-2

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MCCOY GEOTHERMAL

DRILL HOLE MAP

Figure 4. Location map of drill holes used in geochemical study of drill cuttings.

- 9 -

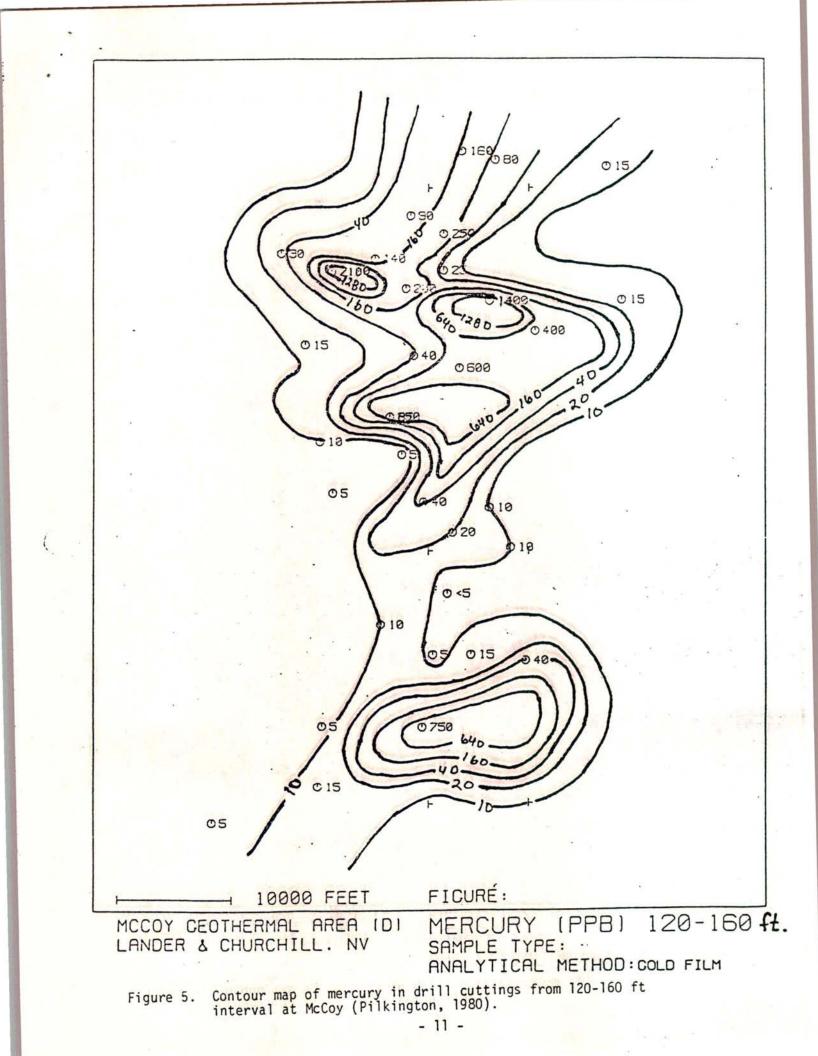
Chemically the waters from Well 66-8 and the McCoy Mine well are sodium bicarbonate waters with projected subsurface temperatures of $98-148^{\circ}C$. based upon the conductive quartz geothermometer and $153^{\circ}-206^{\circ}$ using the N-K-Ca geothermometer.

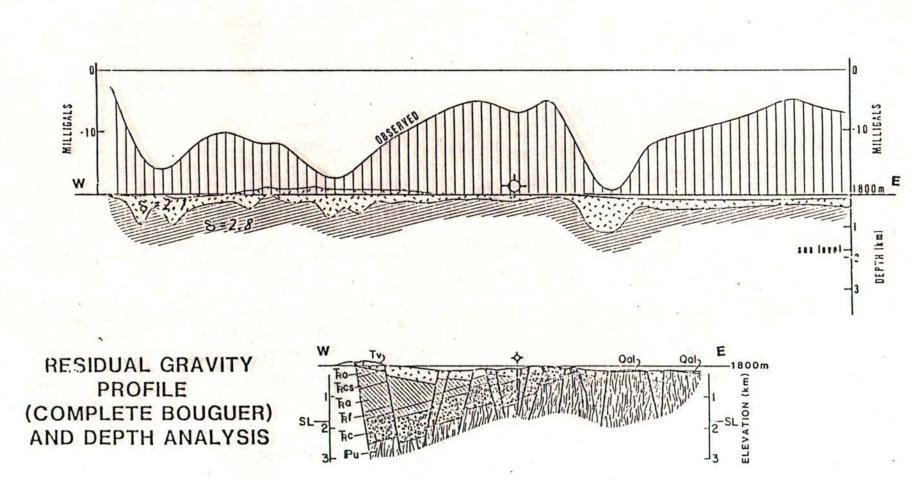
In 1979 Joe Moore of the University of Utah Research Institute proposed to undertake a geochemical study of the drill cuttings from McCoy. AMAX agreed to provide Joe with a split of our samples. UURI prepared composite samples for the intervals of 0-40, 40-80, 80-120 and 120-160 feet for each shallow thermal gradient hole (Figure 4). The geochemical study was done under DOE contract DE-AC 07-80 ID 12079. Multielement geochemical analyses using Inductively Coupled Plasma-Atomic Emission Spectroscopy (Moore 1980) were performed on each composite sample. Preliminary analysis of the data from Joe Moore indicates that Zn, As, Pb, F and Hg show some correlation with the thermal anomaly and also with the known areas of hydrothermal alteration and mineralization. Figure 5 shows the contour map for mercury for the interval 120-160 feet. The contour pattern appears to emphasize certain structural directions and the linear trends become more pronounced with depth (Pilkington, 1980).

Geophysical Studies

A gravity survey of 340 stations was conducted by AMAX and Microgeophysics in 1979 and Fred Berkman has been involved in the analysis of the data. Berkman prepared a residual gravity profile (Complete Bouguer) and depth analysis along an east-west line through well 66-8 (Figure 6) as reported by Lange, 1980. The gravity interpretation (top) is compared with the geologic cross-section by Pilkington (1980) on the bottom part of Figure 6.

- 10 -







 Gravity profile with automatic interpretation for densities (top) compared with geologic cross-section (bottom).

- 12 -

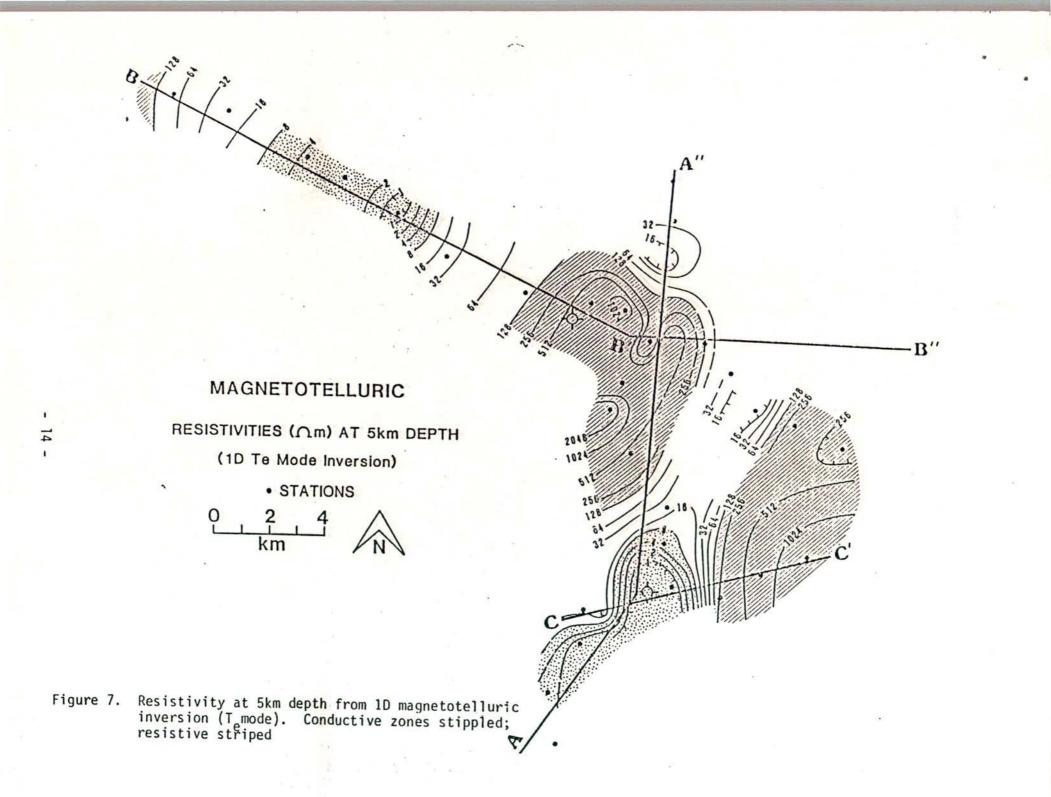
A tensor MT survey was run by Terraphysics in February 1980 (Lange, 1980) at the McCoy property (Fig. 7). The location of the lines along which the MT survey was done are shown on Figure 7. The resistivity as deduced by Lange (1980) at a depth of 5 km from the ID inversion of the MT (Te mode) is also shown on Figure 7. The MT section along line C-C' (Figure 8) illustrates the correlation with geology. Lange (1980) believes the MT sees a deep reservoir, (three or more kilometers) along line C-C' which is leaking fluids up the faults bounding the horst block east of well 66-8.

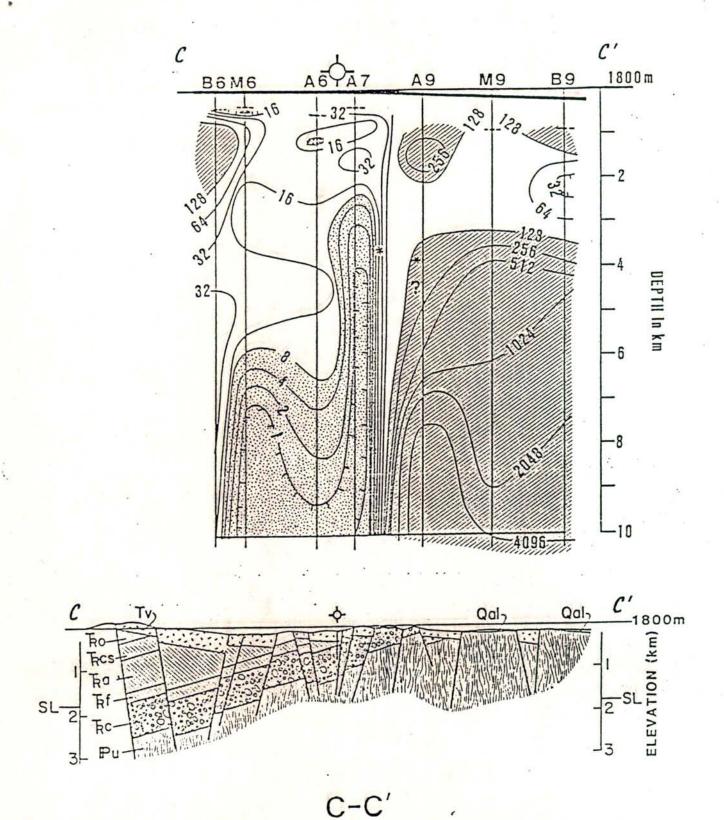
As a part of the Department of Energy's program to stimulate the development of geothermal resources Lawrence Berkley Laboratory (LBL) conducted a survey with the EM-60 frequency domain system over the McCoy prospect (Wilt, M. et al, 1980). The stations for the LBL survey are shown in Figure 9. The survey consisted of 19 frequency-domain electromagnetic soundings from three transmitter loops. A comparison between the EM data from LBL and the AMAX MT data is shown in Figure 10. Wilt et al (1980) conclude that the EM results agree well with the data gathered from well 66-8. A conductor was found at the approximate depths that boiling water was found in the well. The EM does give information on the shallow depths where MT does not give reliable results.

Exploration Drilling

Two intermediate depth exploration wells were completed in 1980. Well 66-8 located in the NWSE Sec. 8 T22N R40E had a TD of 765 meters (2510') and well 14-7 located in the SE NW Sec. 7 T23N R40E had a TD of 613m (2010').

- 13 -





MAGNETOTELLURIC 1-D INVERSION WITH GEOLOGIC PROFILE

Figure 8. MT section (Te mode, 1D inversion) along Line C, compared with geologic section - 15 -

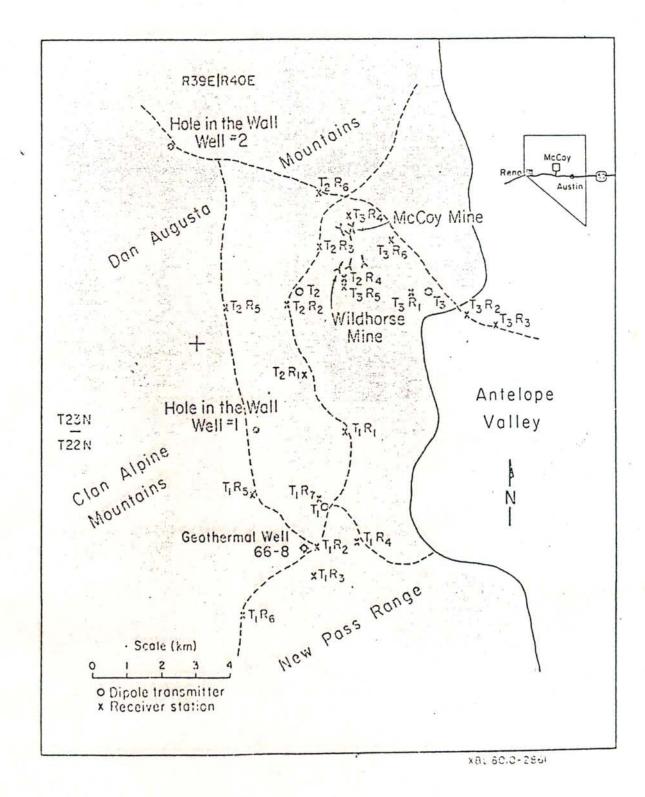
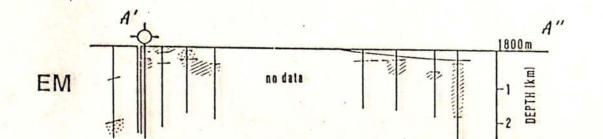
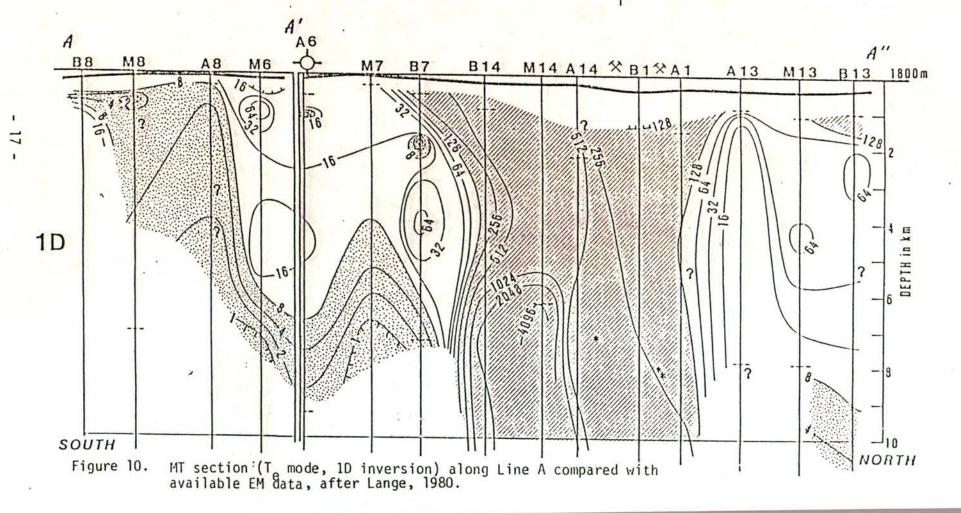


Figure 9. Survey location map of the McCoy prospect. after Wilt et al, 1980.

MAGNETOTELLURIC 1-D INVERSION WITH EM PROFILE





Both wells were drilled as unit wells for the McCoy Federal Geothermal Unit under a Plan of Operation approved by the U.S. Geological Survey. The generalized drilling plan for the holes was:

1. Move on. Rig up.

- 2. Drill 17-1/2" hole to +20 ft.
- 3. Run 20 ft of 13-3/8 conductor pipe, cement.
- 4. Drill 12-1/4 hole to +505 ft.
- 5. Run 500 ft of 8-5/8" casing, cement.
- 6. Install BOP equipment, test.
- 7. Drill 6-1/4-6-3/4 hole to TD.
- 8. Run electric logs.
- 9. Equip hole for flow test if appropriate
- Equip hole for temperature observation if no suitable production encountered - 3" Black iron pipe installed capped top & bottom filled/water.

Well 66-8 encountered a low temperature geothermal reservoir between 1630 feet and TD. Numerous hot water entries were recorded; with first entry at 1630 feet being the hottest at about 100°C, or just slightly above boiling for the elevation. The chemistry of the fluids was discussed in the section on geochemical studies. Lost circulation was encountered at several depths below the upper water zone. After completing the well as a thermal observation well it was discovered that drilling mud left in the hole had flowed out into the formation, and the 100°C water flows into the hole and out one of the lost circulation zones near bottom of hole. Therefore, no valid temperature gradients can be measured. It is proposed to back fill the annulus in 1981 to try and obtain reliable bottom hole temperature data. Well 14-7 encounterrd lost circulation in the Triassic carbonates at a depth of 409 feet. With considerable difficulty the 8-5/8" casing was set and cemented at a depth of 495 feet. A flow of warm water $\pm 50^{\circ}$ C was encountered just below the casing. Below 800 feet drill blind with no returns to TD. The well was completed as a temperature observation well, but the $\pm 50^{\circ}$ C water at 500 feet is going out into formation somewhere near the bottom of hole so that hole is isothermal below 500 feet.

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